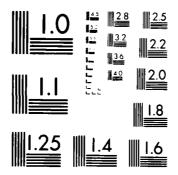
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(12)

Report No. cg-D-33-84

ATLAS OF THE BEAUFORT SEA

Ivan M. Lissauer L.E. Hachmeister B.J. Morson



FINAL REPORT October 1984

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Prepared for:

U. S. Department of Transportation United States Coast Guard Office of Research and Development Washington D.C. 20593



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SAMUEL F. POWEL, III

Technical Director

U.S. Coast Guard Research and Development Center Avery Point, Groton, Connecticut 06340



Technical Report Documentation Page

1. Report No. CG-D-33-84	2. Government Accession	No	3. Secipient's Catalo	g No.		
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		\$	October 1984			
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9. Performing Organization Name an	d Address		10. Work Unit No. (Ti	RAIS)		
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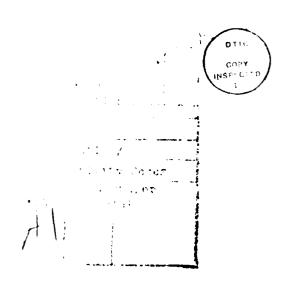
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SECTION A OCEANOGRAPHY

1.0 INTRODUCTION

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In the event of an oil spill in the non-existant or inadequate, or when the responsible party is unknown, the OSC shoreline, the Coast Guard predesignated are taken. In most cases, especially in this region, responsible parties will probably take the appropriate cleanup responsible party's actions are initiate cleanup action using Federal pollution funds. In either case, the OSC will be operating in a where cleanup actions are expensive and responsible for ensuring that timely and adequate containment and removal actions unique, remote, and hostile environment, action and the Coast Guard OSC's role will be to monitor these actions. (080) are conditions Coordinator Beaufort environmental American On-Scene the

In order to effectively respond to a spill on the Beaufort Sea, information

material for decision making in response needs. It can also help the OSC, who may not have special oceanographic actual field reconnaissance. It does, however, provide a means by which the reference quickly and is easily understood. It is important to emphasize that an Atlas, no replace essential. This Environmental Atlas has Alaska. This Atlas is designed so that information can be found become familiar with conditions in the area. also provides reference on the conditions that could affect oil seen compiled to provide the OSC with this information for the North Slope of information in a straightforward manner. and oil cleanup natter how complete, cannot the obtain a] so behavfor the necessary environmental can The Atlas training, user

The Atlas is divided into four sections: Oceanography, Meterology, Ice, and Climatology. It is designed to

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answer two questions the OSC responding to an Arctic oil spill might ask: (1) into what areas can the spill be expected to drift and how soon; and (2) what environmental conditions will personnel be facing at the spill cleanup site?

Current weather conditions and the specific geographical location of the spill source would form the Atlas entry points for calculating estimated trajectories. This information is located in the oceanography section.

questions regarding expected environmental conditions can be answered from information available from the Atlas' meterology, ice, and climatology sections. These sections contain comprehensive tables on mean, average, and frequency of occurrence of environmental conditions and, therefore, operational conditions that response personnel can expect to encounter. The Atlas has been prepared so that it may be updated and expanded as the need be updated.

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2.0. THE HIYSICAL ENVIRONMENT

T

The physical setting of the Beaufort Sea is such that the Beaufort Shelf is virspring and summer. Spring melting of the sea ice coincides with a massive The ice cover tends to insulate the to provide a source of dense brine in the winter and fresh meltwater in the influx of freshwater run-off from the land, both of which tend to stabilize thally completely covered by ice for all It is this open water portion of but two to three months of the year. underlying waters from both the atmosthe upper surface water, retain solar nest, and further enhance sea ice meltparticular year, the open water lead along the coastline may be as narrow as 10 km or as wide as 100-200 km and extend the entire length of the Beaufort the Beaufort to which this document is pheric temperature and wind fields and Depending on the wind field for a applicable. Sea.

Several studies (Aagaard, 1981; Matthews, 1981; Hachmeister and Vinelli, 1983) have concluded that the major driving mechanism for water exchange and

inner shelf, including the nearshore and lagoon systems, is derived from atmosnines both the direction and the intentention or removal of warm nearshore wind field in the western Beaufort is typically dominated year-round by easterly or northeasterly winds, whereas the erly and northeasterly winds in the In the western mean currents will be very similar to prevailing winds, primarily from the east. In the eastern Beaufort, winds are bimodal with prevailing winds from The wind field detersity of the longshore current, the remixing and horizontal exchange of laeastern Beaufort exhibits dominant eastsunner and westerly and southwesterly Beaufort this implies that year-round and from the east in the spring and Stonns tend to produce winds from the northwest in both regions and waters from the coast, the vertical goons, and (probably most importantly) the movement of ice and water on, off the west through the fall and winter, transport on the eastern Beaufort and along the nearshore region. winds in the winter. pheric forcing. sumer.

can result in atrong longshore currents, especially in the fall when a large expanse of open water exists along the Beaufort coast.

The shelf in the central and western Beaufort is relatively narrow with the shelfbreak typically occurring 80-90 km offshore. Lagoon/nearshore systems which characterize this region of the Beaufort coastline have been termed "open", that is, open to the wind-driven longshore transport and to onshore/ offshore exchange through numerous large openings in their offshore barrier island systems.

In the eastern Beaufort the shelf is slightly more narrow (approximately 40-60 km). The barrier island systems tent to be closer to the coastline, more extensive, and more closed to direct flowthrough by the longshore current, thus limiting the exchange of water between the longshore currents and the lagoons to a small number of openings in the barrier island system (limited exchange lagoons). In many of these limited exchange lagoons, the exchange

of water is restricted to one or two major entrances. These lagoons, which typically have very narrow entrances and poor exchange characteristics, exhibit highly localized current jets at the entrances in response to periodic tidal forcing and have been tenmel "pulsing" lagoons.

1

region of the shelf landward of the 40-m "inner" shelf; the region seaward of the That area of the inner shelf ity marked by the ice and/or surface temperatures and lower salinities than the water between 20 and 40 m, is referred to as "nearshore" after Truett baths) and large zonal variability in Following Asgaard (1981), the acteristics will be referred to as the 40-m isobath will be termed the "outer" landward of the 20-m isobath with activwaves, and which exhibits higher sunner tend to exhibit strong continuity in the the cross-shelf sense (crossing isoisobath which exhibits one set of char-Patterns of water movement on the shelf longshelf direction (paralleling isoshelf. baths). (1981).

GENERAL CIRCULATION

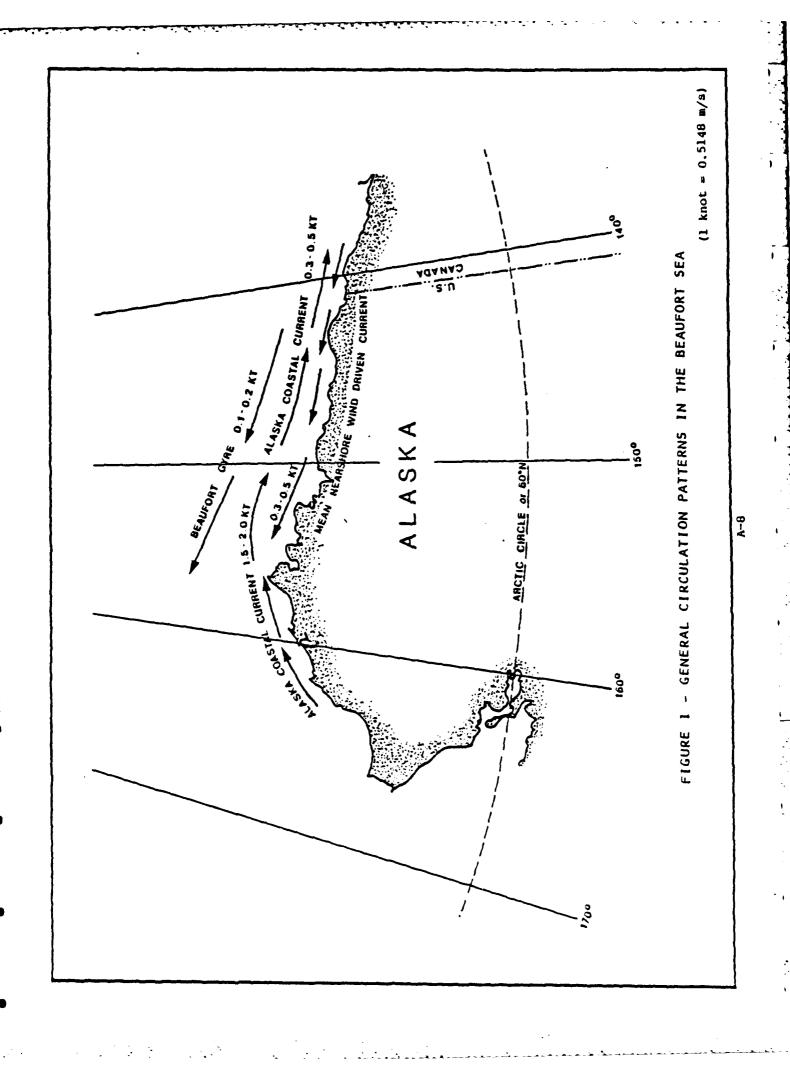
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3.6. GENERAL CIRCULATION

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Three current regimes are present along the Beaufort coastline (Figure 1). In deeper waters, the clockwise Beaufort Gyre moves waters from the Canadian Basin westward. Gyre velocities reach 5-10 cm/s north of the Alaskan coast (Agaaard, 1979). Inshore of the 40- to 50-in isobath currents are wind-driven and generally parallel to the coast (USDI, 1979). The Alaska Coastal Current enters the Beaufort Sea along

the Barrow Sea Valley. The current jet then follows the 200-m isobath through the Alaskan Beaufort Sea (Aggard, 1983; Thomas, 1983). Velocities are usually on the order of 15-25 cm/s to the east, but frequent reversals in direction result in a lower net eastward movement of 7 cm/s. In the nearshore region, mean currents are primarily wind driven and flow from 15-25 cm/s.



3.1 OIL SPILL TRANSPORT

Ourrent speeds in the Continental Shelf and Lagoon and Nearshore sections are conditions with 10 kt winds from the northeast and storm conditions with 30 In the may be rescaled for higher or lower wind appeals by the included formula. Ourrent directions typically follow local Intended to represent mean open water current on the specie or wind directions than those figures are primarily wind induced and Vary significantly for either different Dot case of the mean conditions, kt winds from the northwest. upedu which are displayed will arxl topography indicated. A simple vector technique is presented

below to quickly estimate the near surface transport of oil given the mean current field and the wind vector. This technique provides a rough approximation of resulting oil transport speed and direction which will be useful in estimating times and locations of spill landfalls.

Hean currents in the Beaufort Sea typically follow the bottom topography. Estimates of oil transport, however, will vary from the mean current field and instead follow a more wind driven pattern. This transport may be estimated by adding 3 percent of the wind speed to the mean currents as shown in Illustrations 1 and 2.

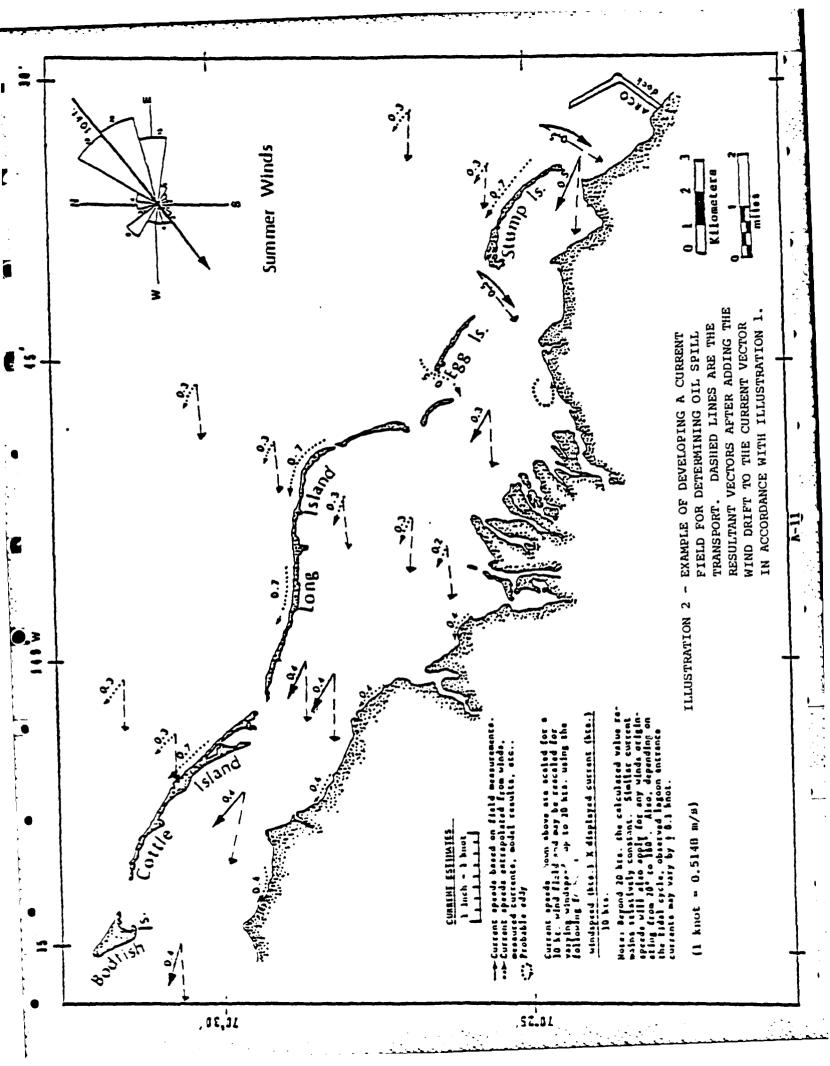
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The following technique may be used for determining the longshore and onshore transport components of the oil which was calculated above: λ_{lon-1}

(See next page for an example of this method.) overlays (such as mylar or acetate materials) to produce a complete transparent Note: This technique is most effectively used in the field by utilizing clear overview for the region of concern.

ILLUSTRATION 1 - TECHNIQUE FOR DETERMINING THE TRANSPORT OF OIL.

(1 knot = 0.5148 m/s)



CONTINENTAL SHELF CIRCULATION

4.0. CANTINENTAL SHELF CIRCULATION

4.1. Outer Shelf

associated with an eastward flowing core The most prominent hydrographic feature on the outer Beaufort Shelf is the summer subsurface temperature maximum typically observed at or seaward of the 40-m This temperature maximum is Alaskan coastal water can typically be seen as far east as 150 °W longitude The nearshore limit of this water typically follows the 40- to 50-m isobaths throughout the American sector of the Beaufort Sea during the sunner months and provides a good demarcation zone for separating the inner and outer shelf of water originating in the Bering Sea, referred to as Alaskan coastal water. where it mixes with local surface water. rejines.

The mean circulation pattern of subsurface waters on the outer shelf is predominantly eastward, paralleling the ischalfs. Aggaard (1981) and others have suggested that this mean flow, observed during both summer and winter months, is driven by the difference in

fic Oceans and extends across the entire Beaufort Shelf, possibly as far east as Baffin Bay. There is some evidence that sea level between the Atlantic and Pacisurface waters above this eastward flow may have mean westward motion, although no direct current measurements confirmperiods on the order of 3-10 days and may show reversals in the mean eastward flow. These reversals are in turn correlated with deep upwelling events on Currents in both the surface and subsurface waters can be affected by perialic meteorological forcing with typical ing this hypothesis have been made. the outer shelf.

4.2. Inner Shelf

Although current measurements on the inner shelf are extremely sparse due to the difficulty of maintaining moorings in the presence of sea ice, recent drifter data reported by Matthews (1981) and current meter measurements made by Aagaard (1981) give some indication of

both open water and ice-covered water novements. It has been generally agreed that water movement on the Beaufort This hypothesis is further supported by Mitthew's drifter data which suggests that the motion of all recovered drifters resulted from prevailing wind-driven currents, both for open water and underice releases. Drifter travel times and conputed current speeds were consistent with values of approximately 3-4 percent of the wind transport for the same periods with under-ice motion being signifiinner shelf is wind-driven. cantly less.

If these conditions can be extrapolated to the eastern Beaufort Sea, then current patterns on the inner eastern shelf would be expected to show a more even distribution of both easterly and westerly currents. As discussed previously, prevailing winds along the central and western Beaufort are from the ENE during all seasons. However, in the eastern Beaufort the distribution of winds is

more bimodal. At Barter Island, for example, the average winds are from the BNE to E for 35 percent of the time and from the WSW to W for 25 percent of the time (Seary and Hunter, 1971) with winds predominantly from the west during the winter and from the east during the winter and from the east during the water season (Brower et al., 1977). If the inner shelf waters in the eastern Beaufort follow the local wind patterns one could expect to observe mean current patterns to the east in the winter and to the west in the summer following local wind patterns.

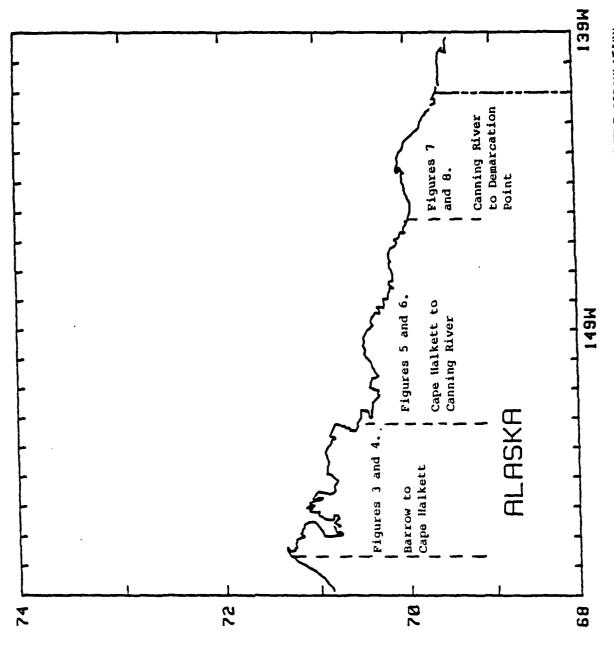
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In the western Beaufort, transverse circulation would be predominantly offshore in the upper layers and onshore in the lower layers due to the predominantly ENE winds. In the eastern Beaufort, the effect would be similar to the western Beaufort pattern during easterly wind events, and opposite (onshore flow in the upper waters and offshore in the lower layers) during westerly wind events. Seasonal variability would

therefore show a net offshore transport of surface waters in the summer and onshore transport in the winter in the eastern Beaufort.

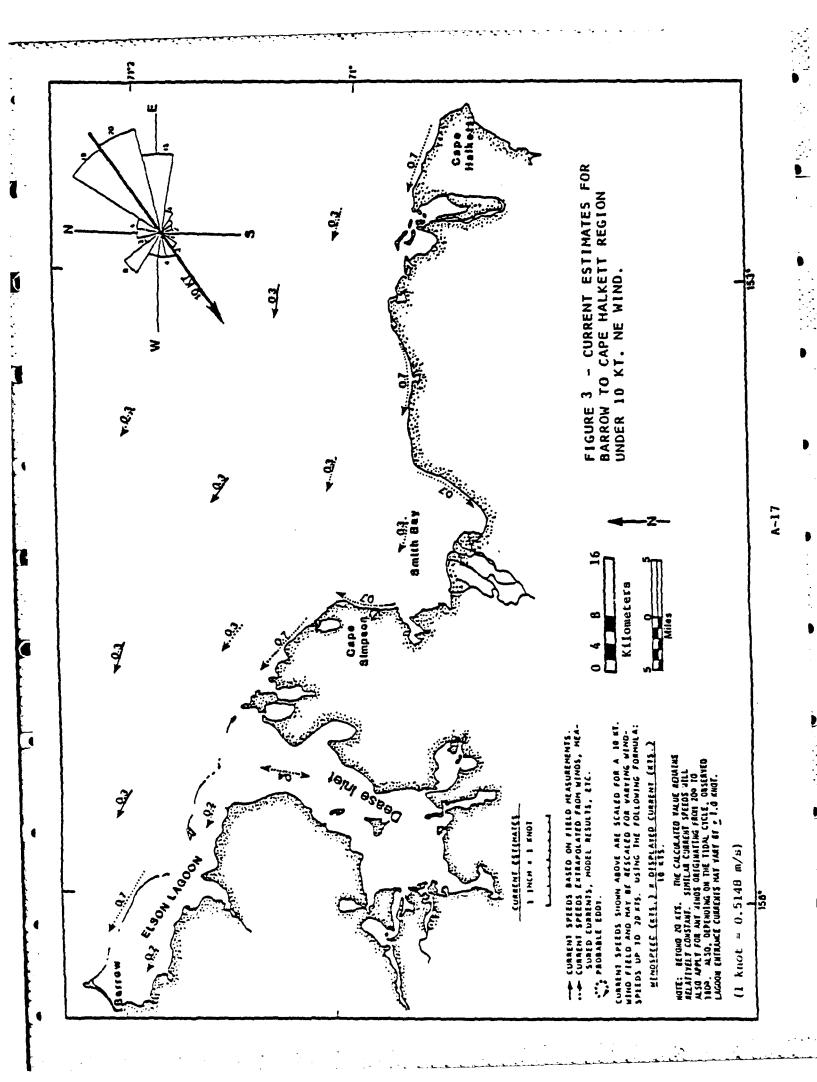
been demonstrated to be primarily wind with Longshore transport of water has also follow the local bathymetry at roughly three percent of the wind speed. In the the the driven with alongshore currents which nean mean Beaufort. In the eastern Beaufort mean Summer longshore transport is dominated weather patterns are bimodal in nature. by mean westward motion with reversals Passage of weather systems across are modified by implies easterly reversals occurring as longshore transport to the west this western Beaufort, wind conditions

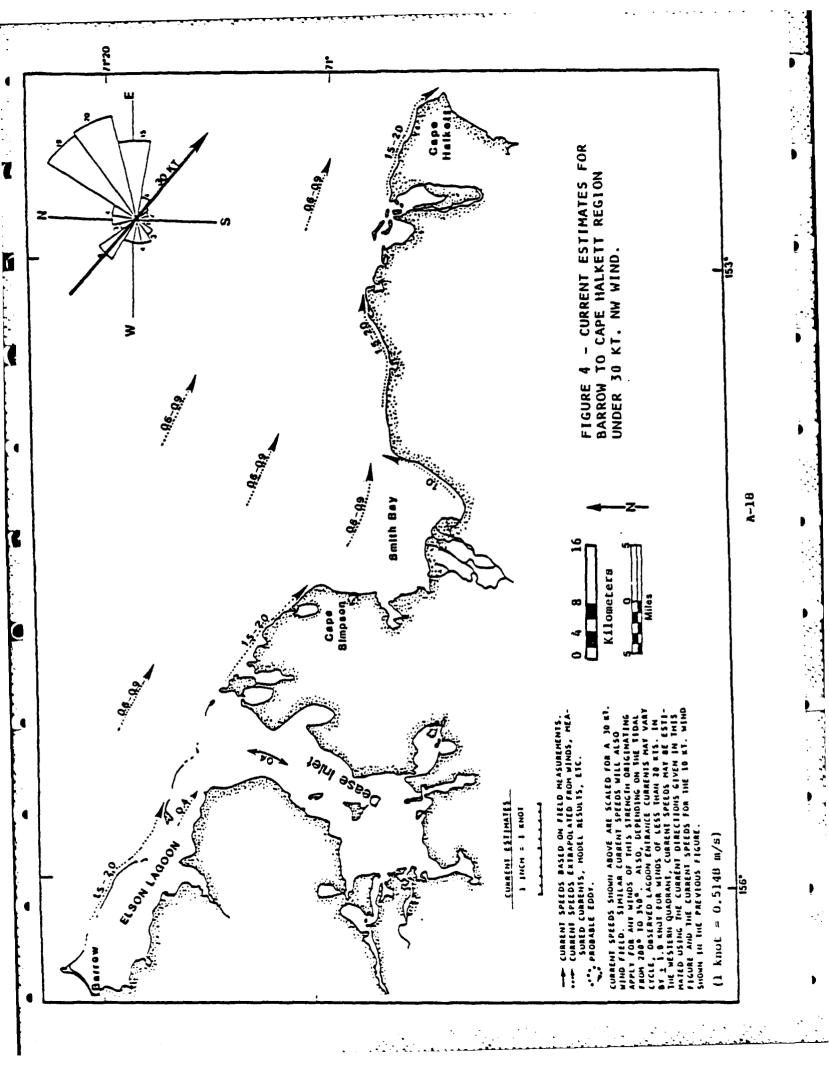
Fall and early winter water in this area with very high transport and wind-induced storm Westward transport also may be The maps following (Figures 2-8) show mean surface currents during 10 kt NE and 30 kt NW wind conditions. Wind roses were generated from Brower et al. (1977) using percent Prudhoe Bay. The pie-shaped wedges are labeled with the percentage frequency of occurrence of wind. Thus 15% of the time the wind is easterly, 20% of the of mean summer winds accompany periods to the east as weather systems show mean eastward fall with associated time it is ENE, etc. through the area. winds. ţ conditions occurrence novement easterly systems. eastward observed set-up

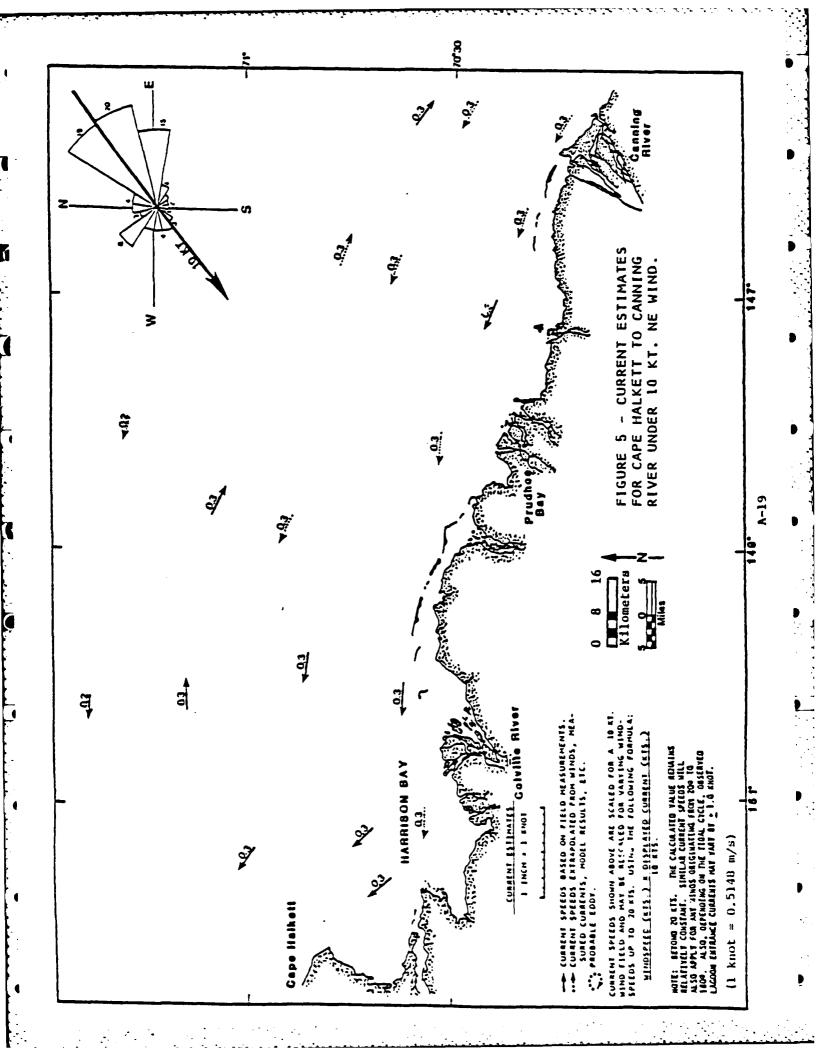


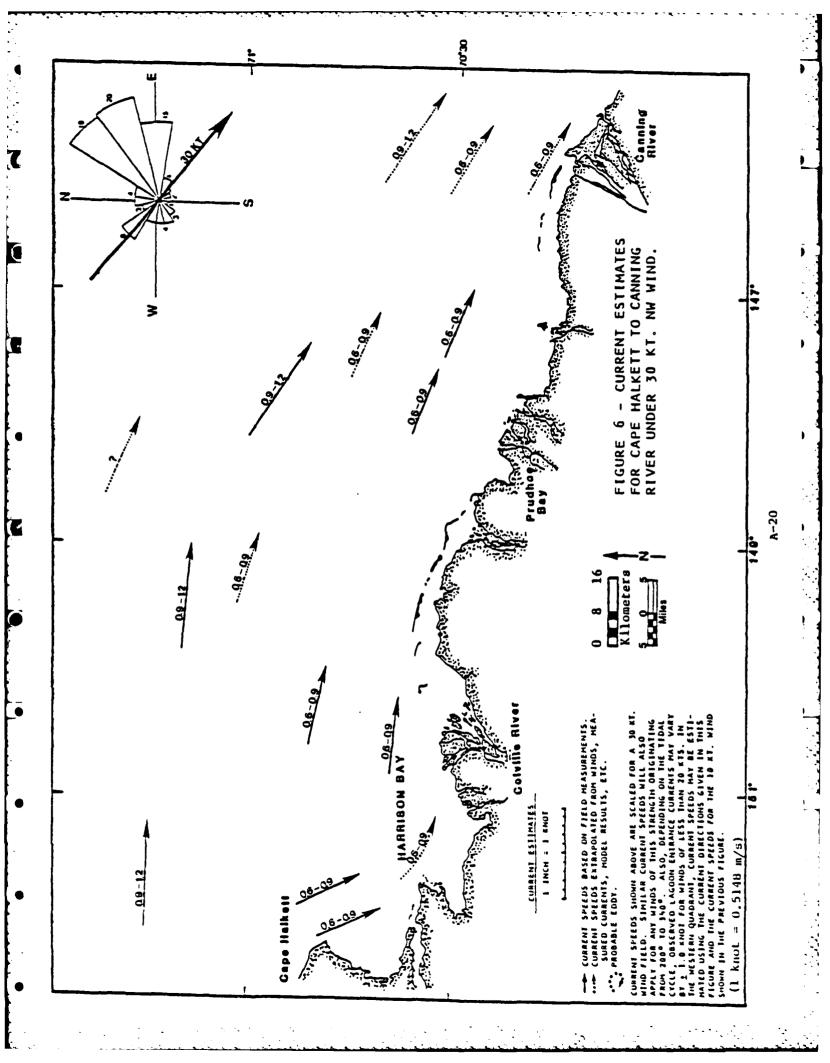
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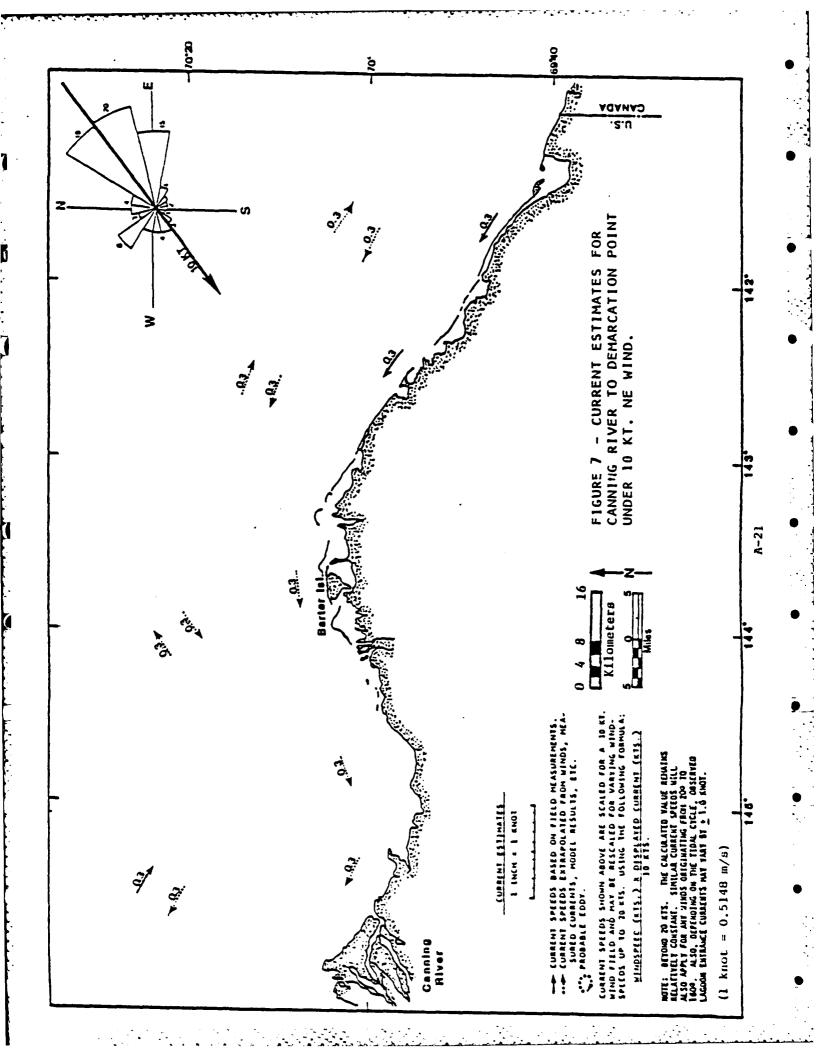
FIGURE 2 - LOCATIONS OF CURRENT ESTIMATE CHARTS FOR CONTINENTAL SHELF CIRCULATION.

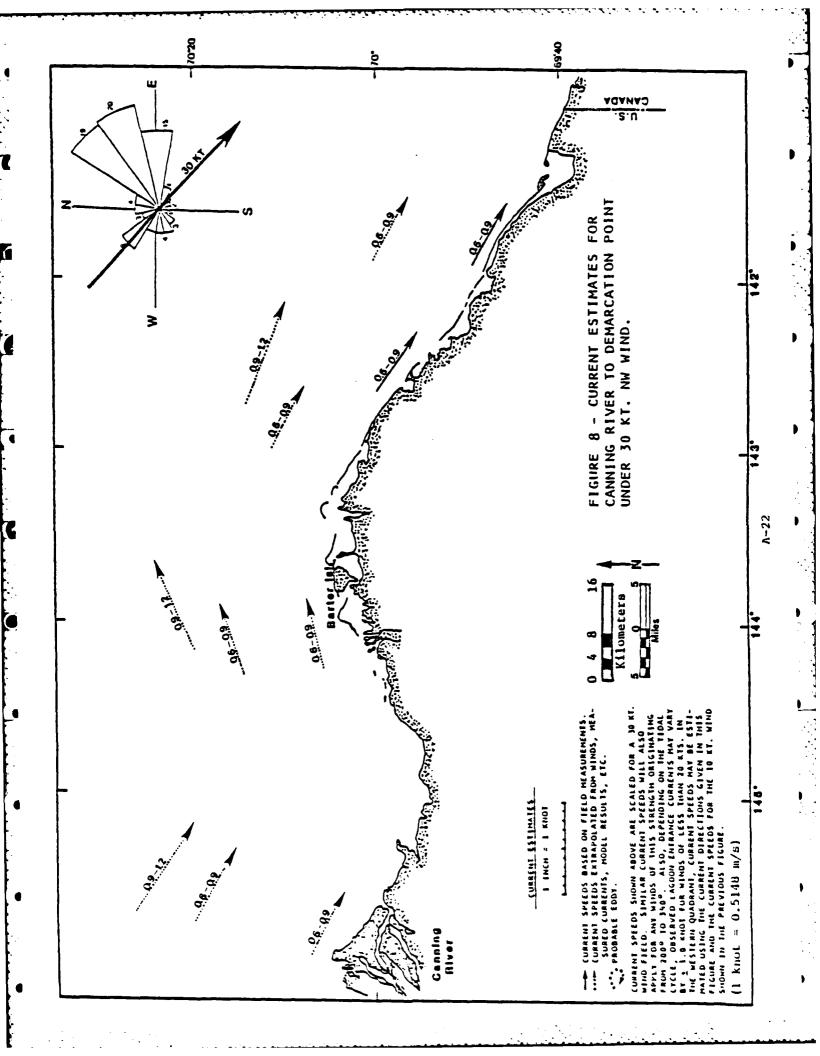












NEARSHORE CIRCULATION LAGOON &

5.0. IACCON AND NEARSHORE CIRCULATION

plished in the central and eastern Causiderable research has been accom-Beaufort nearshore and lagoon regions (Matthews, 1979; Hachmeister and Vinelli, 1983). Observed currents in the lagoons and nearshore appear to be spents approximately 3-4 percent of the win1 speed. Superimposed on these mean wind-driven currents are short-term lagoon/barrier island systems on the eastern and western Beaufort Shelf show tries and to differences in their surpredominantly wind-driven with current effects which are dominated by diurnal (M2) forcing. Circulation patterns and exchange properties of nearshore and many dissimilarities. Observed differences may be attributable both to differences in coastal and lagoon geomeeffects of storm passages and tidal rounding external physical environments. The basic lagoon types appearing on the Beaufort coastline are illustrated in Figure 9. The first type discussed in this report is the open lagoon, i.e. that which is open to longshore transport as well as to cross-shelf exchange between multiple large openings in the barrier islands.

The second lagoon type discussed is the pulsing lagoon, i.e. that which has one major entrance through the barrier islands. The pulsing lagoon is closed to longshore current throughput; exchange with the nearshore waters occurs primarily via tidal pumping of water through a single major entrance, with less exchange also occuring through shallow breaks in the barrier islands. One or more small rivers or streams typically empty into each type of lagoon providing a source of freshwater in early spring.

The third lagoon type is termed a limited exchange lagoon in that it has only limited longshore current throughput via several larger openings in the barrier island system. These lagoons may or may not exhibit pulsing effects due to tidal pumping. Each lagoon type is discussed below, using specific geographic examples of each.

5.1. Open Lagoons

The most extensively studied example of an open lagoon system is Simpson Lagoon. During summer easterly wind conditions, nearshore water enters the lagoon

through eastern and central openings in through the lagoon in a manner similar the barrier islands and is advected goon barrier island chain and the open Ē to the wind-driven longshore transport The multiple large openings in the lawestern end allow considerable flowchange is therefore largely due to advection of new water masses through the of the barrier island chain. lagoon rather than input/local mixing/ of the nearshore waters. seaward through output.

Tidal effects are superimposed upon the wind-driven component of the circulation and periodically modulate that component's effect. Depending on the strength of the wind-driven currents in a particular opening of the barrier island chain, the tide may only modulate the mean flow in the entrance, or it may actually reverse the flow during the opyosing cycle of the tidal current as has been observed in Angun Lagoon (Hachneister and Vinelli, 1983).

When both mean wind-driven and tidal conjuments of the circulation pattern

are acting simultaneously, the effect is order of 3-4 days for mean wind conditions of 10 kts. During these condirents then selectively introduce this lect near the eastern entrances of the a pulsing flow pattern in the lagoon with a mean flow from east to west through the lagoon. Matthews (1979) has estimated that the flowthrough occurs at 3-4 percent of the mean wind speed, indicating lagoon water turnover on the a net wind-driven transport of waters east to west through the lagoon is accompanied by some offshore transport of the warm fresh surface nearshore waters and replacement by cooler saltier offshore waters at depth. Tidal curnearshore water to the lagoon interior at each entrance on the successive flood tides. On ebb tides, the net westward flow is reduced and lagoon waters collagoon to form pools of warmer fresher tides, these pools of alternating cooler saltier nearshore water and warmer fresher lagoon water (formed by mixing of nearshore water from previous cycles and freshwater from river runoff) experience a net westward transport through On successive flood and the lagoon interior. tions, water.

a reversal in the win1-driven current direction from westof surface waters excurs which holds the Warm, less saline lagoon waters in the The current reversal is accompanies by the disappearance of the alternating nearshore and lagoon water masses (as nearshore and lagoon waters uniform warm intermaliate-salinity water winds also result in a rise in the mean sea level which may be as great as several meters in some cases (see Section 8.0). The reestablishment of mean During mean westerly winds (wind direcis diserved and a net onshore transport become identical) and observance of Persistent westerly winds and westward lagoon flow would buyin to produce an observed differential between lagoon and nearshore The rate of water transport imately the same both before and after the abrupt wind shift. During westerly winds, however, nearshore and lagoon waters appear to be identical while erly to easterly (current to the east) through the lagoon appears to be approxduring easterly winds, they exhibit differences in both temperature and The lagoons in the western tion to the east), in the lagoon. nearshore. salinity. Waters.

Beaufort from Barter Island to Pt. Barrow are almost all of the open type similar to Simpson Lagoon.

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5.2. Pulsing Lagoons

In addition, circulation in the and tidal effects on the exchange of Mean summer conditions are illustrated in this figure with winds easterly winds result in somewhat higher nearshore salinities and lower nearshore temperature for exchange with the lalagoon itself may depend on the lagoon Bay than for a lagoon with the longer Figure 10 illustrates the combined wind lagoon and nearshore water for a pulsing lagoon using Angun Lagoon and Pokok Bay predominantly from the east. In a manner similar to the open lagoon case, geometry with greater mixing observed for a lagoon with the geometry of Pokok more narrow shape of Angun Lagoon. as examples. goon.

During westerly winds, the fresher warmer nearshore water is maintained on the coast and advected eastward in the long-shore current as in the open lagoon case. However, when easterly winds are reestablished the nearshore waters cool

Andrews and the second second

(1 knot = 0.5149 m/s)

as warn water is driven offshore and the pulsing effects of cool-water-in/mixing/warn-water-out is observed. Nearshore current reversals have been observed to occur in as little as several hours (Hachmeister and Vinelli, 1983). Because there is no net flowthrough of the waters entering the pulsing lagoons, the interior of Angun Lagoon and Pokok Bay do not experience the alternating patterns of nearshore and lagoon water observed in Simpson Lagoon.

5.3. Limited Exchange Lagoons

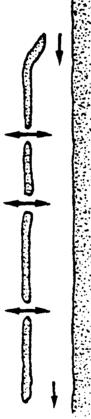
terned the limited exchange lagoon since only limited throughput of The final lagoon type to be discussed nearshore waters is possible. An example of this type is Beaufort Lagoon (east of Arryin Lagoon), actually composed of several small interconnected narrow lagoons with an extensive offshore barrier island chain. The barrier island chain (Icy Reef) has a relatively small number of openings distributed along its length, but several of the openings are quite wide, allowing the possibility of net wind-driven flowthrough of the waters from the longshore peen

followed by (moving eastward) Egaksrak Lagoon, Siku Lagoon, Pingokraluk Lagoon, entrances include Nuvagapak entrance to the west (which actually consists of two Siku entrance, and Other eastern Beaufort Sea lagoons with Beaufort Lagoon is Nuvagapak Lagoon, the main entrance to Damarcation Bay. one narrow and quite deep, the other wider but relatively shallow), configuration include Oruktalik, and finally Demarcation Bay. At the far western Jago and Tapkaurak Lagoons. Egaksrak entrance, openings: current. this

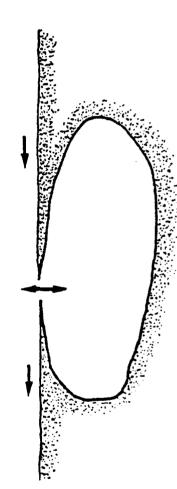
Flowthrough, hence purely advective exchange, is expected to be considerably less than observed for open lagoons. However, considerably more advective exchange is anticipated than for the pulsing lagoons. These limited exchange lagoons encompass over 75 percent of the coastline of the Beaufort Sea east of Barter Island.

The following maps (Figures 11-29) illustrate lagoon and nearshore mean surface currents along the Beaufort Sea coast during a 10 kt NE wind and a 30 kt NW wind.

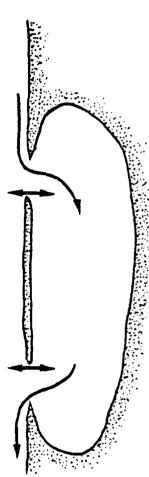
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TYPE 1. OPEN LAGOON (OPEN TO LONGSHORE CURRENT)

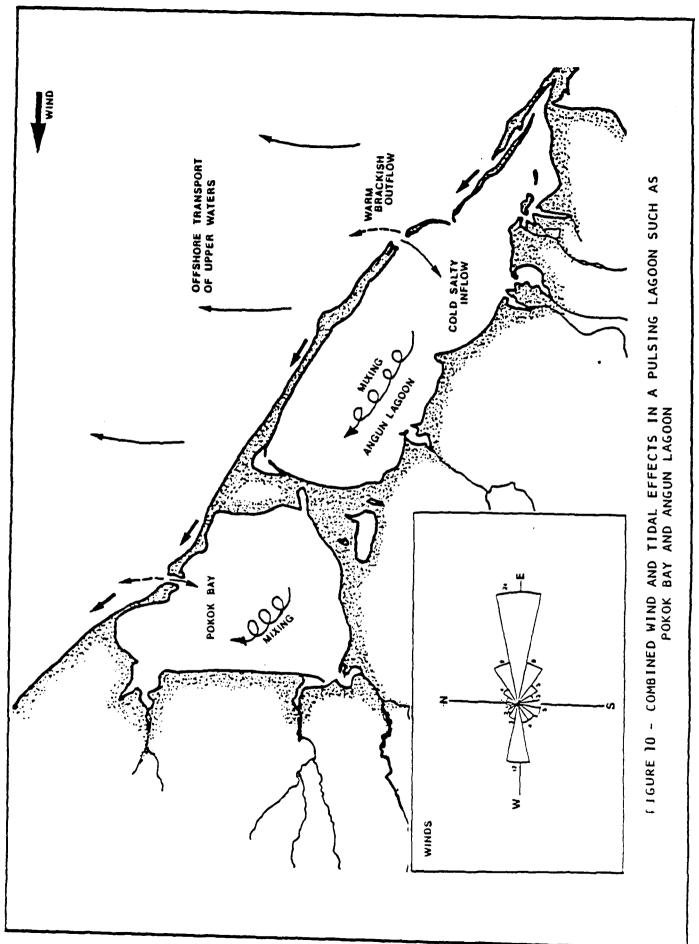


TYPE 2. PULSING LAGOON (CLOSED TO LONGSHORE CURRENT)



TYPE 3. LIMITED EXCHANGE LAGOON (LIMITED LONGSHORE CURRENT EXCHANGE)

OPEN, PULSING, AND LIMITED EXCHANGE. FIGURE 9 - BASIC LAGOON TYPES:



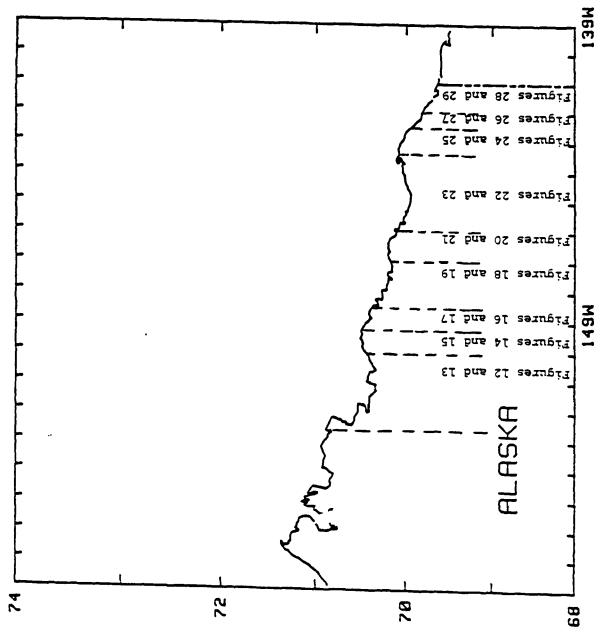
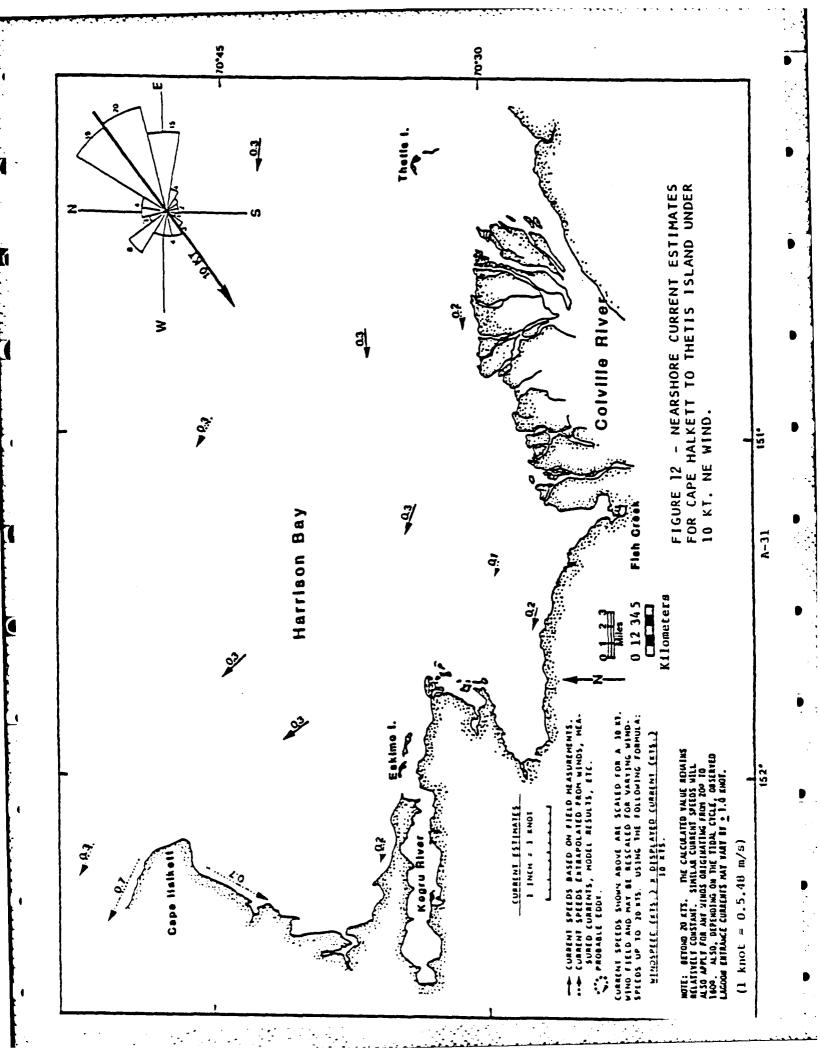
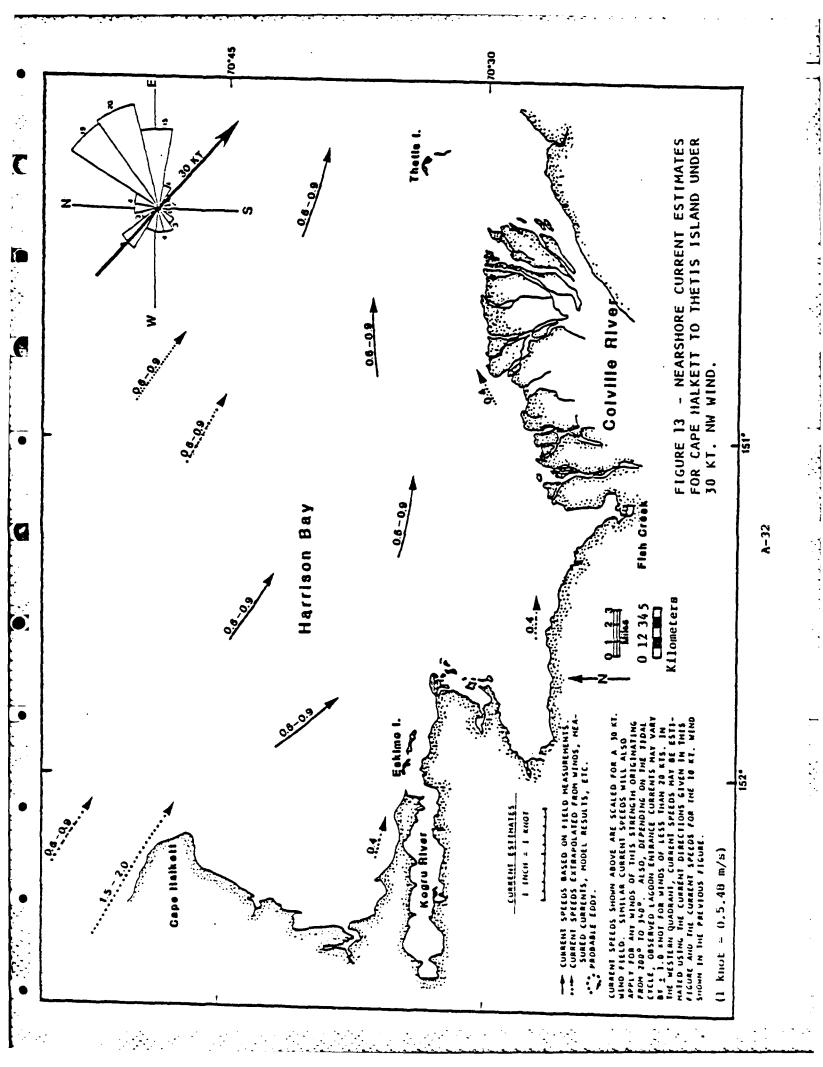
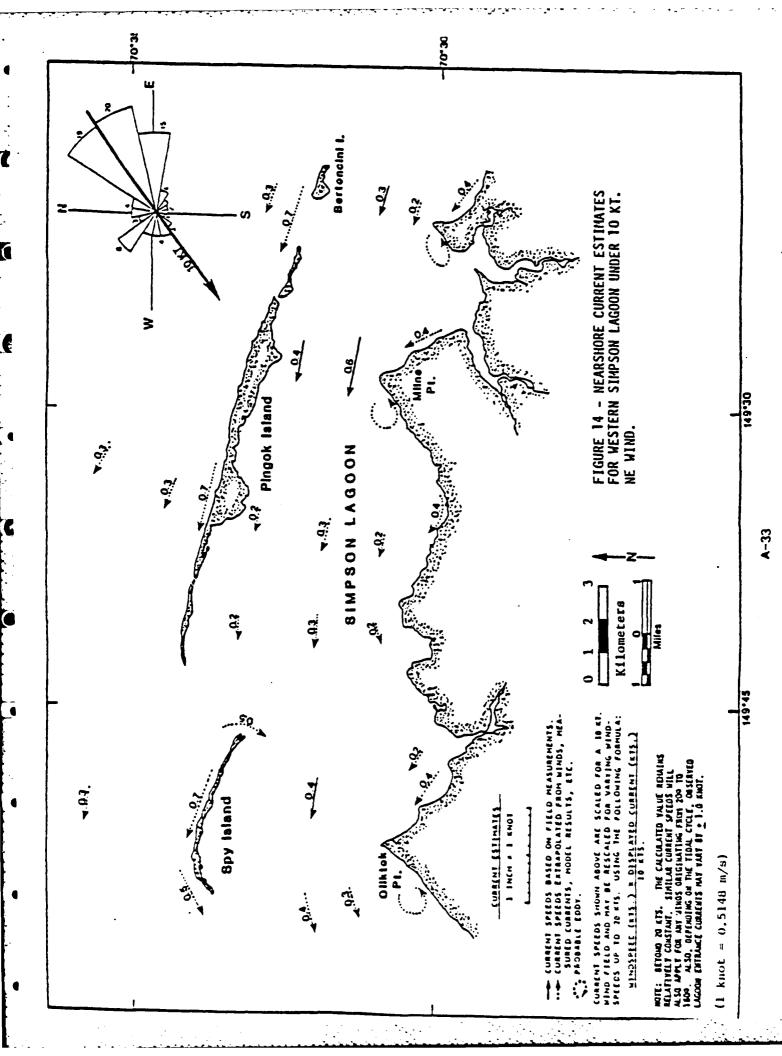


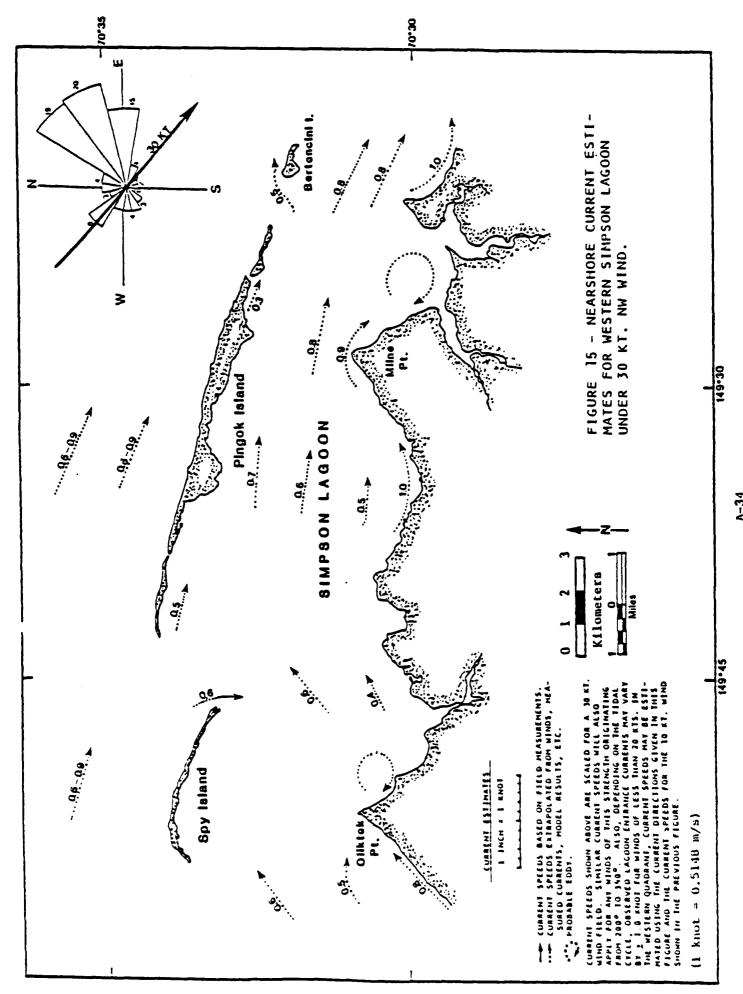
FIGURE 11 - LOCATIONS OF CURRENT ESTIMATE CHARTS FOR LAGOON AND NEARSHORE CIRCULATION.

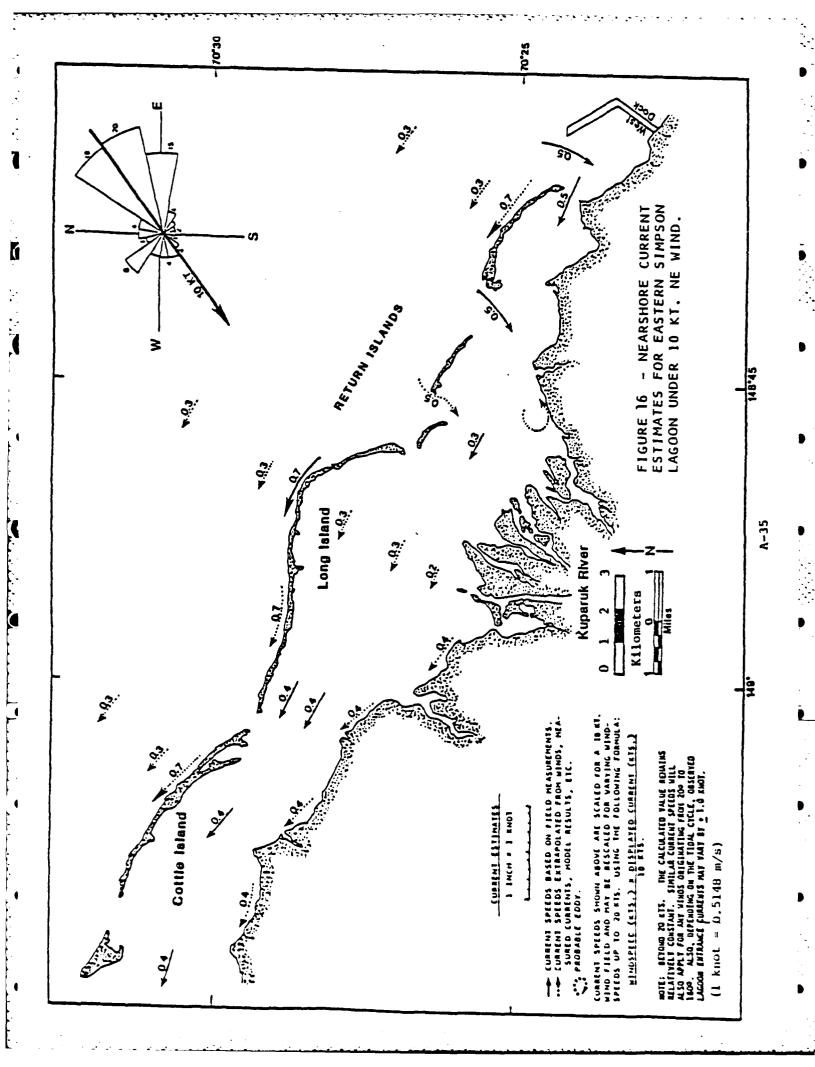
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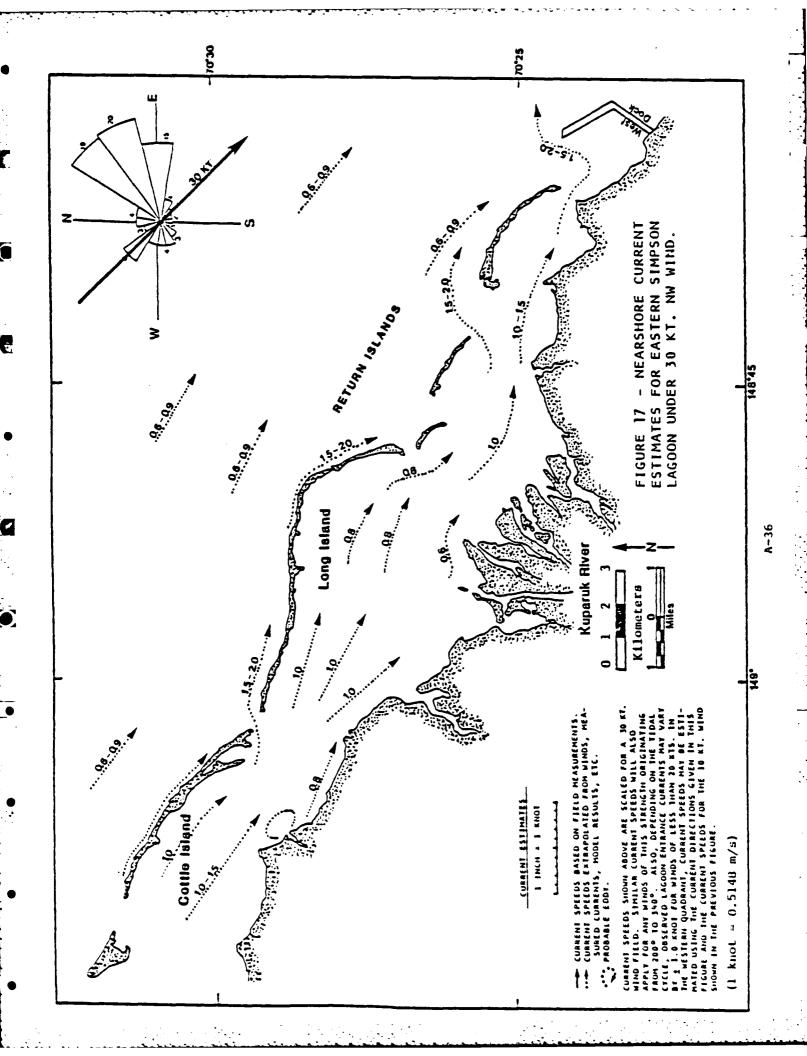


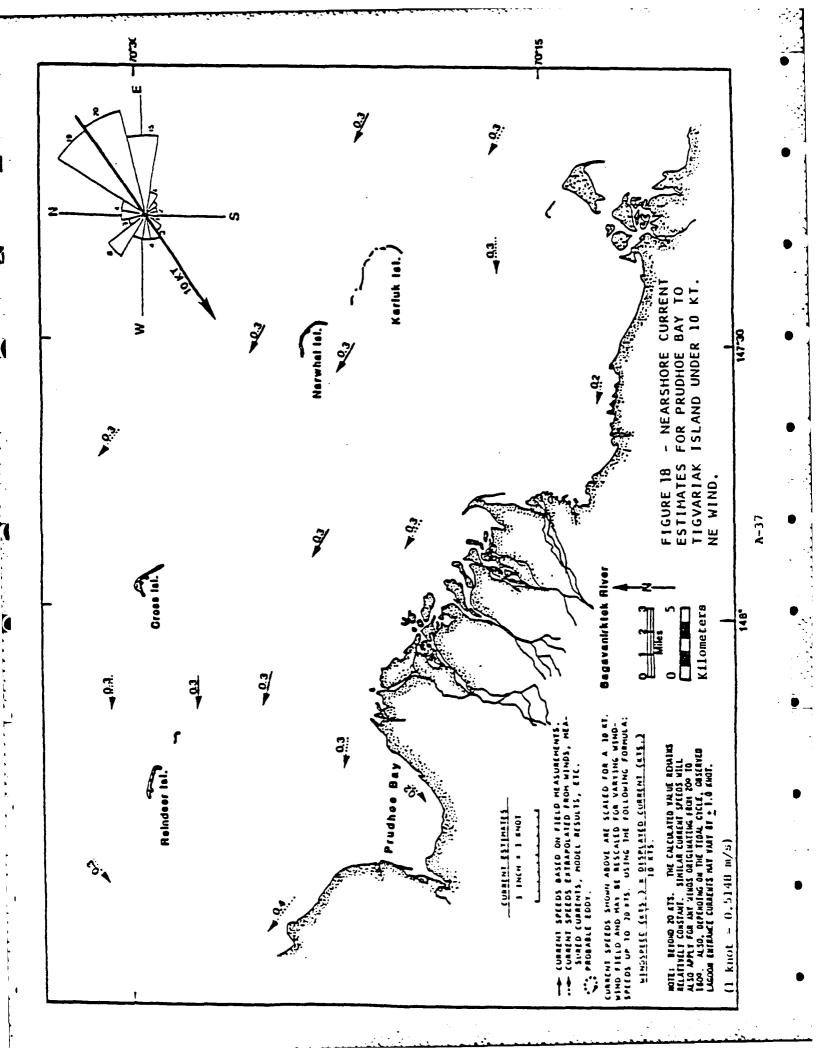


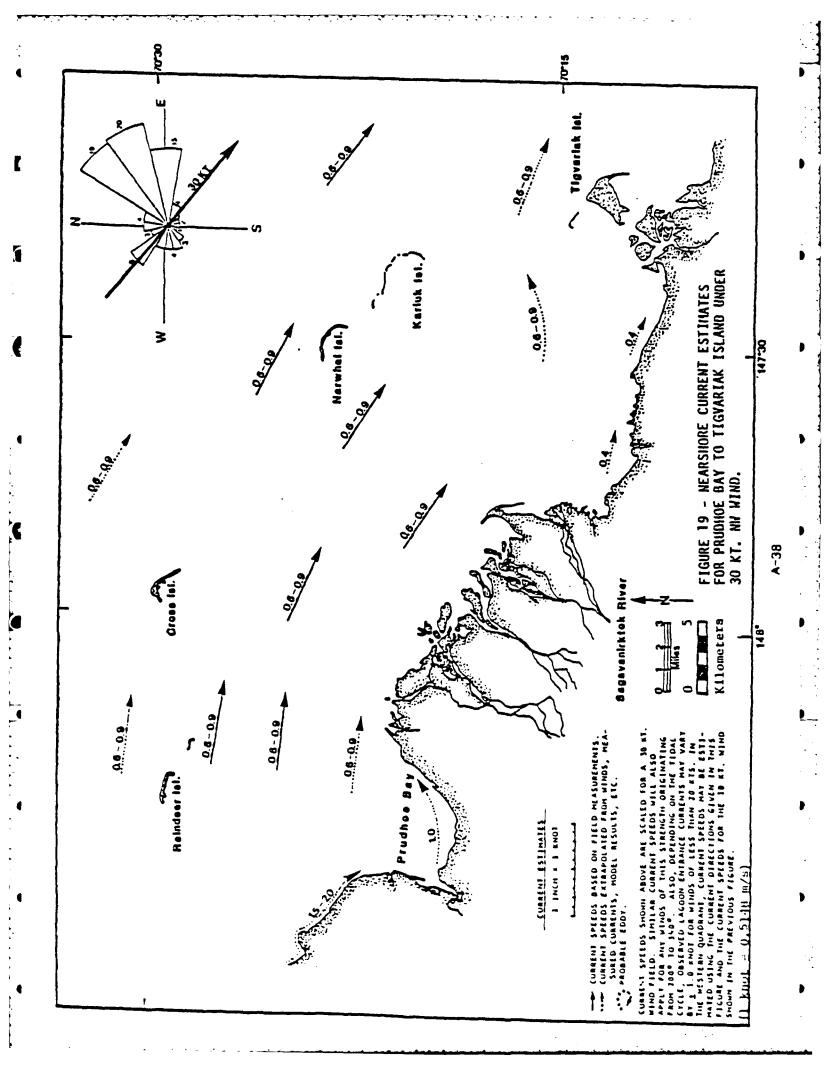


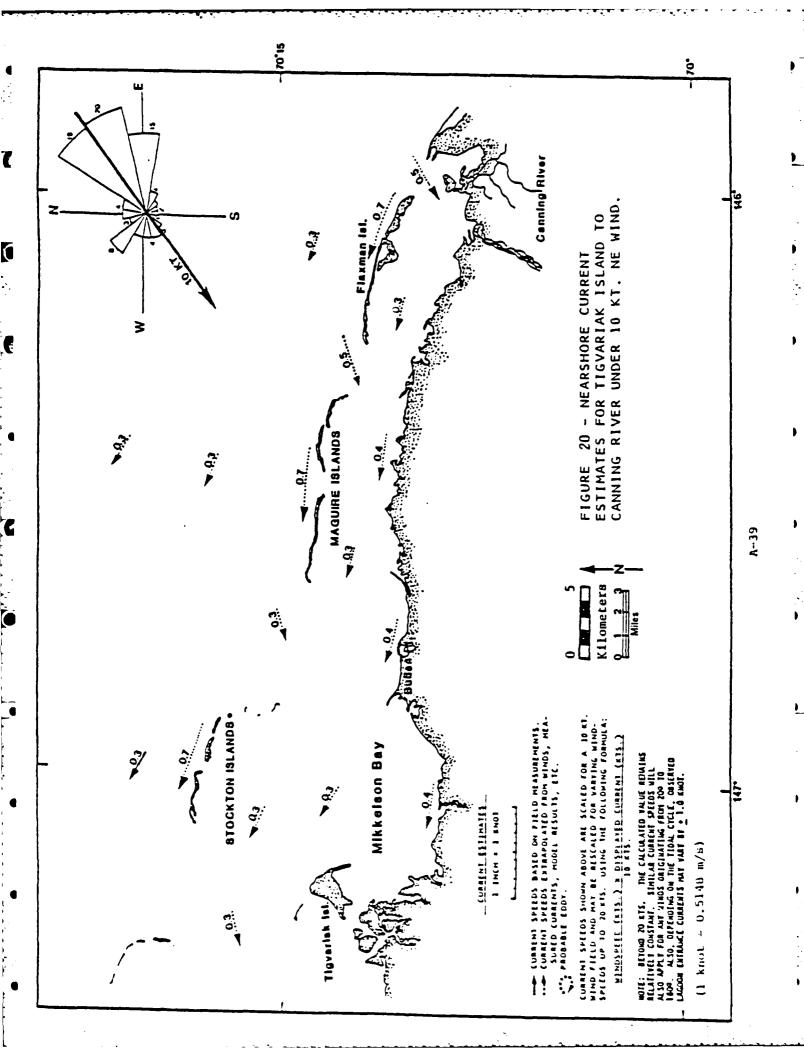




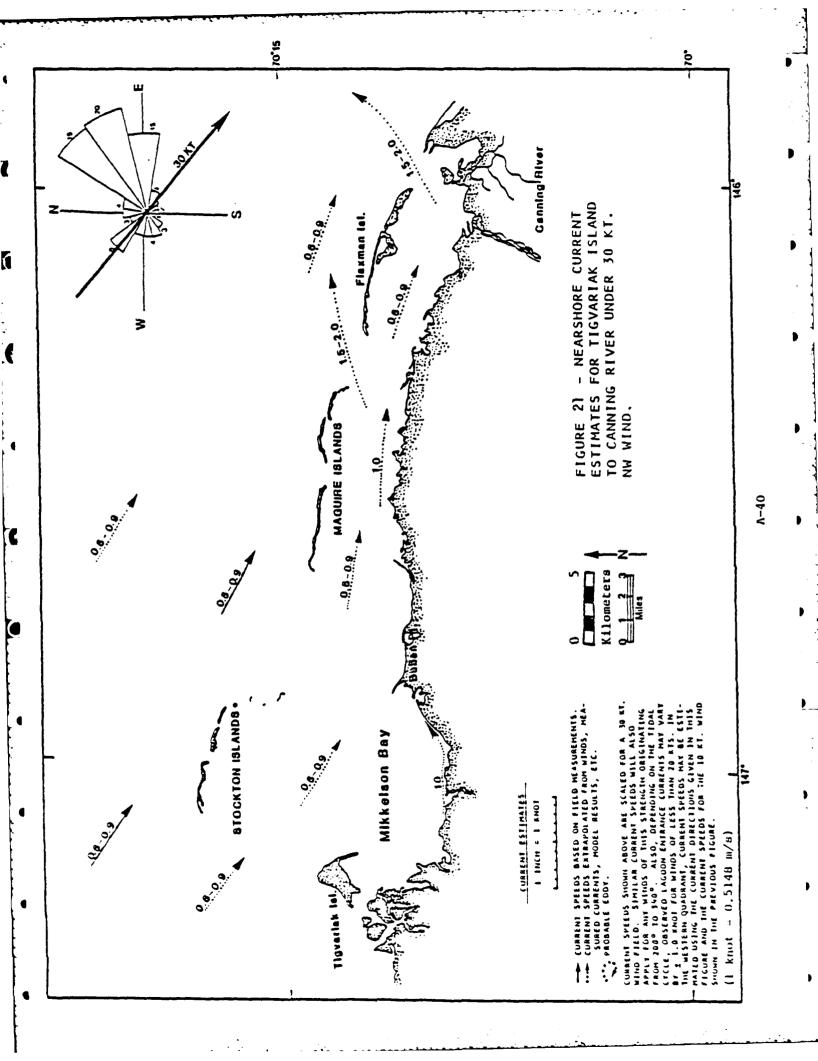


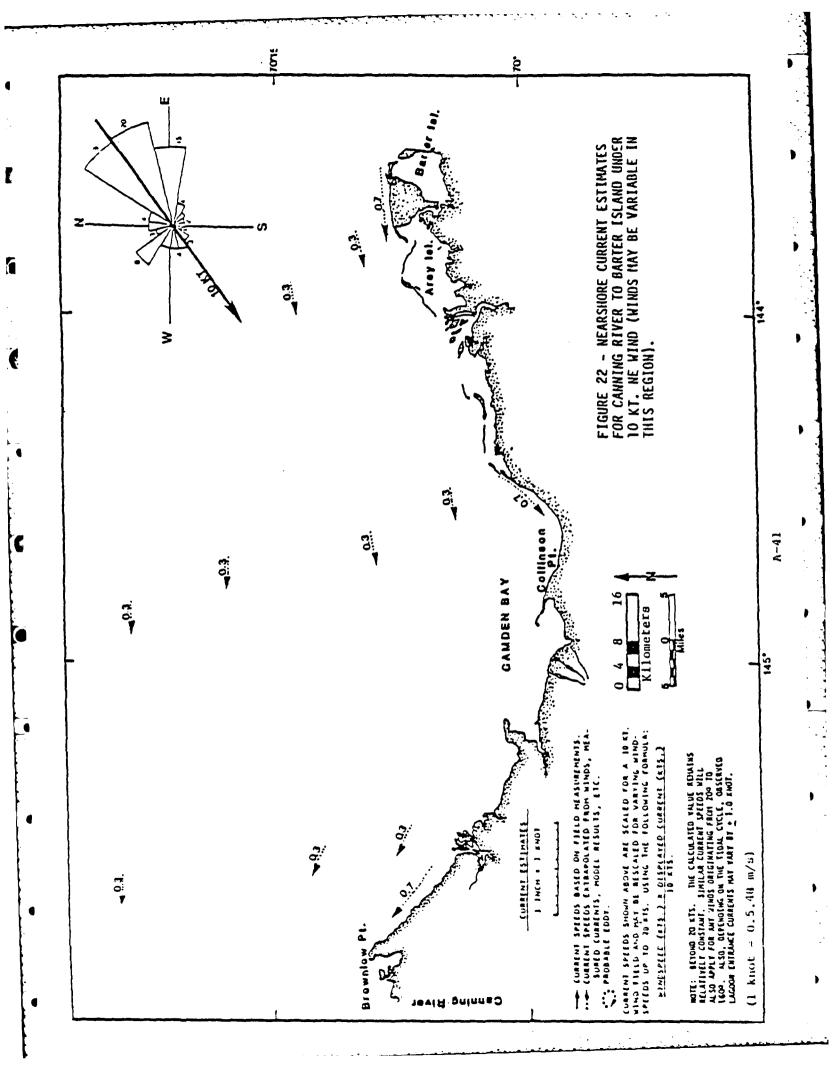


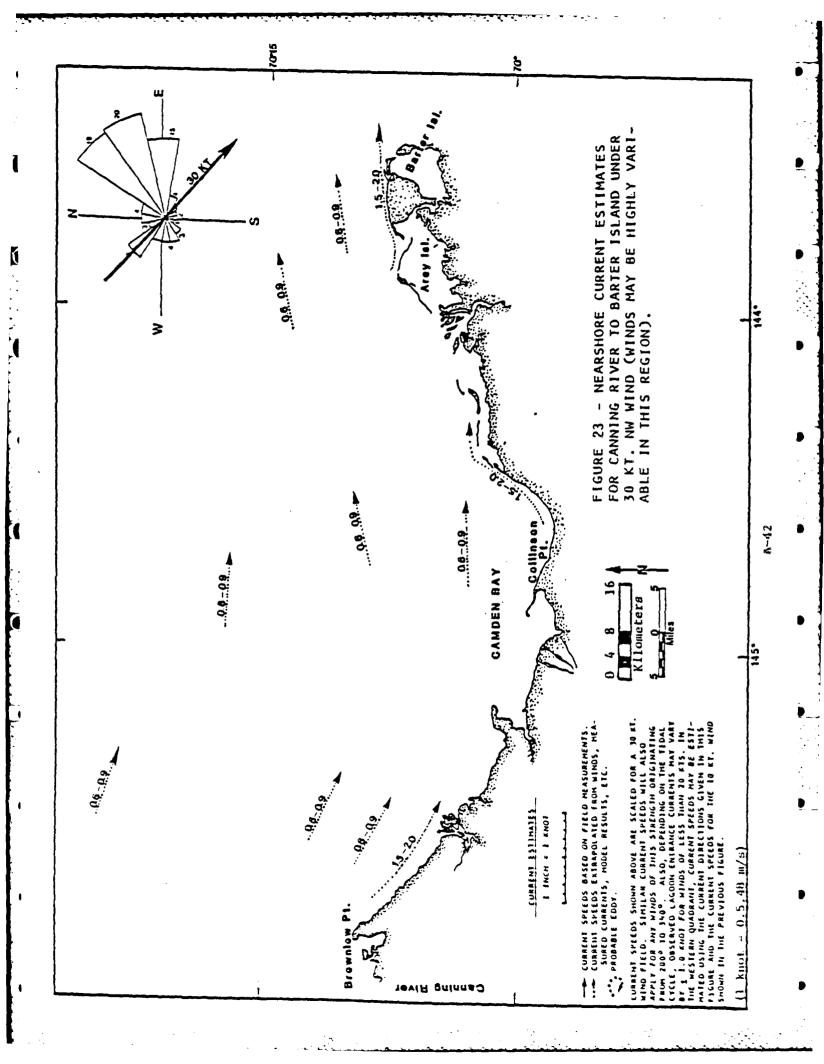


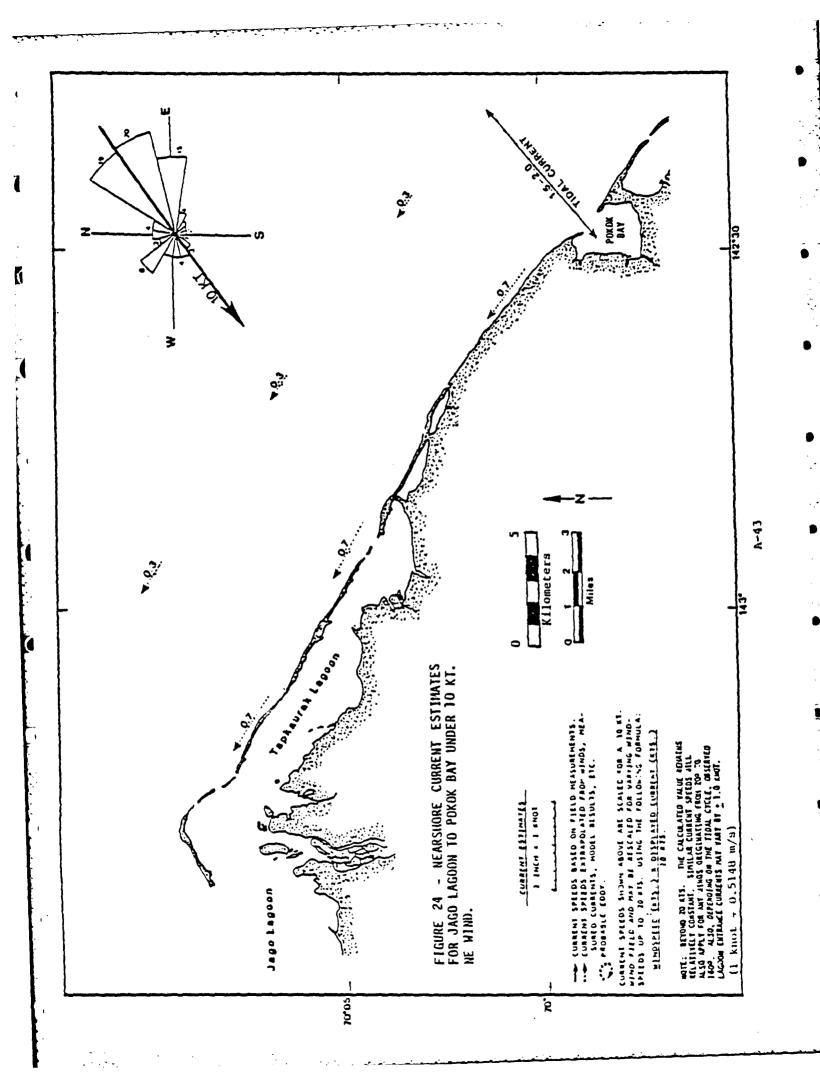


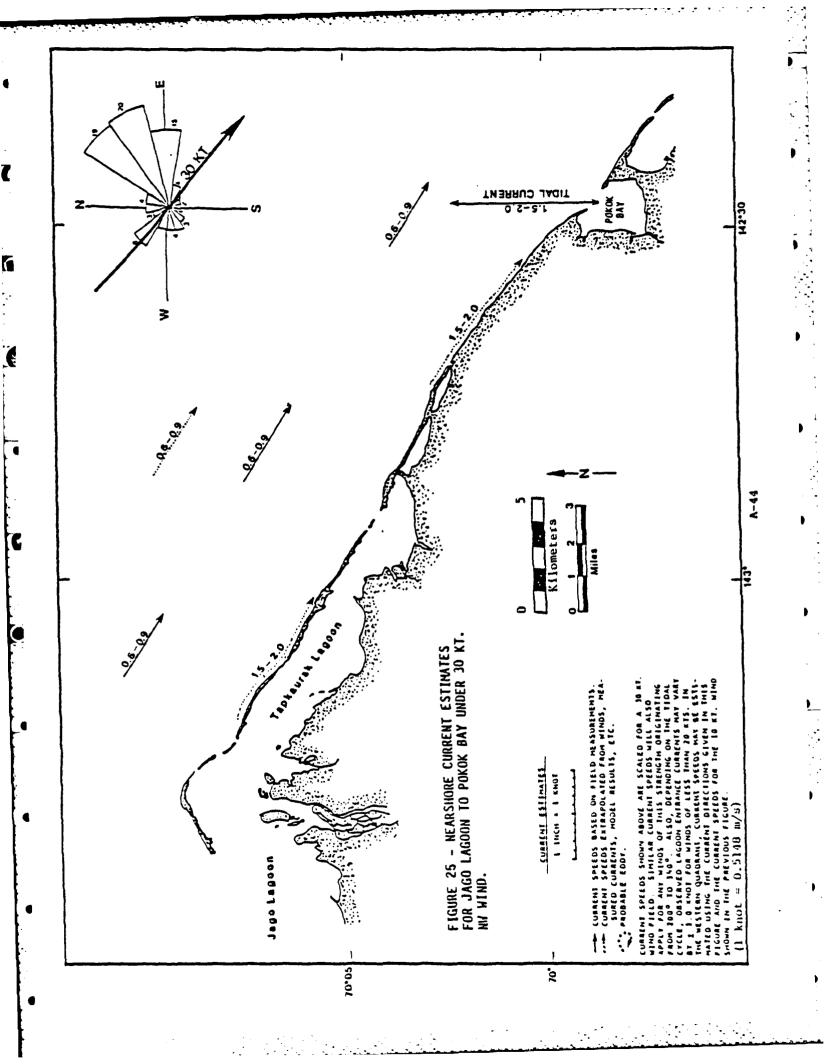
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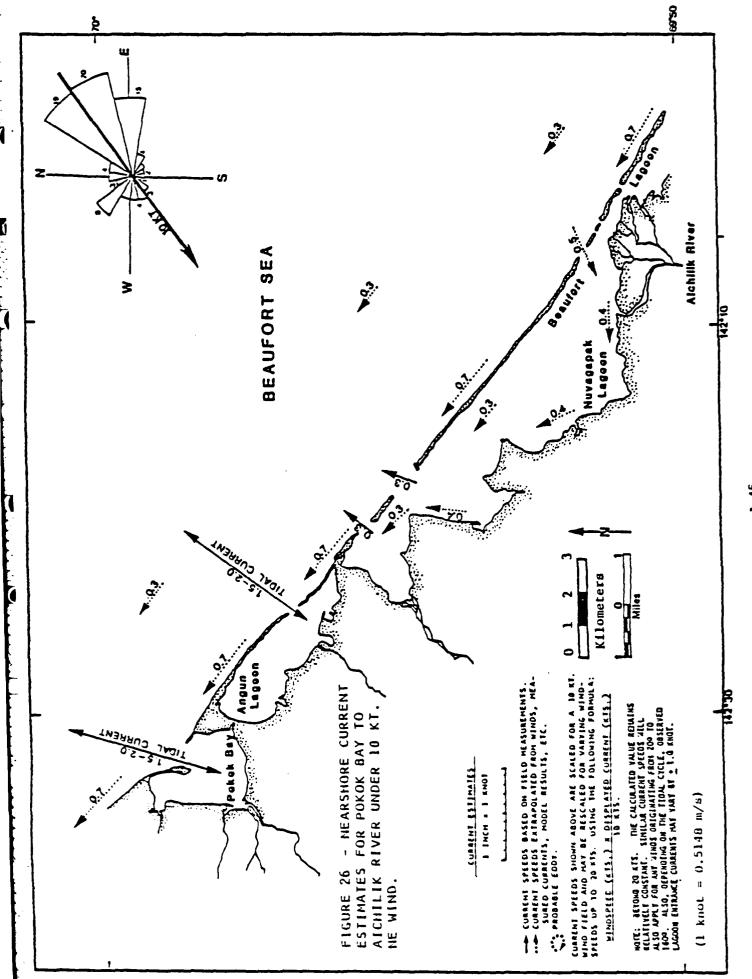


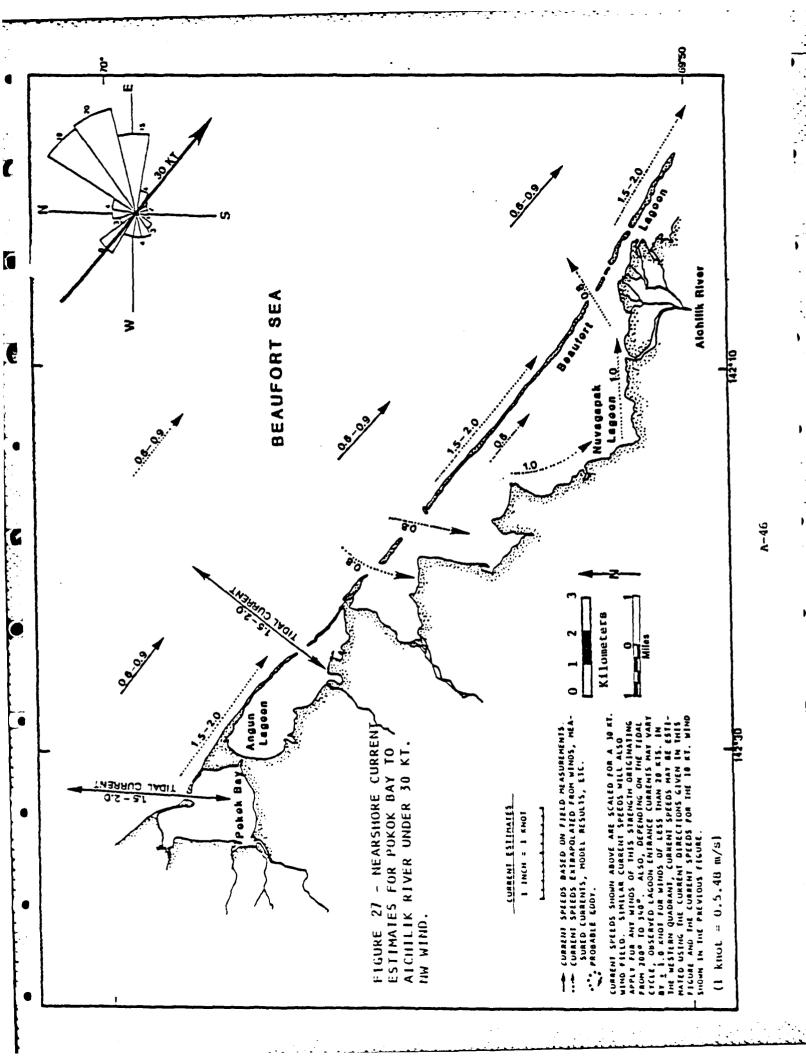


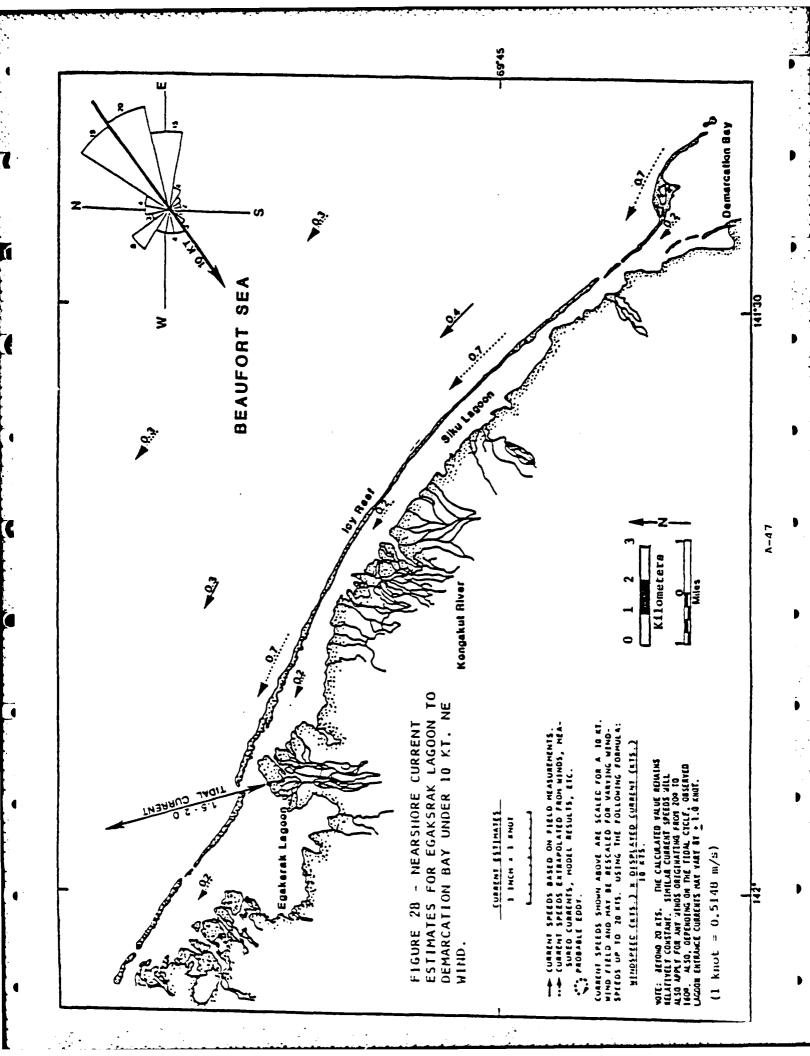


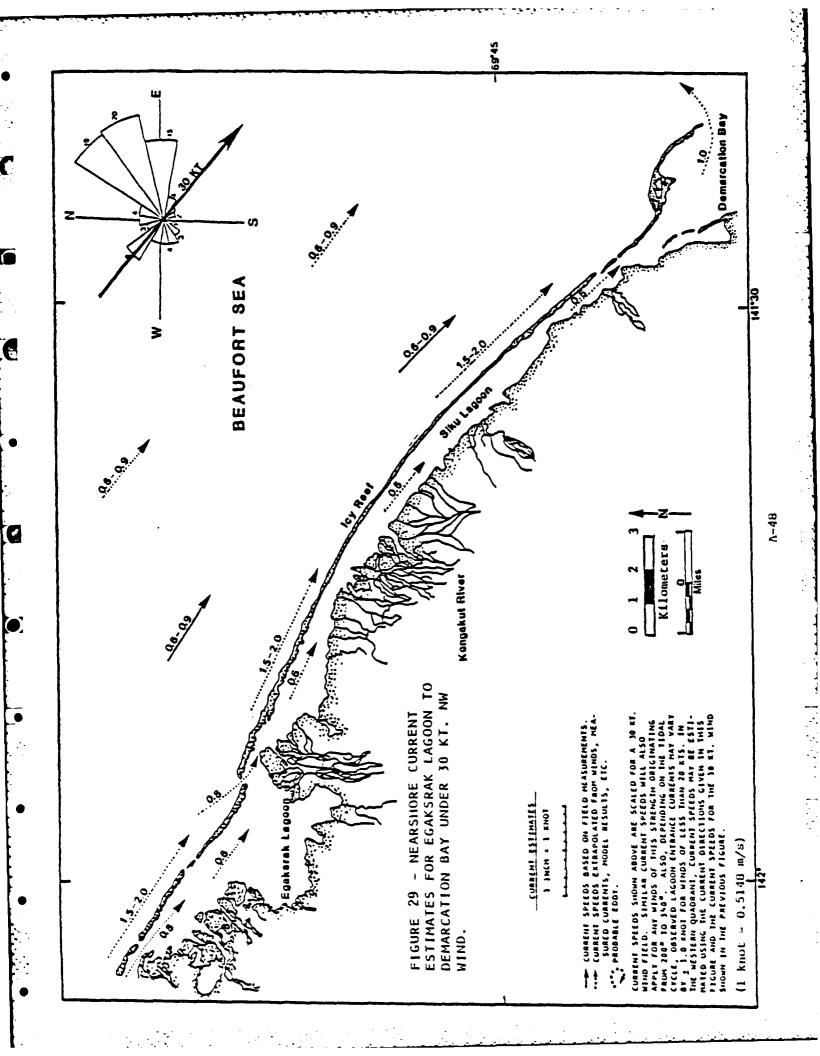












TIDES

Tides in the northeastern Chukchi and Bauufort Seas are generally quite small and are characterized by a mixed semi-diurnal signal with local maxima from 10-30 on in elevation. The largest and most studied tidal constituent in the Beaufort Sea is the M2 or lunar tide, with a period of 12.42 hours.

Although tidal data for much of the Beaufort Sea are scarce, the few meawork which does exist to describe the tidal behavior show the amplitude of the M2 component to be in the range of 5-10 ments indicates that the tide appears to Contours of equal M2 tidal heights (Figure 30) indicate that the tidal range increases from west to east along the Chukchi and Beau-Alaskan Beaufort Sea, the maximum measured M2 tidal amplitude was made near suraments and accompanying theoretical on (Kowalik and Matthews, 1982). Analysis of the phase of the tidal measureapproach the shelf from a north or fort Sea coastlines and begins to decrease slightly again approaching Barter Oliktok Point and recorded at 7.6 cm (Kowalik and Matthews, 1982; Wise and Island (approximately 143 °W). northwesterly direction.

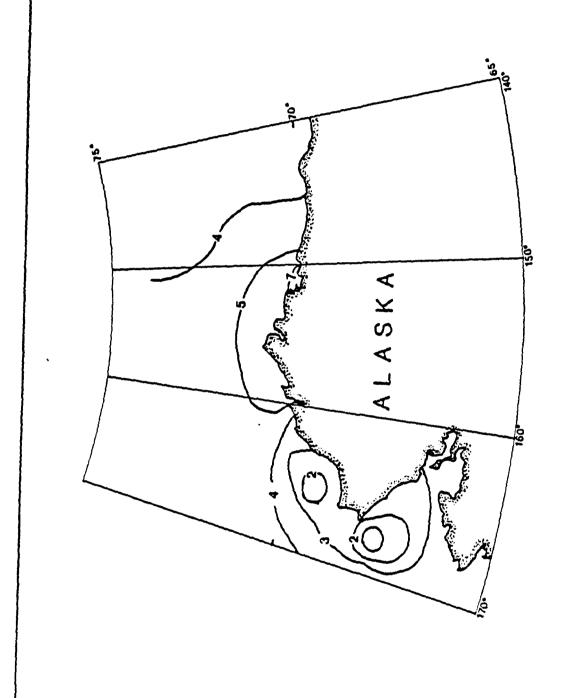
Searby, 1977). The actual daily tidal heights, which include the effects of all tidal constituents, can be obtained from standard tidal tables available from the Department of Commerce (various years). These tables allow determination of local tidal high and low water times and can be used to assess the relative magnitude of tidal effects for a particular day.

The minor axis indicates the direction orthogonal to the major axis erly current is present at high tide on central Beaufort shelf, then one These ellipses can be used to estimate the maximum expected tidal velocities in major axis of each tidal ellipse depicted in the figure indicates the dominant tidal current direction for a particular geographic region and the average magnitude of the tidal current in that direcaverage tidal current magnitude in a tion for each tidal ellipse indicates the temporal procession of tidal current directions as the tidal cycle progress-For example, if a 2.0 cm/sec eastof tidal motion. The direction of rota-Tidal ellipses are shown in Figure 31. a given area of the Beaufort Sea. tion.

quarter of a tidal cycle later the current vector would be rotated counterclockwise to the north and a smaller tidal current of approximately 1.0 on/sec in the offshore direction would be observed for the M2 tidal component. The predominant direction of rotation as indicated in Figure 31 is counter-clockwise for the Beaufort Sea. Local regions of clockwise rotation, however, are observed in the northeastern Chukchi Sea and in the Chukchi Sea south of the Lisburne Peninsula.

Figure 31 shows that the maximum velocity of the M2 tidal current diminishes rapidly at the boundary between the Chukchi and Beaufort Sea shelf and the Canadian Basin. This observed phenomenon is due to the depth discontinuity associated with the shelf break. There is also greater than 50% decrease (from > 2.0 cm/s to < 0.8 cm/s) in the maximum coastal M2 tidal velocity from the central to the eastern Beaufort Sea shelf.

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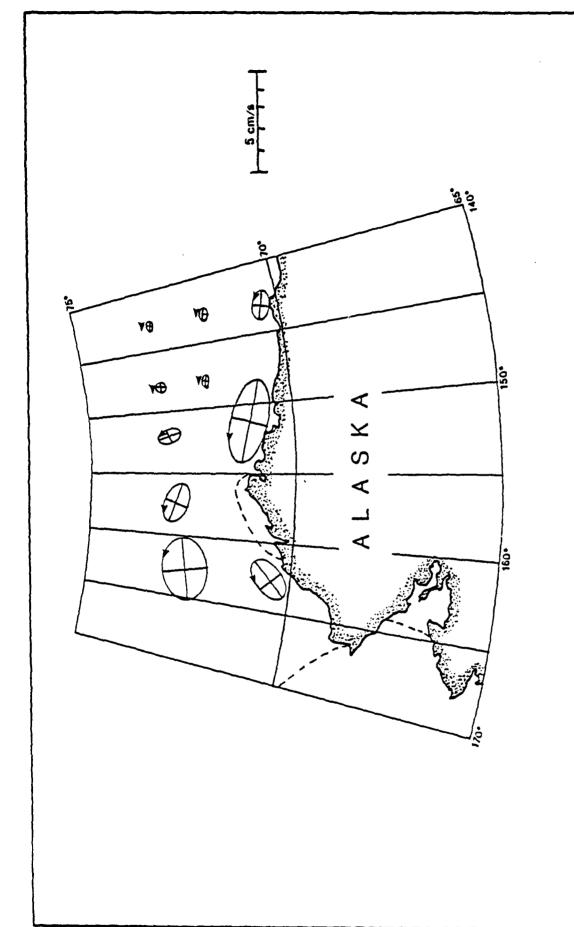


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FIGURE 30 - AMPLITUDE OF THE M2 TIDE IN THE BEAUFORT SEA (FROM KOWALIK AND MATTHEWS, 1982). AMPLITUDE LINE UNITS ARE CM.



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FIGURE 31 - TIDAL ELLIPSES IN THE BEAUFORT SEA; ARROWS DENOTE DIRECTION OF ROTATION ALONG TIDAL ELLIPSE. BROKEN LINES INDICATE BOUNDARIES BETWEEN AREAS WITH DIFFERENT ROTATION (FROM KOWALIK AND MATTHEWS, 1982).

RIVER DISCHARGE

7.0. RIVER DISCIARGE

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Only a few of the rivers along the Beaufort Sea coast are gauged, and these are not always regutering the Beaufort Sea include the Colville, the Sagavanirktok, the Kuparuk, and the Cauning. Hydrographic data for these rivers is given in Table 1. Seawater intrusion into high variability in seasonal annual flow with spring runoff being the major con-Spring and sunner discharge of greatly affects the salinity and tur-Arctic coastal river discharges have not The major rivers en-None of these rivers has year-round flow; all cease flowing by late Januar, and begin to flow again in late May and river deltas occurs from mid-autumn All rivers have the Colville River and lesser rivers tributor to annual flow volume (Carlson, hidity of the nearshore Beaufort Sea. been well documented. through the winter. larly monitored. early June. 1976).

The Colville River is the largest river entering the Beaufort Sea. During spring thaw in June, the Colville River discharges 50 percent of its annual flow. It continues to flow as late as January, with no measurable discharge from then until late April or early May. In the winter months, beawater reaches as far upstream as the Itkillik River.

Annual discharge of the Colville River is 12 km³ (Schell and Horner, 1981), which is about 73 percent of the total discharge of all rivers between the Colville and Canning Rivers.

inundation may be between 0.5 and 1.0 m (Reminitz and Bruder, 1972). Measured four weeks prior to breakup of fast ice along the coast. During this time there river size and volume of snowmelt. The front of fresh water may move at a rate of 2 to 3 m/s over the ice. Depth of Over-ice flood extent from the fron year to year, and is a function of areas of flooded sea ice are shown in River flow from snow melt begins about is a large amount of fresh water forced both under and over the ice in the near-The extent of the overdata are not available for ice flow of river water varies greatly either the Sagavanirktok or Kuparu Colville River is shown in Figure shore region. Table 2. Graphic Rivers. The flow of the Kuparuk and the Sagavanirktok Rivers have been well documented because of the industrial activity in the region. Maximum and minimum mean daily flows of these rivers are shown in Figures 33-34.

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HYDROGRAPHIC DATA FOR MAJOR RIVERS ENTERING THE ALASKAN BEAUFORT SEA

COLVILLE RIVER

Average Annual Flow: 12,000 cubic feet per second (cfs).

extent of under-ice an a Seaward extent of flooded sea ice is During spring thaw in June the Colville River discharges 50 Percent of its annual volume. Winter undocumented, but generally occur the last week of May to flow is assumed to be 0 or negligible, This is plume is 40 km. the first week of June. are 8-18 km; maximum unganged river. dates Breakup water

KUPARUK RIVER

Average Annual Flow: 1305 cfs.
Miximum Flow: 82,000 cfs.
Minimum Flow: 10 cfs.

Seaward extent of flooded sea ice is to the barrier islands. The hydrographic regime is shown in Figure 34. River starts flowing from 29 May - 8 June; flow usually ceases between 20-29 September. Sixty to eighty percent of annual runoff occurs in the month of June.

SAGAVANI RKTOK RIVER

Average Annual Flow: 2770 cfs.
Maximum Flow: 20,000 cfs.
Minimum Flow: 10 cfs.

Hydrographic regime is shown in Figure 33. Flow starts 13-22 May; freeze-up occurs between 27 September and 6 October. Sixty to eighty percent of annual runoff occurs in June.

CANNING RIVER

Average Annual Flow: 1125 cfs.

This is an ungauged river. No ice overflow has been mapped; no seasonal discharge information.

(1 cubic foot = 0.028317 cubic meters)

Table 2

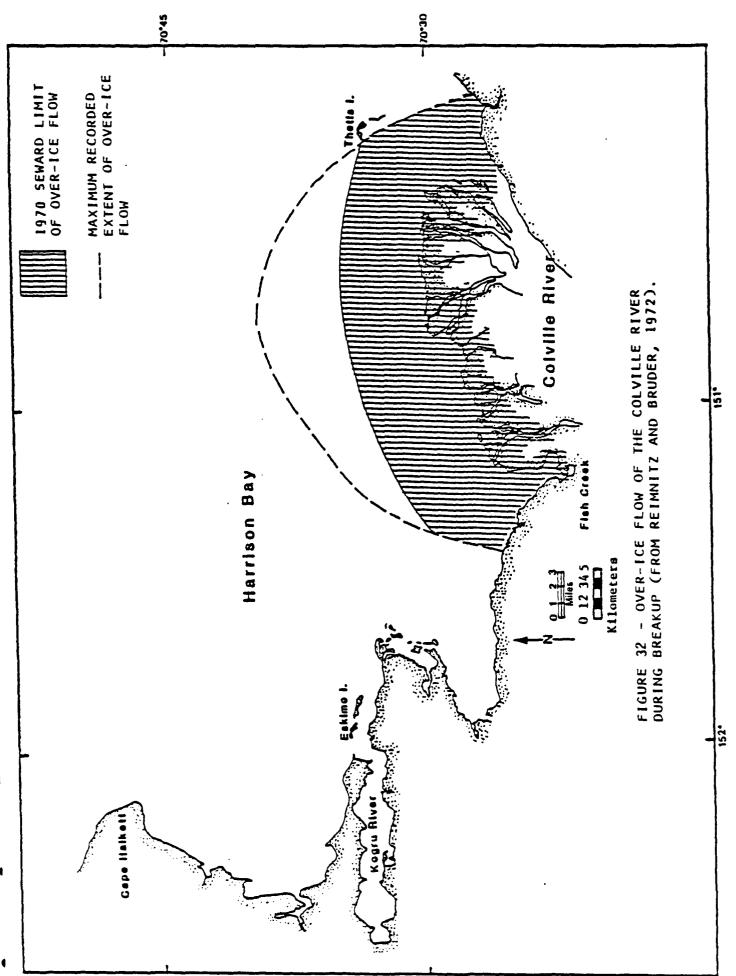
AREAL EXIENT OF RIVER OVERFLOW ON SEA ICE IN SPRING

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Colville River (km ²)	15	50	61	120	219	276	100
Sagavanirktok River (km²)	19	151	185	40	208	179	100
Kuparuk River (km ²)	I	10	()	30	101	69	1
Date	21 May 1974*	26 May 1974*	4 June 1974*	6 June 1974*	4 June 1975*	9 June 1975*	6 June 1976**

*From Carlson (1977).

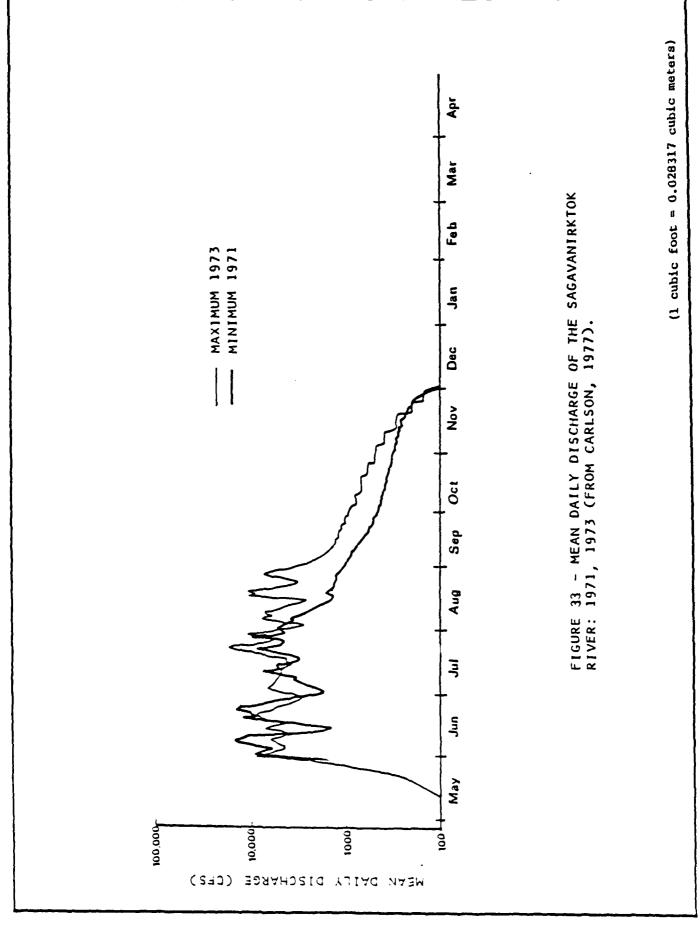
**From Barry (1979).



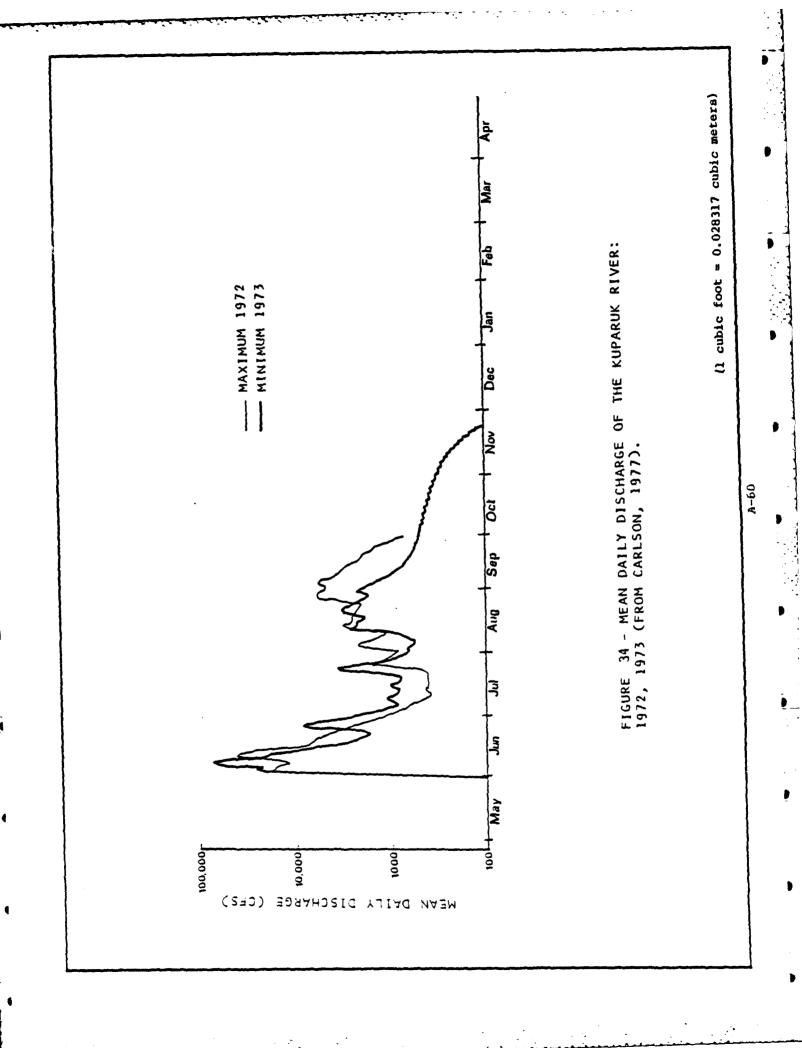
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STORM SURGE

B.O. STORM SURGE

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local wind regime, because of water experiences relatively anall changes in ations may range from as much as +3.0 m unpublished data). The largest positive when long stretches of open water are common and the winds have become preduninantly westerly, driving water onto little or no warming. Water may begin to rise before a change occurs in the being driven onto the shelf ahead of the wind. Maximum wave size (up to 3 m) is reached within a few hours of the onset The entire Beaufort coast in general however, meteorologically-induced varithe shelf. Storm surges may occur with storm surges typically occur in the fall Hea level due to astronomical tides! to -0.9 m (Schaeffer, 1966, Matthews,

of the storm. Swift easterly currents of 2-3 kts may occur, piling up water on the windward side of the coast.

The best documented storm surge occurred on 13 September 1970 when NW winds of 80 km/hr were recorded at Oliktok Roint (Figure 35). Many of the barrier islands off Simpson Lagoon were inundated, and some barges in Prudhoe Bay were lifted out of the water and onto the causeway to which they were secured (Reimnitz and Maurer, 1978). Intervals at which events similar to this occurrange from 25-50 years in the central portion of the Beaufort Sea coast. Near Barrow, surges are recorded much more often (Reimnitz and Maurer, 1978).

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FIGURE 35 - HEIGHT OF 1970 STORM SURGE ABOVE MEAN SEA LEVEL, AS MEASURED FROM THE ELEVATION OF DRIFTWOOD FOUND ON THE MAINLAND AND ON THE ISLANDS. NOTE PILE-UP OF WATER ON THE EAST SIDE OF SHALLOW EMBAYMENTS AND THE LAGOON NEAR THE COLVILLE RIVER (FROM REIMNITZ AND MAURER, 1978).

ICE DRIFT

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9.0. ICE DRIFT

G

The Beaufort Sea shelf is essentially ice covered for all but two to three Breakup typically begins in early June which coincides with the initiation of spring river As breakup proceeds, a nearshore lead forms from Pt. Barrow to Demarcation Bay which may kilometers to several hundred kilome-Open water conditions persist through September and the refreezing vary in offshore extent from several typically begins in early discharge (Section 7.0). months of the year. process October. Nearshore ice takes on several forms as the freeze-up process proceeds. Adjacent to the coastline, shorefast ice begins to form and depending on the time of the year may extend out to 70 km from the coastline in the central and western Braufort. This ice remains relatively motionless throughout the winter months. Offshore of the fast ice is the grounded ice. This is typically ridged and quite thick and stretches along the nearshore at approximately the 20 to 30 m depth contour. This ice is also relatively

motionless throughout the winter months. Beyond the grounded ice is the shear zone where the moving ice and the shorefast ice meet.

Few measurements of ice motion exist in or near the shear zone. Beyond the shear zone the ice pack moves at varying speeds and directions under the influence of the wind field. However, mean ice drift similar to the mean wind field exhibits motion to the west. Recent measurement programs have studied this westward drift to determine mean ice speeds and the spatial variability of ice motion.

As part of the Arctic Basin Buoy Program, automatic data buoys were deployed on the pack ice of the Arctic Basin in the years 1979-1982 to monitor large-scale ice movement (Thorndike and Colony, 1979; 1980; Thorndike et al., 1981). Position data from these buoys were transmitted at one-minute intervals and are available in several documents published by the Arctic Basin Buoy Program.

Manthly averages of the proy motion were determined from these data and are displayed in Figures 36-39. As anticipated, the general drift of the loe is from east to west. Average monthly ice specis were also determined for one degree latitude bands between 71-74 °N. These values are displayed in Table 3. Manthly averages range from a low of

3 cm/s to a high of 20 cm/s based on the buoy drift data. The peak in the speed of ice drift occurs in the fall, with the lowest ice drift speeds found in late winter and early spring.

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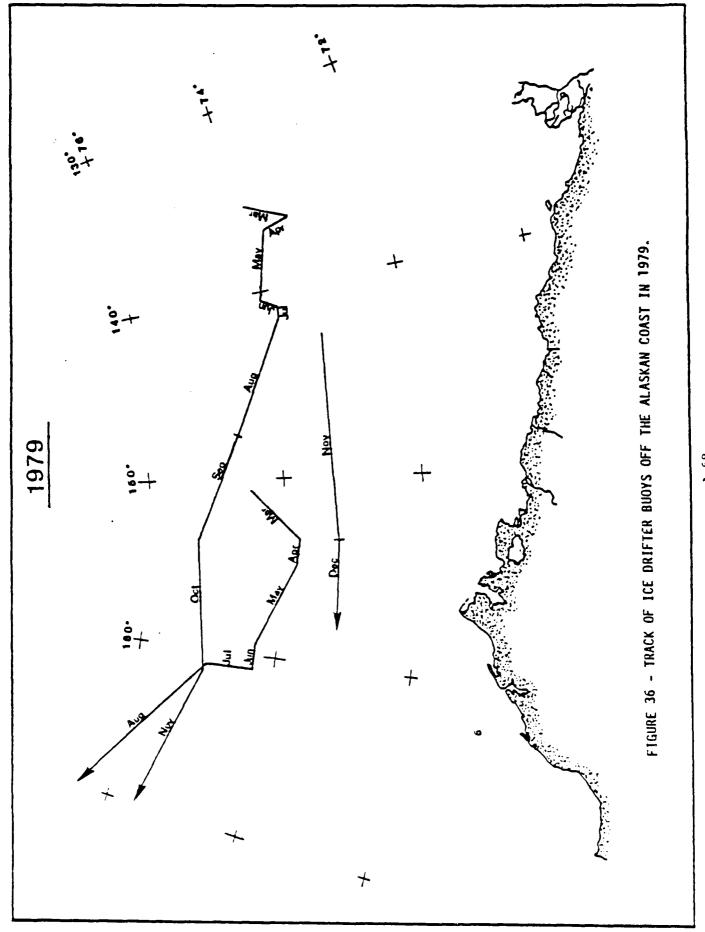
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All data analyzed for this section were for buoys deployed in the Beaufort Sea between 70-76 "N and 150-170 "W.

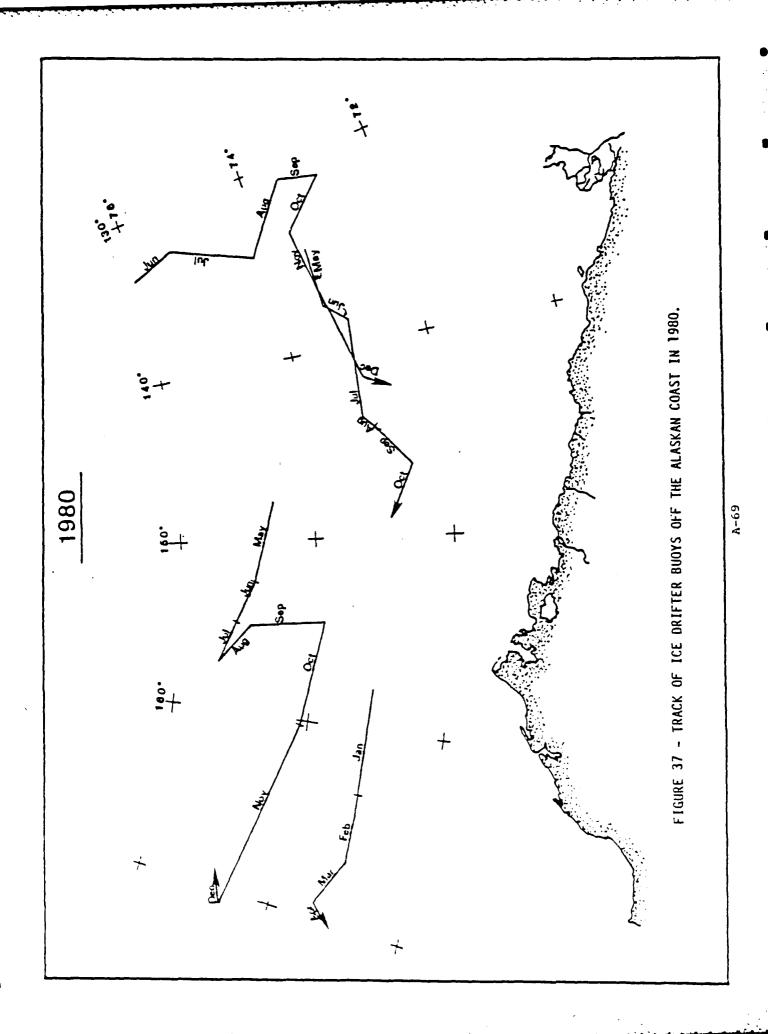
Table 3
AVERAGE MONTHLY ICE DRIFT SPEEDS (cn/s) VS. LATITUDE

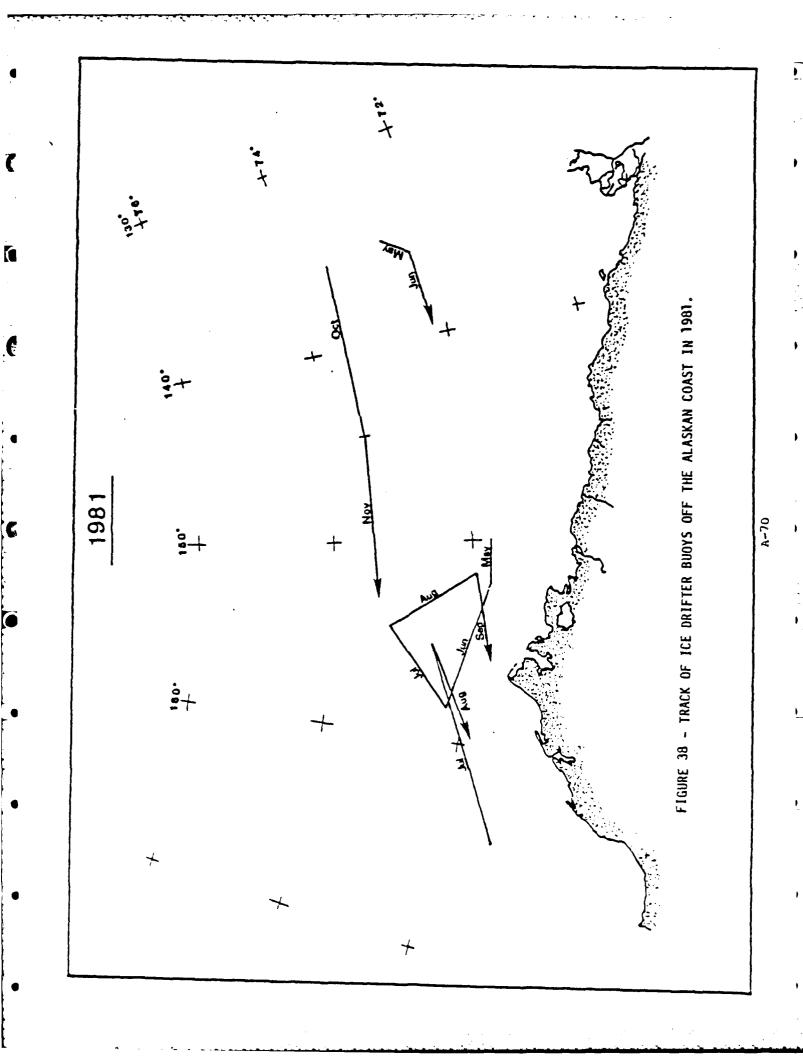
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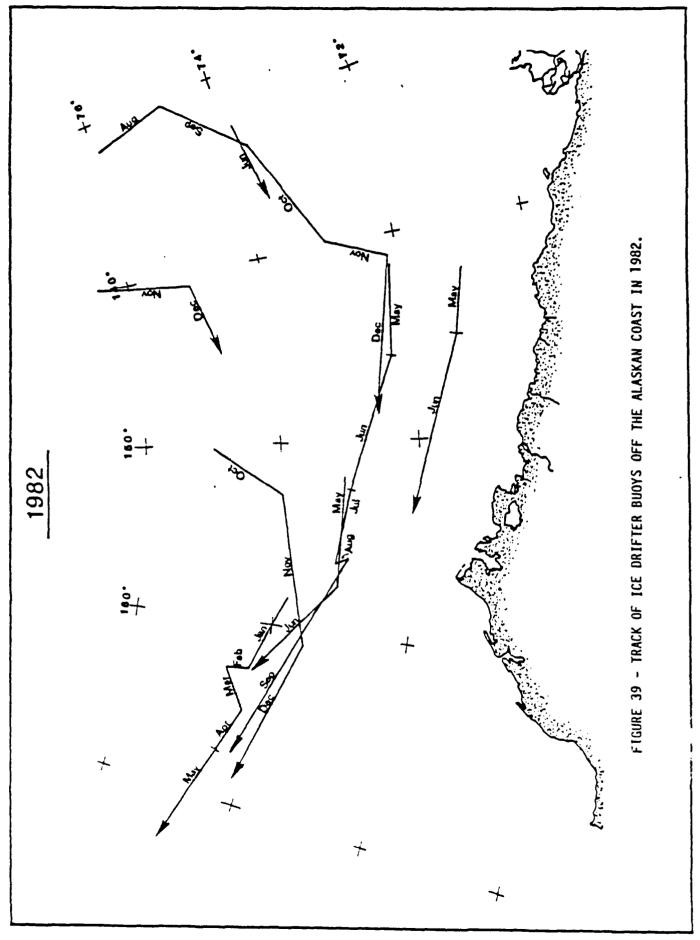
Jan Feb Mar Apr May Jun Jul Aug Sep Oct Nov Dec	13 7 9 9 11
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71 - 72°N 3 6 12 14 13 13	



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BATHYMETRY

10.0. BATHTYMETRY

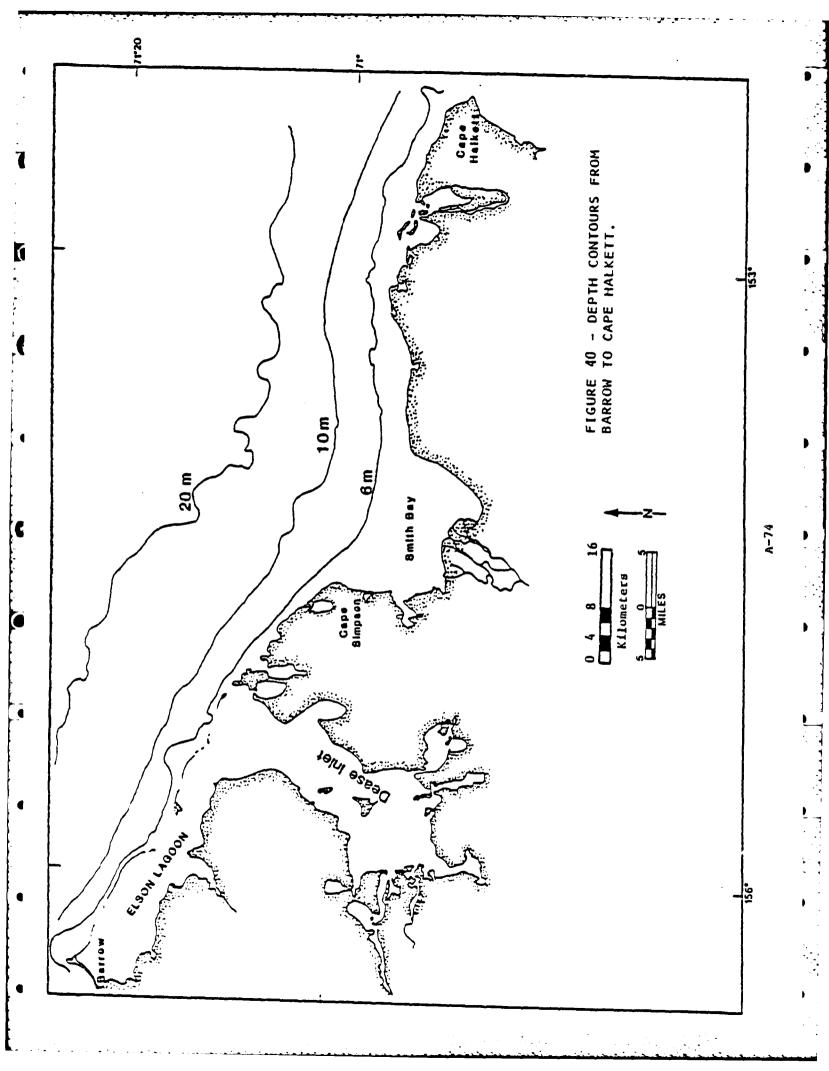
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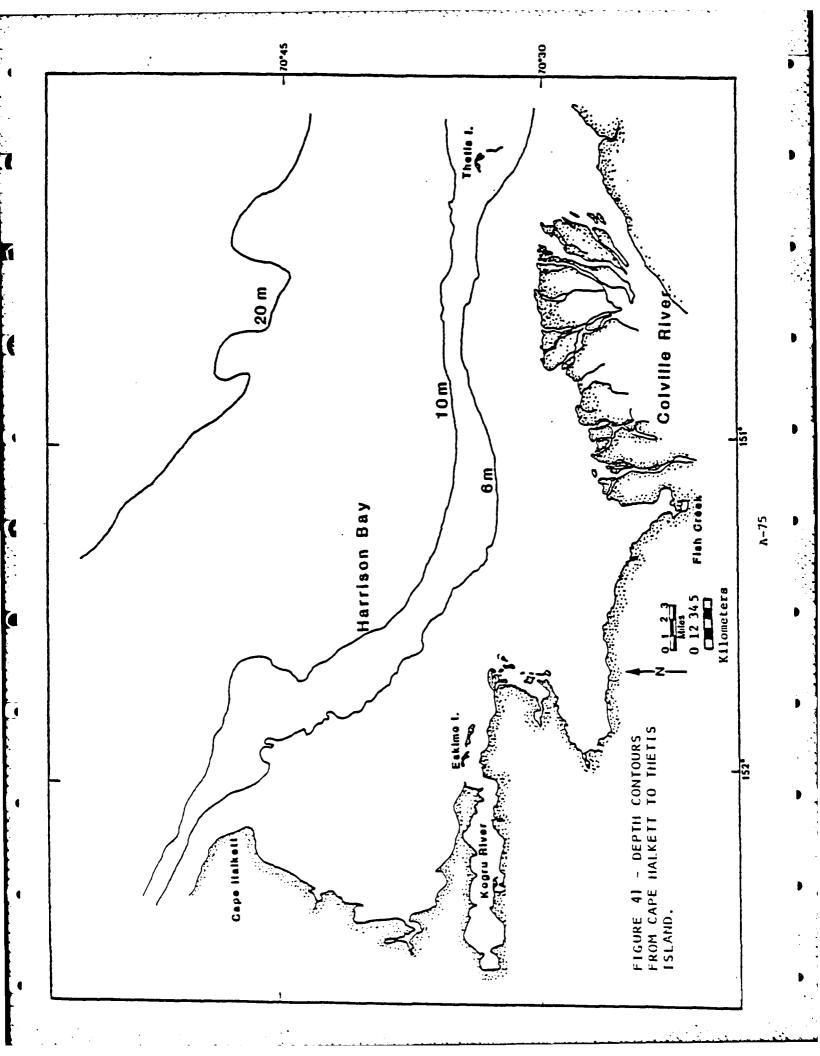
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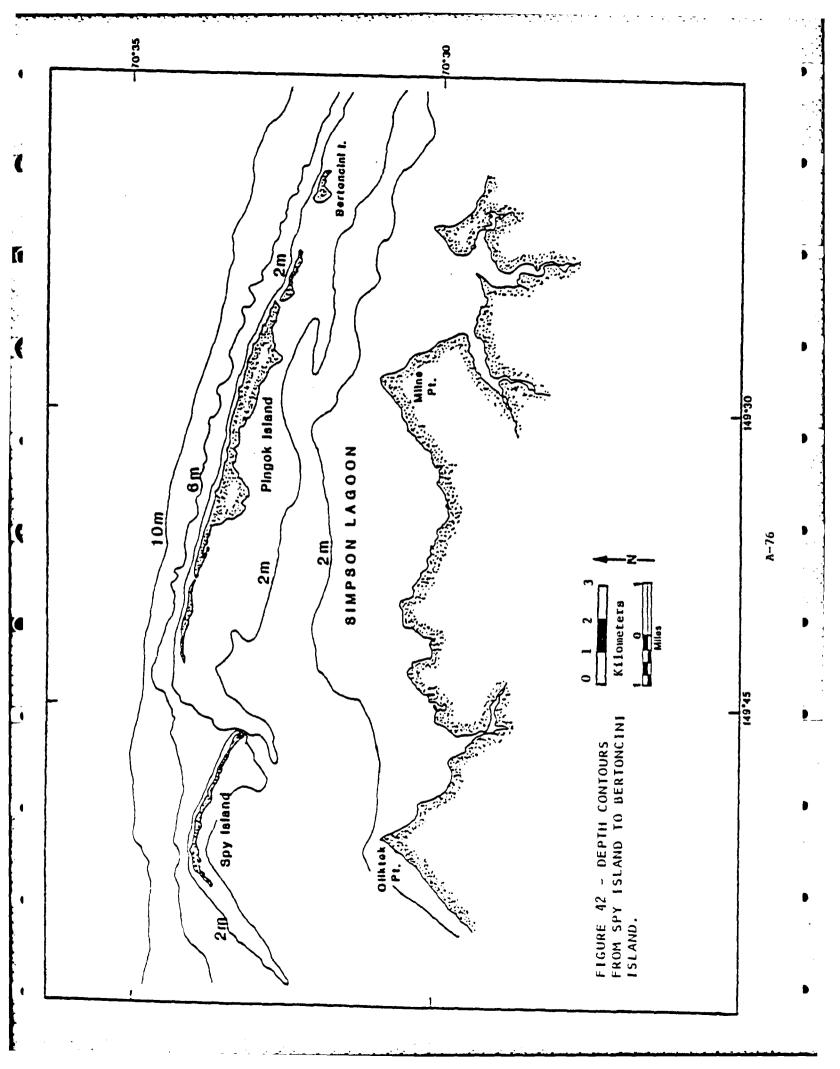
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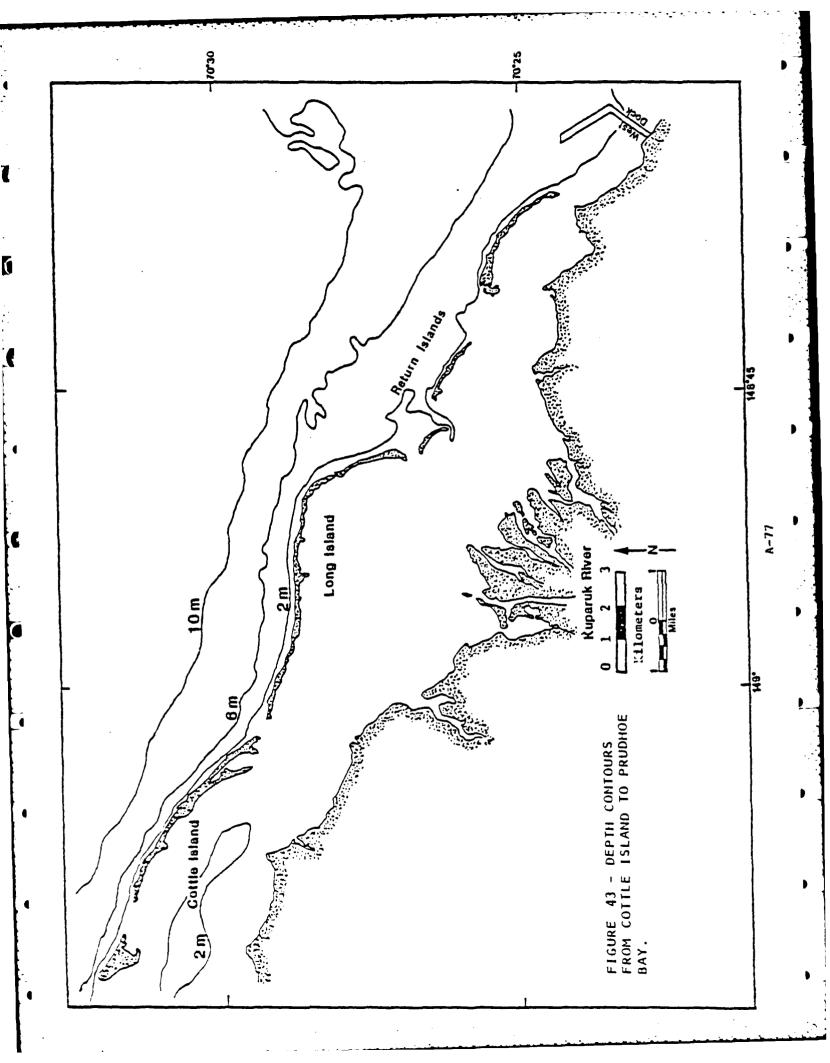
Figures 40 through 49 which follow display generalized bathymetry of the nearshore Beaufort Sea. These contours were derived from NOS Nautical Charts and U.S. Geological Survey 7-1/2' maps. In the nearshore Beaufort Sea, bathymetry changes rapidly and, at times, diastically. Coastal erosion rates are high (1.6 m/yr average over the whole coast), and result from both mechanical and thermal processes. (Ropkins and Hartz, 1978). As a result of this

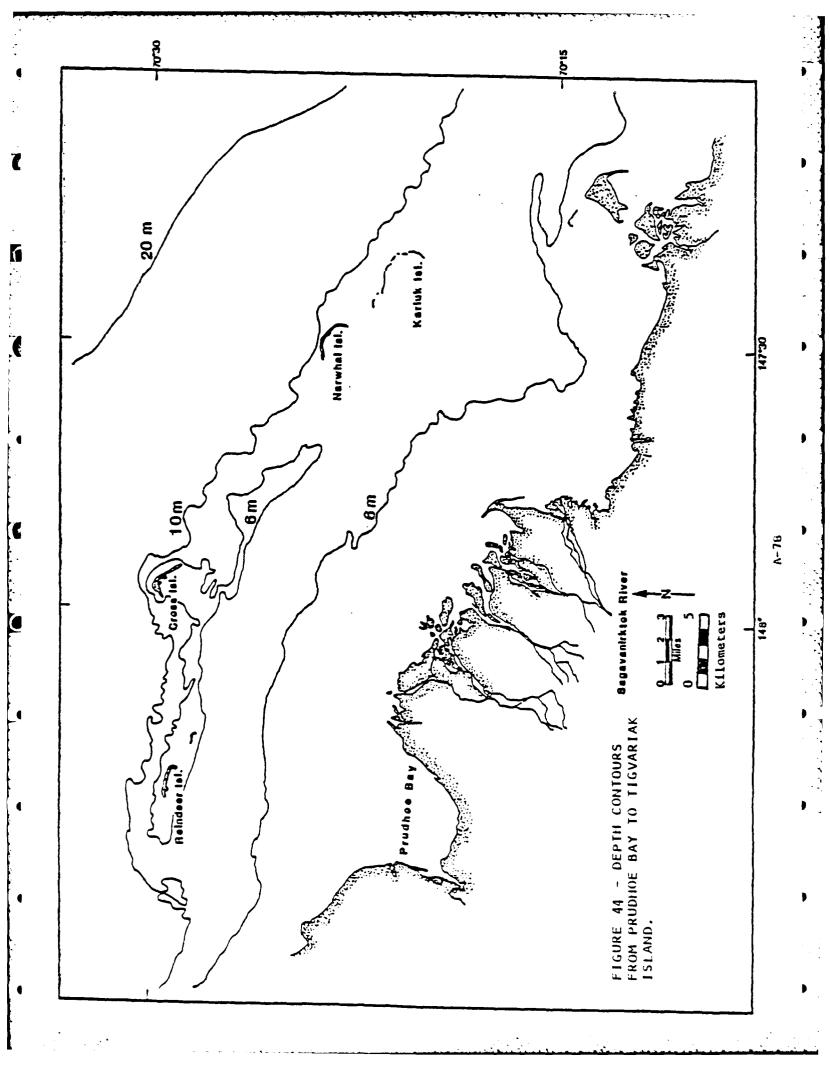
erosional pattern, sediment moved longshore during the open water season can rapidly change water depths and location of navigable channels. This makes navigational charts and bathymetric maps unreliable for all but small boats. The maps which follow are intended to give the reader a general overview of the expected depth contours in the coastal zone.

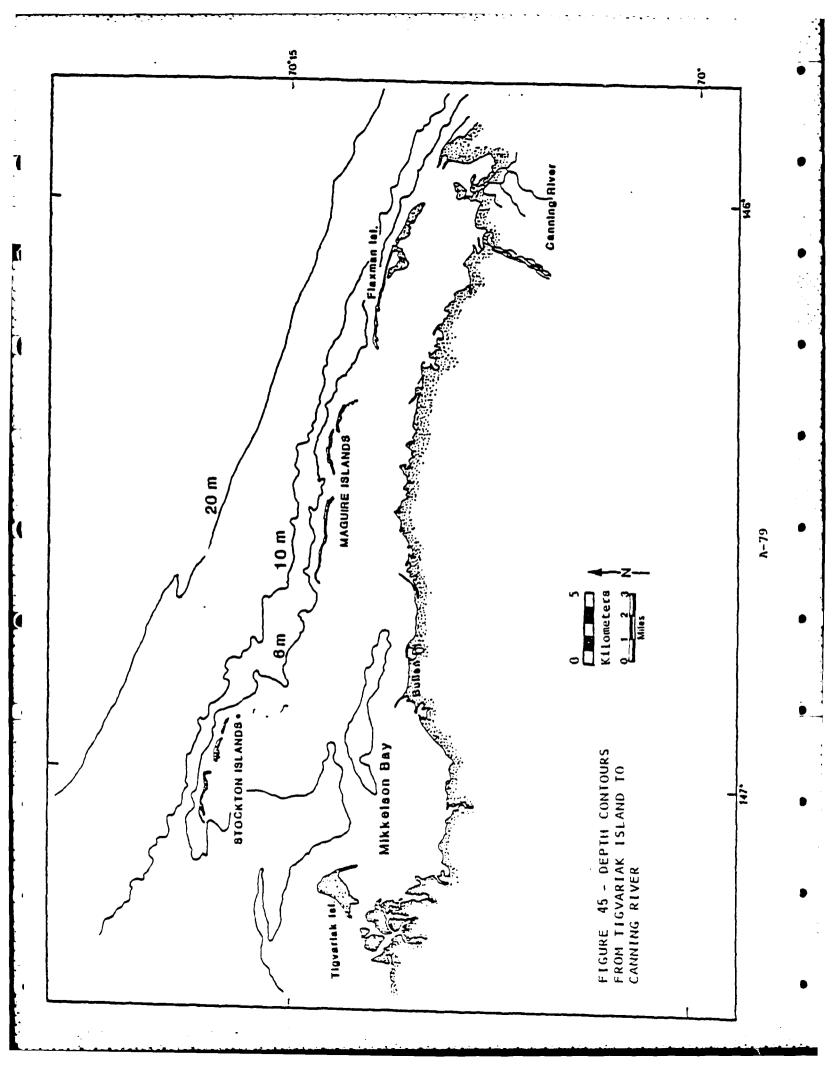


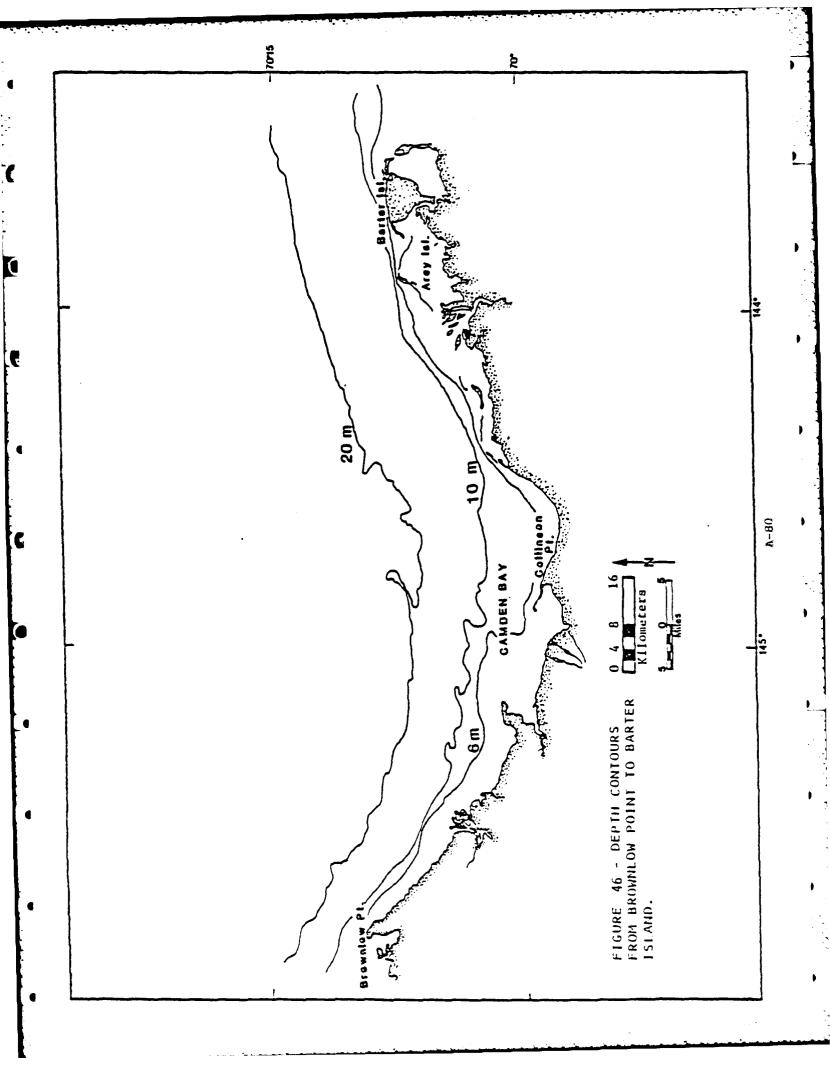


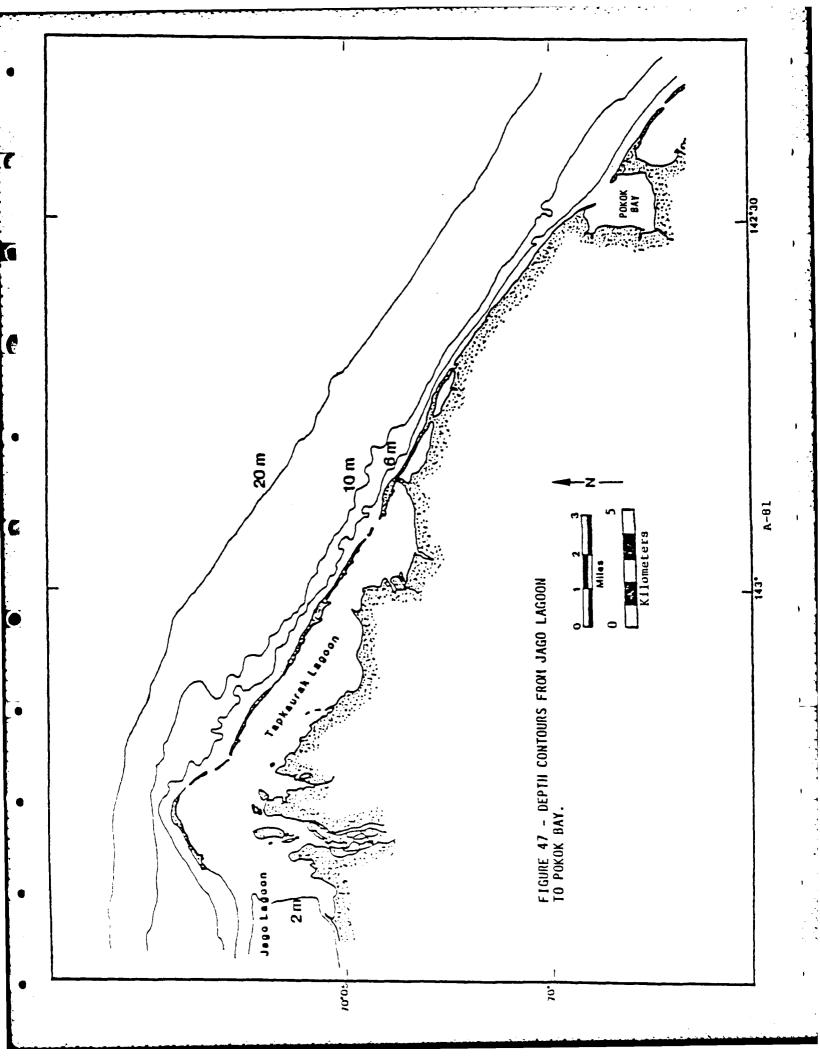


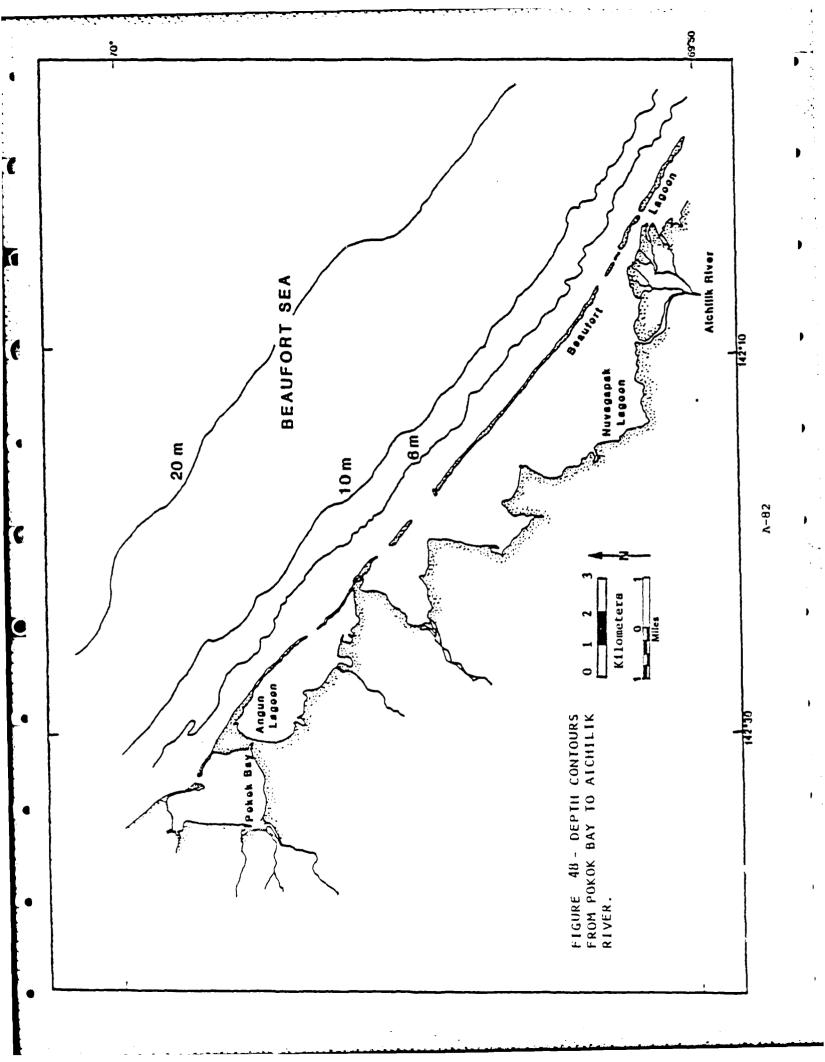


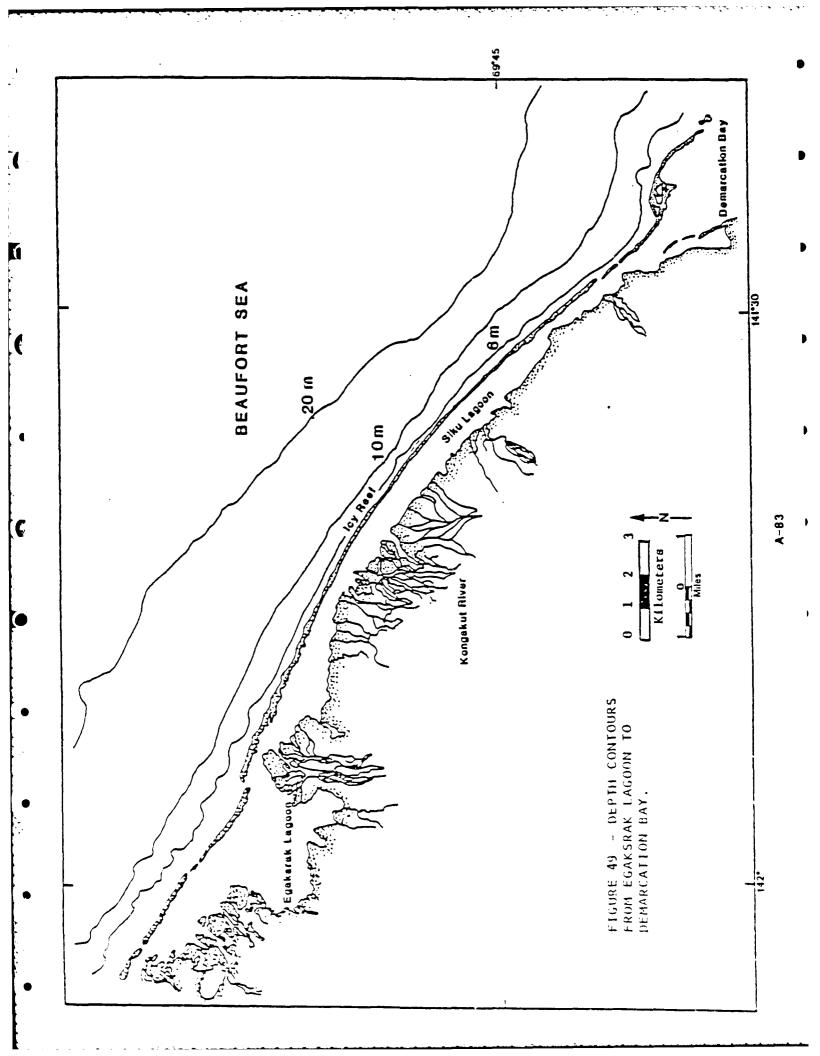












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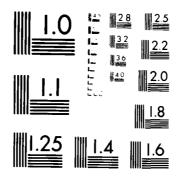
ATLAS OF THE BEAUFORT SEA(U) COAST GUARD RESEARCH AND DEVELOPMENT CENTER GROTON CT I M LISSAUER ET AL.

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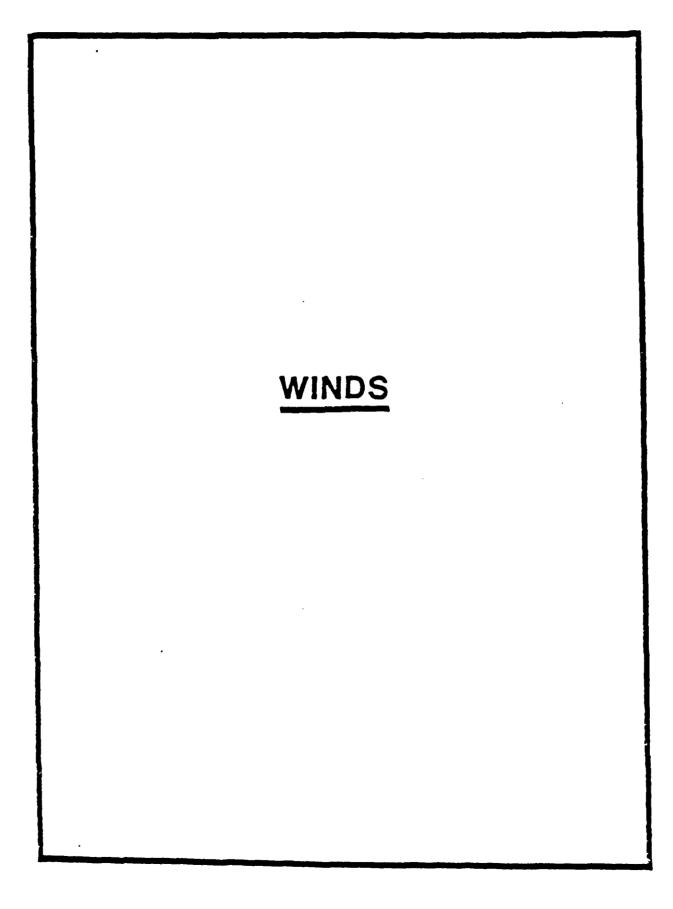
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SECTION B METEOROLOGY



11.0 WINDS

Figure 50 shows the coastal areas of Alaska. Figure 51 presents the major geographic areas along the Beaufort Sea coast. For this entire coastline, which runs several hundred miles, long-term wind data is available from only two sites, Point Barrow and Barter Island. This data collected every three hours is available from 1942 to the present. A source of wind data, over shorter term, is from Prudhoe Bay, Alaska. This data covers the last fifteen years. Other sporadic sources of data are available but because of their nature cannot be used to develop climatological data bases. Hufford et al. (1976) examined wind data from several sources and concluded that the most apparent and consistent feature of the surface winds along the north Alaskan coast is the persistence of easterly winds throughout the year (figure 52). The cause of these predominantly easterly winds is the atmospheric high pressure area centered over the Arctic in the eastern Beaufort Sea. The generalized effect of this atmospheric system can be seen in the prevailing wind directions for the Arctic Basin during both the mid-winter and midsummer months (figures 53 and 54). For both these time periods the prevailing easterlies are apparent along the North Alaskan Coast. The predominant easterly winds are punctuated by westerly winds particularly during the late summer and early fall. These westerly winds are more apparent at Barter Island than at Point Barrow (figure 52). The westerly winds are attributed to low pressure systems moving eastward over the Arctic Ocean (figure 55). Searby and Hunter (1971) analyzed the winds at Barter Island and Point Barrow and provided mean monthly wind speeds for sixteen directions of the compass (tables 4 and 5), and the monthly percentage frequency occurrence of wind directions for sixteen directions of the compass (tables 6 and 7). The strongest winds occur from the east-northeast and west directions during the late summer through the fall when low pressure systems, discussed previously, reach a maximum in intensity and numbers. Table 8 provides percentage frequency of occurrence by speed groups of the wind for each month and for sixteen directions of the compass for Point Barrow and Barter Island. The predominant wind speed group on a monthly basis is 4 to 12 miles per hour (\sim 2 to 6 m/s). The secondary wind speed group is 13 to 24 miles per hour (\sim 6 to 10 m/s). These two speed groups account for approximately 90 percent of the measured winds at Barrow and 80 percent at Barter Island. Table 9 shows the frequency of occurrence of wind speed and direction for Barter Island. Table 10 provides further information on the average winds in the Beaufort Sea for Point Barrow, Prudhoe Bay and Barter Island. The mean wind speeds range between 4.2 and 6.8 meters per second (~ 9 to 15 mph). Thus the energy of the winds on the average is not high. However, more energetic winds can occur as indicated in table 11.

Recent research work by Kozo (1980 and 1982) has investigated mesoscale phenomena along the Beaufort Sea coast which affect the wind field which in turn is responsible for near shore surface water movement. The phenomena studied were mountain barrier baroclinicity and sea breeze forcing. Mountain barrier baroclinicity is a predominantly wintertime phenomenon caused by the Brooks Range. The basic cause of this phenomenon is a piling up of cold air against the Brooks Range. This causes a pressure gradient which is favorable to west winds. The ultimate result is a 180° surface wind shift along the Alaskan Arctic coast between Point Barrow and Barter Island during moderate wind conditions. Such shifts are not explained by the National Weather Service synoptic charts. Kozo (1980) indicates that this phenomenon is a major physical process responsible for the wintertime abundance of westerly

winds from Prudhoe Bay to the east. Likewise the frequency of westerly winds during wintertime is less from Prudhoe Bay west to Point Barrow.

The second mesoscale phenomenon investigated by Kozo (1982) was sea breeze forcing. This phenomenon results in a pressure gradient that accelerates the movement of air from sea to land. Kozo's data indicates that the sea breeze can dominate the surface wind direction at least 25% of the time during the summer in a zone at least 20 km landward and 20 km seaward of the coast.

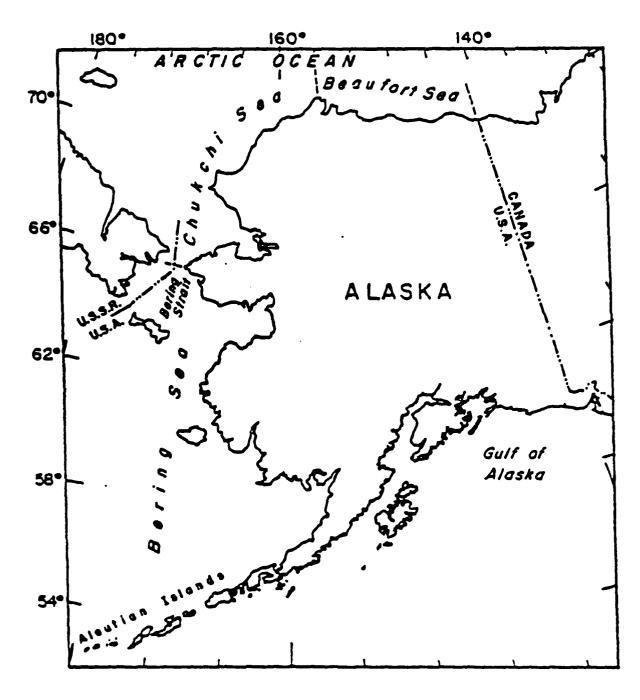


FIGURE 50 - GENERAL LOCATION MAP OF COASTAL AREAS OF ALASKA

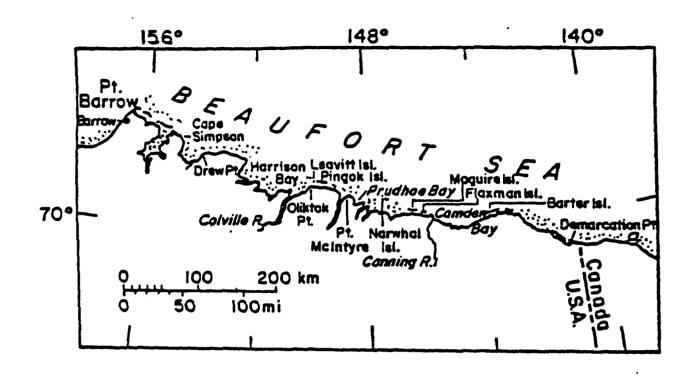
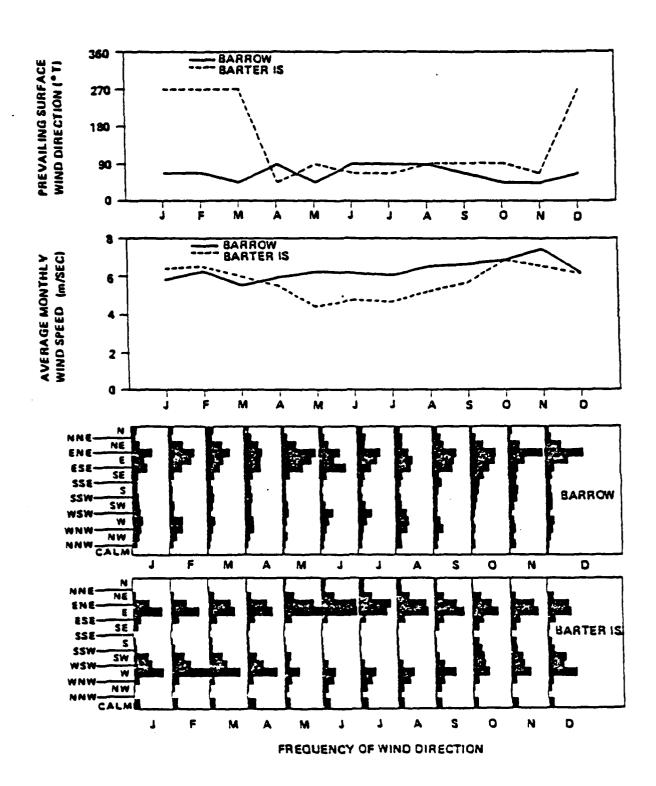


FIGURE 51 - GEOGRAPHIC LOCATIONS ALONG THE BEAUFORT SEA COAST FROM THE U.S.-CANADIAN BORDER



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FIGURE 52 - CHARACTERISTICS OF THE WIND AT POINT BARROW AND BARTER ISLAND (BASED ON 42 YEARS OF DATA)

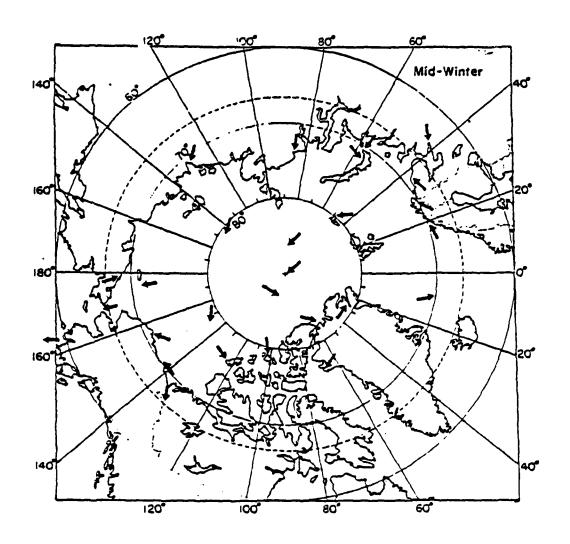


FIGURE 53 - PREVAILING WIND DIRECTIONS IN THE ARCTIC BASIN DURING THE MID-WINTER MONTHS (FROM BILLELO, 1973)

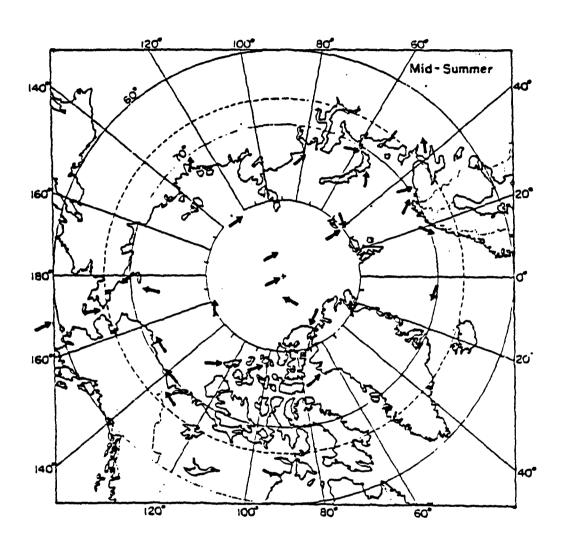


FIGURE 54 - PREVAILING WIND DIRECTIONS IN THE ARCTIC BASIN DURING THE MID-SUMMER MONTHS (FROM BILLELO, 1973)

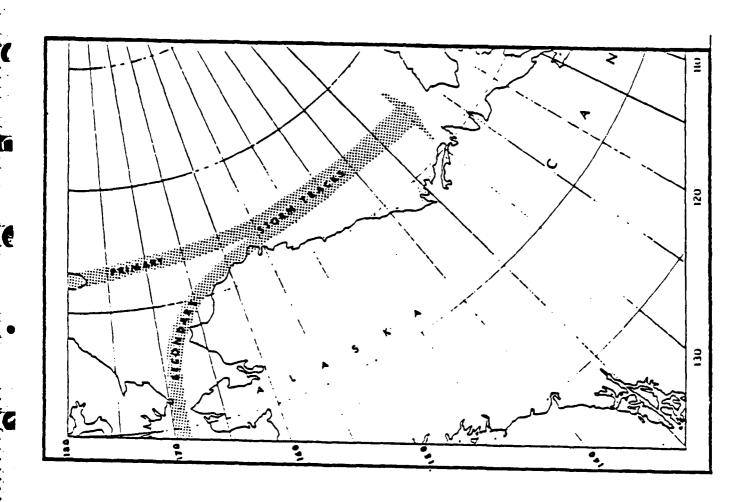


FIGURE 55 - PRIMARY AND SECONDARY STORM TRACKS FOR THE NORTH SLOPE OF ALASKA

Table 4

MEAN MONTHLY WIND SPEED VALUES FOR SIXTEEN DIRECTIONS OF THE COMPASS BARTER ISLAND

BARTER ISLAND WIND SPEED - MEAN VALUES (IN METERS PER SECOND)

	J	F	M	A	M	J	J	A	S	0	N	ם
N	2.7	2.9	2.2	2.4	2.4	2.3	2.5	2.6	3.2	3.8	2.6	2.5
NNE	3.0	2.5	2.9	2.9	2.6	2.7	2.7	2.9	3.2	4.2	3.2	2.9
NE	3.8	3.9	3.9	4.2	4.1	3.4	4.0	4.2	4.9	7.4	5 0	4.2
ENE	6.4	6.0	7.1	6.3	5.8	5.7	5.4	6.4	6.6	9.0	•1	6.6
Ε	6.9	6.6	6.8	6.5	7.0	6.2	5.7	7.0	7.1	10.0	8.7	7.7
ESE	4.5	5.1	5.0	4.9	5.2	5.2	4.7	4.8	5.5	5.9	7	5.0
SE	4.2	3.0	3.3	3.7	3.2	3.7	3.2	3.5	4.4	4.2	3.6	3.3
SSE	5.0	2.9	3.0	3.5	2.6	2.9	3.3	3.0	3.1	3.0	2.8	3.1
S			3.4	3.3	2.6	3.2	3.1	3.2	3.5	3.3	3.4	
SSU	3.2 3.2 2.9 3		3.0	2.8	2.7	3.3	3.3	3.3	3.6	3.6	3.5	
SW .	3.9	3.9	3.8	3.7	3.8	3.3	3.8	4.3	4.1	4.2	4.1	4.4
พรพ	8.0	6.2	5.2	4.9	4.4	5.1	5.6	5.9	6.3	6.9	6.0	6.1
น	9.7	10.0	8.5	7.9	6.3	5.4	5.7	5.7	7.2	8.7	8.1	8.6
นทน	7.7	9.1	7.1	6.9	5.5	4.6	5.0	4.9	5.4	7.1	6.9	7.3
NW	4.0	4.3	3.1	3.5	3.0	3.1	3.2	3.3	3.9	5.2	4.8	4.8
NNW	3.7	2.5	2.8	2.6	2.9	2.6	2.7	2.8	3.2	4.1	2.8	3.0
				MC	NTHLY	AVERAC	E SPEE	מב				
	6.4	6.5	6.0	5.5	4.4	4.7	4.6	5.2	5.6	6.8	6.5	6.1

B-11

Table 5

MEAN MONTHLY WIND SPEED VALUES FOR SIXTEEN DIRECTIONS OF THE COMPASS POINT BARROW

BARROW WIND SPEED - MEAN VALUES (IN METERS PER SECOND)

	J	F	м	A	М	J	J	A	S	0	N	D
N	4.4	3.8	4.6	4.6	3.7	3.8	3.7	4.1	5.6	6.1	6.1	4.2
NNE	4.3	4.1	5.8	5.4	5.3	4.4	4.4	5.3	6.0	6.8	6.1	4.8
NE	5.1	4.9	5.3	5.2	5.6	5.4	5.6	5.3	5.9	6.7	7.3	5.8
ENE	6.4	7.2	6.4	6.4	6.8	6.8	7.4	7.9	7.4	8.0	9.0	7.3
E	6.4	7.5	6.0	6.6	7.4	7.6	6.8	7.9	7.5	8.1	7.5	6.2
ESE	8.1	6.7	6.4	7.2	7.0	7.3	6.9	7.0	7.2	6.3	6.2	5.4
SE	5.5	4.2	5.3	5.4	6.1	5.4	6.0	5.4	5.2	5.2	4.9	4.0
SSE	4.5	4.0	4.6	4.9	3.9	4.7	5.9	5.3	5.6	5.0	5.6	3.5
s	4.2	4.5	4.4	5.2	4.1	4.6	5.7	5.2	4.9	5.9	6.0	4.2
SSW	5.3	6.2	5.2	6.4	4.9	4.6	6.3	6.8	6.1	6.1	7.5	5.9
SM	4.8	6.1	4.8	5.4	5.7	5.4	6.2	5.9	6.3	5.9	7.8	5.4
WSW	6.4	8.1	6.5	6.5	6.1	6.4	6.5	7.7	6.6	8.2	9.7	6.3
W	6.9	6.7	5.6	6.1	5.4	5.2	5.7	6.5	7.5	7.4	8.7	7.7
MNH	5.6	6.4	5.4	6.5	5.6	4.9	4.6	6.3	7.8	6.9	7.3	7.5
NW	4.8	5.5	3.9	5.2	3.8	4.5	4.4	5.8	€.4	5.9	6.9	6.1
NNW	4.9	4.6	4.3	4.2	4.4	4.3	4.1	5.4	6.0	6.4	6.5	4.2
				MO	NTHLY	AVERAC	E SPEE	ED .				
	5.8	6.2	5.5	5.9	6.2	6.1	6.0	6.5	6.6	6.8	7.4	6.1

Table 6

MONTHLY PERCENTAGE FREQUENCY OF OCCURRENCE OF WIND DIRECTION FOR SIXTEEN DIRECTIONS OF THE COMPASS - BARTER ISLAND

BARTER ISLAND

WIND DIRECTION (PERCENTAGE FREQUENCY OF OCCURRENCE)

	J	F	М	A	М	ַ ד	J	<u> </u>	S	0	N	<u>D</u>
N	.4	.4	.5	. 8	1.8	2.6	2.8	2.5	1.9	.9	.6	.9
NNE	.5	. 3	.7	.9	1.5	2.3	3.1	3.0	2.2	1.2	1.0	.5
NE	2.6	2.3	4.3	2.9	7.2	8.7	9.4	8.1	6.6	6.8	4.3	3.9
ENE	10.8	10.6	15.2	13.9	20.1	23.6	21.2	19.8	15.0	13.0	16.2	14.5
E	19.8	18.5	17.5	17.6	29.4	24.3	18.2	20.7	20.0	17.3	18.9	15.7
ESE	5.3	5.3	5.1	5.7	5.1	3.0	4.5	5.7	6.9	7.2	6.7	5.0
SE	3.9	1.1	1.7	2.1	1.1	٠,9	1.5	1.8	3.0	4.1	3.6	1.6
SSE	.9	.6	.4	.8	.5	.5	.7	.9	1.0	1.2	1.0	.7
s	1.9	1.7	1.5	1.9	.9	.7	1.5	1.4	1.8	4.3	3.3	2.1
SSW	2.2	2.4	2.2	2.0	1.2	.6	7	1.2	1.7	6.1	4.5	4.2
SW	10.3	8.9	9.2	7.0	4.6	1.3	1.4	1.8	3.1	7.4	8.0	12.0
WSW	12.4	13.3	12.2	11.6	3.9	3.7	3.6	3.7	6.4	8.1	9.4	10.7
W	<u> 19.1</u>	24.9	21.3	20.5	10.2	9.3	11.0	11.2	13.6	11.8	13.4	19.5
WNW	4.0	4.8	3.8	4.4	5.2	7.7	7.9	8.9	6.4	4.2	3.0	2.9
NW	1.1	.9	.9	1.4	2.6	4.3	5.0	4.5	4.0	1.9	1.1	1.7
NNW	.2	.1	. 3	. 4	1.3	1.8	2.0	2.2	1.7	.8	.4	. 3
CALM	4.7	4.0	4.1	6.1	3.4	4.8	5.5	3.7	4.6	3.7	4.6	3.8

PREVAILING DIRECTION UNDERLINED

TABLE 7

MONTHLY PERCENTAGE FREQUENCY OF OCCURRENCE OF WIND DIRECTION FOR SIXTEEN DIRECTIONS OF THE COMPASS - POINT BARROW

BARROW

WIND DIRECTION (PERCENTAGE FREQUENCY OF OCCURRENCE)

	J	F	M	<u> </u>	14	J	J	A	<u>s</u>	0	N	<u>D</u>
N	3.9	3.4	5.0	3.7	2.1	2.5	4.3	4.3	3.7	2.0	3.1	2.6
NNE	3.8	2.5	4.8	5.3	3.9	3.6	6.1	5.5	5.9	3.7	2.0	3.9
NE	6.2	10.1	11.9	9.4	9.9	6.5	5.8	5.4	7.4	8.3	7.8	11.2
ENE	14.3	17.4	16.0	12.3	22.7	13,4	15,7	9.4	16.2	17.2	23.4	27.7
_ <u>E</u>	8.2	15.0	12.9	11.4	19.7	10.2	12.9	12.0	14.3	16.8	11.7	15.2
ESE	11.4	7.9	9.1	11.3	13.6	17.4	9.3	11.4	13.9	12.0	7.1	6.3
SE	5.4	3.7	5.4	6.0	5.5	5.2	5.3	4.8	4.7	8.6	5.9	4.3
SSE	3.1	1.9	2.4	3.7	2.1	2.2	3.6	4.1	5.3	6.5	5.6	2.2
s	3.1	2.3	4.1	4.6	2.1	1.7	2.5	2.8	3.1	5.5	6.4	2.8
SSW	5.2	2.3	4.7	5.0	2.3	2.0	3.2	3.8	2.8	4.3	6.0	3.0
SW	5.4	3.1	3.9	4.4	2.9	3.1	5.2	4.1	2.8	3.0	3.4	3.1
WSW	5.9	2.9	3.3	4.5	3.4	8.5	9.5	7.7	2.7	2.6	3.1	2.7
<u>W</u> _	7.6	9.5	5.7	5.7	2.4	6.4	6.5	8.9	4.6	2.2	3.7	4.9
WNW	6.8	9.3	3.9	5.7	1.9	4.5	3.7	7.7	5.8	2.6	4.6	4.6
NW	4.8	5.3	3.3	3.6	2.6	3.6	2.9	3.4	3.1	1.7	3.2	2.4
NNW	3.8	2.4	2.4	2.4	1.6	2.2	2.5	4.1	2.7	2.3	4.5	1.6
Calm	1.2	0.8	1.2	1.0	1.1	0.9	1.1	0.7	1.0	0.6	0.4	1.5

PREVAILING DIRECTION UNDERLINED

TABLE 8
WIND-PERCENTAGE FREQUENCY OF OCCURRENCE BY SPEED GROUPS
FOR POINT BARROW AND BARTER ISLAND BY MONTH

EXTREME MONTHLY WIND SPEEDS FOR BOTH LOCATIONS BY MONTH (1 mph = 0.4470416 m/s)

		- מאנו	PERCE	NIAGE	FREQU	ENCY	OF OC	URREN	CE BY	57550	GROUP	5		 	E	XTR	FIF
	i			BARRO						BARTE	r isu	WD			WIN	DS	(MPH)
	Calm	1 to 3	4 to 12 HPII	13 to 24	25 to 31 MPN	32 to 46 HPII	47 HPII and over	Calm	1 to 3 HPII	4 to 12 HPH	13 to 24 PPH	25 to 31 NPH	32 to 46	47 HPH and over	BARRON	UHIAT	BARTER ISLAND
Jan	1.3	6.1	59.6	27.5	4.0	1.4	0,1	4.7	6.4	44.0	29.2	7.4	6.8	1.4	56		75
Feb	0.8	7.5	53.4	31.2	4.7	2.3	1.0	4.0	5.3	43.2	32.8	6.9	6.2	1.7	58		65
Mar	1.2	5.8	60.3	30.9	1.7	0.1	٥	4.7	6.7	44.9	32,8	6.5	4.2	0.8	58	<u>بر</u>	77
Apr	1.0	4.5	57.2	33.1	.3.8	0.3	a	6.1	6.4	47.0	31.2	5.6	3.5	0.2	52	A B	52
May	1.1	3.1	52.3	47.9	1.5	0.1	0	3.4	6.4	48.3	36.7	3.6	1.6	a	43	1	55
Jun	0.9	2.5	54.9	40.4	1.3	0	٥	4.8	5.2	53.9	33.1	1.6	0.3	0	38	Y A	38
Jul	1.0	3.6	52.5	42.0	0.9	0	0	5.5	5.8	56.6	30.6	1.3	0.2	0	56	<	40
Aug	0.6	2.7	53.6	37.7	4.2	1.2	0	3.7	5.5	52.8	32.8	3.7	1.5	0.1	47	0 1	46
Sep	1.0	3.1	46.1	46.4	3.1	0.3	0	4.5	5.9	47.4	34.3	5.1	2.2	0.4	56	=	78
0ct	0.8	2.7	49.2	40.6	5.5	1.2	0	3.7	4.5	40.2	33.9	9.8	7.5	0.4	55	1 1	58
Nov	0.3	3.8	43.4	42.2	6.9	3.2	0.2	4.5	5.9	42.0	37.0	9.0	6.5	0.8	63	VO	67
Dec	7.4	6.9	50.4	38.0	2.8	0.5	0	3.0	6.8	45.9	31.6	6.9	4.0	1.0	70		75
Yr.	0.9	4.4	52.7	37.7	3.4	0.9	•	4.4	6.0	47.2	32.5	5.6	3.7	0.6	70		73

Table 9

PERCENTAGE FREQUENCY OF OCCURRENCE OF WIND DIRECTION AND SPEED FOR BARTER ISLAND (FROM ISAKSON, ET AL., 1975)
(1 mph = 0.4470416 m/s)

PERCENTAGE FREQUENCY OF WIND DIRECTION AND SPEED (MPH) (FROM HOURLY OBSERVATIONS)

BARTER ISLAND

SURFACE WINDS

=	(FROM 110	HOURLY O	UBSE KVAT TONS	LIONS									
	1-3	9-4	7-10	7-10 11-16 17-21	17-21	22-27	28-33	34-40	41-47	48-55	2.56	×	Mean Wind Speed
Z	7.	٠,	.3	0.	0.	0.						1.4	5.1
NNE	.3	.,	4.	.1	0.							1.4	5.7
NE	.5	1.6	1.9	8.	.2	.1	0.	0.	0.			5.2	8.6
ENE	.4	1.9	4.6	5.0	2.0	6.	.3	.1	0.	0.		15.2	12.3
E	.,	2.5	5.1	9.9	3.2	1.9	8.	.3	.1	0.		21.2	13.7
ESE	7.	1.3	1.9	1.2	4.	.2	.1	0.	0.			5.5	10.3
SE		.7	.,	.2	0.	0.	0.	0.	0.			2.0	7.2
SSE	.2	.3	.3	0.	0.	0.	0.					.8	6.2
S	.3	8.	8.	.1	0.							2.0	6.4
SSW	.3	.8	1.0	.1	0.	0.						2.3	6.5
SW	٠.	1.7	2.7	.,	.1	0.	0.	0.	0.	0.	0.	5.8	7.7
MSM	.3	1.3	3.4	1.9	.5	.3	.2	.1		0.	0.	8.1	11.6
3	7.	1.5	3.9	4.7	2.4	7.1	1.0	3	4	4	0,	16.4	15.5
NA.	.3	.9	1.6	1.3	9.	.3		0.	0.	0.	0.	5.2	11.7
NW	.4	1.0	.,	.3	1.	0.	0.	0.		0.	0.	2.4	7.1
NNM	7.	9.	.2	0.	0.	0.						1.0	5.6
ARBL													
CALM	X	X	X	X	X	X	X	X	X	X	X	4.1	
	0.9	18.3	29.4	23.0	9.6	5.5	2.5	1.2	.3	.2	.1	100.0	11.2

162511

TOTAL NUMBER OF OBSERVATIONS

B-16

a) AVERAGE WINDS IN THE BEAUFORT SEA AREA BY MONTH FOR POINT BARROW, PRUDHOE BAY AND BARTER ISLAND

	Sai	TOW	Prudh	ое вау	Barter	Island
ilonth	Mean Speed (m/s)	Prevailing Direction	Mean Speed (m/s)	Prevailing Direction	Mean Speed (m/s)	Prevailing Direction
January	5.1	ESE	6.8	¥	6.6	¥
February	4.9	ε	5.9	¥	6.3	¥
March	5.0	ENE	5.4	W	6.1	W
April	5.1	NE	5.3	Ε	5.3	¥
May	5.2	ENE	5,2	Ε	5.6	Ε
June	5.1	Ε.	4.9	ε	5.1	ENE
July	5.2	Ε	4.7	Ε	4.7	ENE
August	5.5	Ε	4.2	E	5.2	ε
September	5.8	£	5.7	¥	5.9	ε
October	5.9	E	5.2	W	6.5	E
November	5.5	E	6.4	¥	6.7	Ε
December	5.0	Ε	5.4	W	6.2	ε
Annual	5.3	E	5.4	ε	5.9	Ε

b) EXTREME (ONE-MINUTE) WINDS IN THE BEAUFORT SEA AREA BY MONTH FOR POINT BARROW, PRUDHOE BAY AND BARTER ISLAND

	Bai	rrow '	Prud	hoe Bay	Barter	Island
Honth	fean Speed (m/s)	Direction	Speed (m/s)	Direction	Speed (m/s)	Direction
January	22	Ε	23	W	36	W
February	21	Ε	17	W	28	W
March	26	w	17	EZE	34	¥
April 1	18	NSN	16	W	23	u
May	17	MSM	15	Ε	25	¥
June	16	271	20	£	17	w
July	16	271	14	WNW	18	MZM
August	16	SM .	22	W	20	¥
September	20	NM	21	E	35	н
October	25	W	23	ENE	26	¥
November	24	¥	18	٤	36	ч
December	25	MSM	26	W	32	w
Maximum	26	¥	26	¥	36	¥

(FROM U.S. MATIONAL CENTER AND U.S. MATIONAL WEATHER SERVICE RECORDS)

Table 11

WIND SPEED DATA BY HONTH FOR POINT BARROW AND BARTER ISLAND (FROM BILLELO, 1973)

Mean windspeed (m/s) and direction, Northern Hemisphere

	٦	Ŀ	X	<	Σ	C	A		S	0	2	Q	Year
Barrow													
mean speed (30) prevailing direction (7) fastest mile (30) direction (9)	5.7 ESE 56	5.8 ENE 51 SW	5.6 NE 48	5.9 E 52	6.1 NE 43 SW	5.9 38	6.1 SW 41	6.5 E 47	7.1 ENE 56 W	7.2 NE 51	6.4 NE 63	5.6 ENE 70	6.2 NE 70
Barter Island													;
mean speed (7) prevailing direction (7)	7.6	7.8	6.0	6.5 ENE	6.1 E	5.7 ENE	5.3 ENE	5.7 E	6.6 F	7.6 E	7.7 ENE	6.1 ENE	6.5 ENE

() Number of years of record

STORM SURGES

12.0 STORM SURGES

Storm surges may occur with little or no warning. They are caused by the wind pushing water towards the coastline thus causing a rise in sea level at the coast. This rise in sea level can cause severe flooding along a coastline. A brief description of storm surges was presented in section A. This section provides a more detailed study of storm surge events and a practical method for estimating how high the sea level might rise and therefore how severe the flooding might be.

Surges are highest when there is little or no ice (open water) over the ocean. Intensive storms can produce 12-foot (3.7 meters) surges over almost any portion of the western and northern Alaska coasts. This seems to be the upper limit of surges in most places. In ice-covered seas the surge is usually less than 3 feet (~ 1 meter). A hazard associated with ice-covered surges is the flooding of shorefast or bottomfast ice as the rising water comes through the cracks in the ice. Another hazard is ice push-up, which occurs when floating ice rises above shorefast ice and is pushed on-shore by wind or current forces.

Finally, there is the condition were the sea is ice covered, but the ice is relatively thin or unconsolidated. Some storms are capable of obliterating thin ice and thus creating an open water condition. Surges have exceeded 4 feet with these situations.

Autumn is the season for most storm surge flooding along the Beaufort Sea coast. The most likely storm for a surge is one moving from west to east well offshore with the surge being caused by the west to northwest winds in the southwest quadrant of the storm.

The two worst recorded cases of coastal flooding were caused by fall storms. One, in October 1963 (Section E, Case #4), had a surge of 12 feet (3.7 meters) in Barrow and lesser surges from Point Lay to Barter Island. The other, in September 1970 (Section E, Case #6), was judged to be as high as any previous storm as determined by driftwood lines. The elevation of the highest driftwood line varied from 4.5 to 11.2 feet (1.4 to 3.4 meters) above sea level. The variation of height is partially due to differences in exposure. Since the onshore winds during the storm were from the west, eastern shores of stream mouths, bays, etc., had higher surges than western shores. Figure 56 shows the paths of meteorological systems which have caused large storm surge occurrences. Documentation of some of these events as well as others is provided by Wise et al., 1981 (Section E).

12.1 RAPID MANUAL FORECAST FOR THE BEAUFORT SEA COAST

- 1. Favorable wind direction for surge formation measuring clockwise: 270 to 020 (°T).
- 2. If the wind is from a favorable direction determine the following from meteorological forecasts:
- a. The number of hours the coastal area will be subjected to the favorable winds. The wind direction and speed should be reasonably constant and not vary past the following limits:
 - (1) The wind direction or orientation of the isobars does not change direction at a rate greater than 15° per 180 nautical miles and the total changes does not exceed 30°.
 - (2) The wind speed does not vary more than 20 percent from the average wind speed in the area of the direction fetch being considered. Example: average wind is 40; acceptable range is 32 to 48.
- 3. Using the wind speed compute the surge height from Figure 55. The surge height is then adjusted for duration of wind speed, ice cover and barometric pressure.
 - a. If wind speed duration is less than:
 - (1) 3 hours reduce surge by 60 percent
 - (2) 6 hours reduce surge by 40 percent
 - (3) 9 hours reduce surge by 20 percent
 - (4) 12 hours reduce surge by 10 percent
 - (5) 12+ hours no reduction
 - b. If ice cover is less than:
 - (1) 1.5 tenths no reduction
 - (2) 3.0 tenths reduce surge by 20 percent (cumulative to above)
 - (3) 5.0 tenths reduce surge by 50 percent (cumulative)
 - (4) 10.0 tenths reduce surge by 75 percent (cumulative)
 - (5) Surges to 3 feet with 10 tenths ice cover have been reported with ice to 3 feet thick between October and January.
- c. Raise the surge height one foot for every 30 mb pressure increment below 1000 mb in surge area.

The above surge prediction method contains some subjectively derived information. Therefore it should not be used to predict an exact value of the surge height, but rather, a general estimate of the potential magnitude of a surge event.

4. Example

A possible surge condition is developing in the Prudhoe Bay region. Predicted wind is 290° T at 38 knots. Using Figure 57 the surge height is predicted to be 10 feet. Wind duration is forecast to be five hours. Reduce surge by 40 percent (10 -4 = 6). Ice cover is 4 tenths. Reduce surge by 50 percent (6 -3 = 3). Lowest pressure coincident with surge is 960 mb. Raise surge height 1.3 feet (3 +1.3 = 4.3). Estimated surge height is therefore predicted to be between 4 and 5 feet.

(1 foot = 0.3048006 m) (1 nautical mile = 1.85326 Km)

- 1. 2-3 Oct, 1963, Barrow 2. 5-7 Oct 1963, Barrow
- 3. 15-17 Nov 1966, Barrow
- 4. 20-25 Sep 1968, Barrow
- 5. 26-27 Aug 1975, Icy Cape 6. 17 Aug 1975, Icy Cape, negative surge 7. 8-12 Oct 1972, Point Lay 8. 25-26 Oct 1969, Barrow

Two additional surges not on the map: one with a low center 177° $178^{\circ}N$, and another caused by a large persistent high pressure area centered north and east of the area.

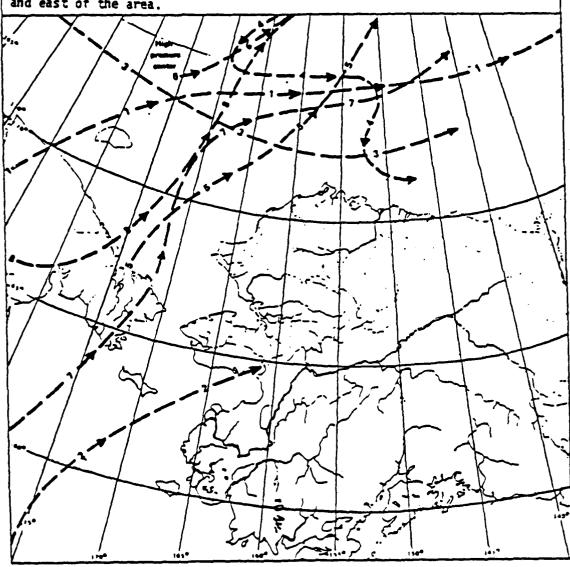
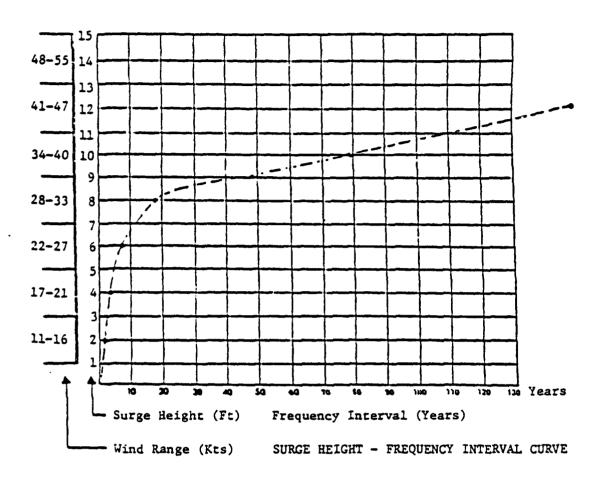


FIGURE 56 - PATHS OF METEOROLOGICAL SYSTEMS THAT HAVE CAUSED STORM SURGES (FROM BROWER, ET AL., 1977)



- (1 foot = 0.3048006 m)
- (1 knot = 0.5148 m/s)

FIGURE 57 - CURVE OF SURGE HEIGHT VS FREQUENCY OF OCCURRENCE FOR STORM SURGES IN THE BEAUFORT SEA

WAVES

13.0 WAVES

Wave action along the North Alaskan coast is inhibited by the presence of the pack ice and floating ice during periods of open water. The pack ice during the summer and fall months may vary from only a few kilometers to 100 kilometers offshore depending on meteorologic conditions. The pack ice reduces the open water fetch and floating ice on open water areas reduces the development of waves. The surface winds determine the prevailing direction of the seas.

When ice conditions provide an appreciable fetch, the generally small sea height conditions can be surpassed (tables 12-14). For the Western Beaufort Sea area waves of greater than 30 feet (9.1 meters) were recorded for winds from the east. Other wave occurrences of more than ten feet are indicated for all of the areas along the coast. However, for all these areas the tabulations of sea height versus direction indicate that over 85% of the time sea height is less than 3 feet (\sim 1 meter). For the nearshore area of Pingok Island, Wiseman, et al. (1979), found that the most common nearshore wave field had an energy peak at periods between 2 and 3 seconds and a significant wave height of 20 to 30 cm. These nearshore conditions are generally true for the entire Beaufort Sea coast.

Tables 15-17 provide further information on sea height conditions in the Beaufort Sea. For the western, central and eastern areas of the coast, sea height values are provided for different wind speed groups for eight directions of the compass.

Table 18 provides estimates of maximum sustained winds, maximum significant waves, and extreme waves for various selected return periods.

Table 12

SEA HEIGHT (FT) VS DIRECTION FROM 11 JULY TO 31 OCTOBER FOR THE CENTRAL BEAUFORT SEA COAST (FROM GATTO, 1980)

(1 foot = 0.3048006 m)

SEA HEIGHT

	<3	3	5	6.5	8	9.5	11	13	14	16	17-29	30+	TOTAL	%
N	2												2	0.2
MME	4	4											8	1.0
NE		1							1	\			9	1.1
ENE	32	2	2							·			34	4,4
E	68	23	, 2							ł			94	11,5
ESE	25	•								J			34	4.2
SE	4	3											7	0,7
SSE		۵							1	1			14	1.7
S	5	1											•	0,7
S SSW	3	•		2									5	0.6
SW	7									T			7	0,9
wsw	2	3											5	0.4
w	34	7	4										45	5.5
WNW	34	10	3	4		1							43	7.7
NW	37	6		3	1		1						48	5.9
NNW	5			2					1.				7	0.7
CALM	428					1		1					428	52.3
C -15														
C +15									7					
TOTAL	704	77	'12	27	1	Ī	1	1					878	1
%	84-3	9-4	1-5	2-4	0-1		0-1			\top				100

SEA HEIGHT (FT) VS DIRECTION FROM 11 JULY TO 31 OCTOBER FOR THE EASTERN BEAUFORT SEA COAST (FROM GATTO, 1980) (1 foot = 0.3048006 m)

6

SEA HEIGHT

	< 3	3	5	6.5	8	9.5	1.1	13	14	16	17-29	30+	TOTAL	%_
N	4	1						1	1		1		5	0.5
NNE	4		1	1					1		1		4	0.4
NE	24	1	2										34	1.2
ENE	48	7	3			1							48	4.4
ε	57	45	9										111	10.4
ESE	16	3											19	1.1
SE	4	•	2						1				15	1.4
322	1												1	0.1
S SSW	3								3			<u> </u>	1 • 1	٥
SSW	2			<u> </u>				1	<u> </u>				2	0.
Sw	8	1											1 1	٥
WSW	6												•	٥
W	33	•	2	2									43	4,
MMM	21	7								1	1		28	2
NW	14	5								1			19	1.
NNW	2		i										2	٥
CALM	494								1.				494	45
C -15														
C +15			1										1	
TOTAL	933	91	13	1 2	}	1			4				1048	
%	89.2	1.5	1.7	0.2	T				0.4	ī				100

Table 14

SEA HEIGHT (FT) VS DIRECTION FROM 1 JULY to 20 NOVEMBER FOR THE WESTERN BEAUFORT SEA COAST

(FROM GATTO, 1980)

(1 foot = 0.3048006 m)

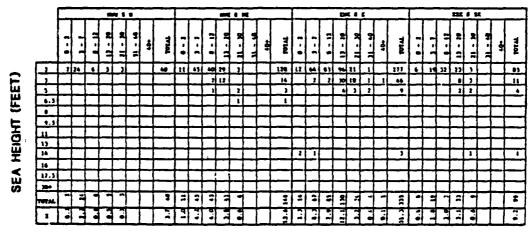
-			_	_								_	_	_		_	,	-	_					
	*	3:1	7.4	•	7.7			7	0.0	0.0	0.7	0.5	0.7	1.3	6.	1.7	2.0	1		20.				
	TOTAL	3	57	119	150				2	20	7	2	=	37	94	=	*			2/6			2415	
	ġ					ľ	•																2	-0
	17-29						1																	
	91												,											
	•															-	·						-	0.0
_	13																							
SEA HEIGHT	•			-	·																		-	0.0
SE	9.5				1																			
	•		1	1			~								•	•	-						^	0.0
	6 9		Ī						-						•	•	^	•					2	10
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	~	=			-	3.0	3.6	•	-	~	-	7		1	•	2	=	02	•				141	
	3	1		=	0	Ξ	121	:	5-	=	-	٩		= =			2	28	3.6	1370			21.74	
		1	2	NA.	N.	ENE	_	ESE	SÉ	556		. 3		N C	RCR	≥	FIR	3	MNN	CALM	0 - 15	414	101	
	_	<u> </u>		<u> </u>		_			_		NOI	133	32416	a 1	735									

TABLE 15

WIND SPEED AND DIRECTION YS. SEA HEIGHT FROM 11 JULY to 31 OCTOBER FOR THE CENTRAL BEAUFORT SEA COAST (FROM GATTO, 1980)

(1 knot = 0.5148 m/s)(1 foot = 0.3048006 m)

WIND SPEED (KNOTS) AND DIRECTION



	Г		_	-	SE				Т			_	101		,			Γ	_		YEN	•							-		~				
	1:0	1.1	11 . 0	N - 11		3 . 2	*	17 84	!		1	4:4	13 - 10	2	. \$	ģ	3 2	-	1 - 1	8 - 12	11 - 10	X - =	31 - 16	*6*	1014	7 - 0		11 - 1		0 - 11	•	ģ	TOTAL	17101	-
F	~	70	•	7				7	,	•	15	12	Ţ	•			4	79	41	9	75	16			221	3	*	24	• >	•	7		120	936	en.
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	3	3	Ē	3					:†	1	3	3	-		Г		3	1:	3		1.0	-			-		7	-							8

TABLE 16

WIND SPEED AND DIRECTION YS. SEA HEIGHT FROM 11 JULY to 31 OCTOBER FROM THE EASTERN BEAUFORT SEA COAST (FROM GATTO, 1980)

(1 knot = 0.5148 m/s)(1 foot = 0.3048006 m)

WIND SPEED (KNOTS) AND DIRECTION

- 1	_	_	_	_	N. 1				\bot					K						_	-							u	. 36			
	1 - 0	1:1	11 - 0	13 - 70	21 - 30	27 . 12	404	TATAL	-	1.1	6 - 12	13 - 19	•	21 - 45	•	77201	• . 1	1 - 1	4 : 11	11:18	21 - 36	3 - 1	\$	POTAL	1.1	1:1	0 - 11	11 - 10	21 - X	21 - 19	100	
7		10	•	13	-			47		16	13	17				*	,	•	73	.,	14	7	┪	74	-	10	70	33	3	T	7	i
			1							1	3	-							7	19	٠,	71		11		\neg	7	٠,	٦,	\neg		i
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6.5																							\neg		_	_		\neg	╗	╗		Ì
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TABLE 17

WIND SPEED AND DIRECTION VS. SEA HEIGHT FROM 1 JULY TO 20 NOVEMBER FOR THE WESTERN BEAUFORT SEA COAST (FROM GATTO, 1980)

(1 knor = 0.5148 m/s)(1 foor = 0.3048006 m)

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TABLE 18

ANNUAL MAXIMUM WINDS AND WAVES FOR SELECTED RETURN PERIODS

Return periods for maximum sustained winds and for maximum significant and extreme wave heights are presented in tabular form for the Beaufort Sea coast. Sustained winds are winds averaged over a period of one minute; the significant wave height is the average height of the highest one third of all waves (sea and swell) in view and the extreme wave height is an empirical estimate of 1.8 times the significant wave height. The return period (years) is statistically derived from empirical observations. It estimates how often an extreme event may occur in a particular area. For example, on the average the Beaufort coast can expect annual maximum sustained wind speed to exceed 81 knots once in 100 years.

Return period years	laximum sustained wind-knots	Maximum significant wave-meters (feet)	Extreme wave meters (feet)
5	57	10.0 (33)	18.0 (59)
10	62	11.0 (37)	20.5 (67)
25	69	13.0 (43)	24.0 (78)
50	75	15.0 (49)	27.0 (88)
100	81	17.0 (55)	30.0 (99)

SECTION C

14.0. ICE ZONES

Three ice zones can be broadly defined for the southern Beaufort Sea.

- 1. Fast ice zone consists of seasonal ice which is an extension of the land, because it generally remains immobile during the winter. Its extension seaward varies but typically progresses to the 20 meter depth contour by late winter. The ice is generally two meters thick.
- 2. Seasonal pack ice zone begins at the edge of the fast ice and continues out 100 to 200 km. There are often strong shear forces within this ice zone. It is mobile and often contains a large percentage of first year ice. In this zone, the undeformed areas of ice have an average thickness of approximately 2 meters.
- 3. Polar pack ice zone mainly composed of thick multi-year floes. This zone generally lies beyond the continental shelf. Approximate ice thickness of the general terrain of old multi-year floes is 2 to 4 meters.

SHORE FAST ICE C-3

15.1 ANNUAL ICE CYCLE WITHIN THE NEARSHORE AREA OF THE BEAUFORT SEA (AVERAGE CONDITIONS)

Time

Event

Late September to early October

New ice begins to form in open water. Ice forms first adjacent to rivers and in coastal lagoons. (During severe ice seasons this process can begin as early as late August.)

Mid to late October

The landfast ice sheet has formed and is continuous. Areas of the ice are unstable because of the thinness of the ice (particularly north of the barrier islands).

November to February

The landfast ice area has thickened and become more stable, particularly inside the barrier islands. Some modifications of the ice sheet occur due to ridging, incursions of older pack ice north of the barrier islands, and grounding of ice masses as they are driven ashore by the winds.

March to May

Ice has generally reached its greatest thickness. Ice has its most stable period during this time.

Late May to early June

Warming trend causes breakup of rivers and overflooding of ice by the rivers in the nearshore zone.

Early to late June

Melt ponds begin to form on the ice. The ice begins to lose thickness and therefore weakens. Toward the end of the period open water areas begin to occur. These areas are generally along the coast and around the barrier islands, particularly the southern sides. Cracks in the ice can be found both north and south of the barrier islands.

June to August

Breakup of ice sheet continues. Significant open water occurs within the barrier islands by mid to late July.

August to September

Generally open water within the barrier islands. Some ice masses remain within the barrier island during severe ice years. Open water north of the barrier island is dependent on the north-south migration of the pack ice. Refreezing occurs.

15.2 FREEZING AND BREAKUP OF NEARSHORE ICE

Table 19 provides data on freezing and breakup of the nearshore ice field at Point Barrow and Barter Island. As can be seen, yearly freezeup may occur at Barter Island at any time between 20 September and 25 October and at Point Barrow at any time between 23 September and 19 December. Similarly, breakup has occurred at Point Barrow as early as 15 June and as late as 22 August, and at Barter Island from 22 July to 14 August. Freezeup and breakup are at present unpredictable parameters.

FREEZEUP AND BREAKUP DATES OF FAST ICE FOR BARTER ISLAND AND POINT BARROW Table 19

	ars	Date		
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•		Average		7/17-23
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	Familian	1621	1122	6/15
	Vears		9	56
	Average		10/5	10/1-5
FREEZEUP	Latest		10/25	12/19
	Earliest Latest		9/20	9/3
	STATIONS	BEAUFORT SEA:	Barter Island	Point Barrow

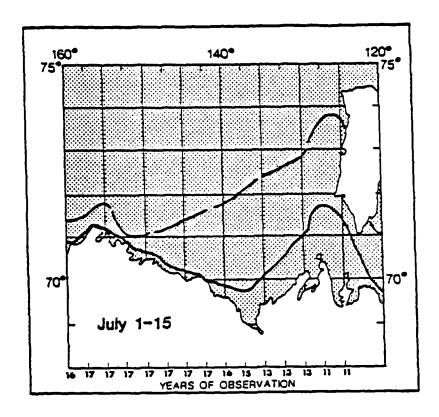
PACK ICE	

16.0 PACK ICE

In addition to the unpredictability of freeze-up and breakup is the erratic movement of the pack ice. The pack ice can be driven toward the coast at any time by a strong onshore wind. During severe ice seasons, i.e., 1975, ships can be prevented from passing east or west along the Beaufort Sea coast. During less severe seasons, the ice edge can retreat more than 50 kilometers offshore leaving a wide expanse of open water along the entire Beaufort Sea coast. Because of the sporadic nature of ice movement and growth "average ice conditions" are difficult to ascertain. However, Brower et al. (1977) prepared a comprehensive climatic atlas of the outer continental shelf waters and coastal regions of Alaska. From their work, diagrams of the extreme northern latitude and extreme southern latitude of the pack ice edge can be shown for the period 1 July to 31 October (figures 58-61). Additionally, the mean location of the ice edge has been extracted for this same period (figures 62-65).

1

Satellite imagery has provided a method of obtaining more detailed information of sea ice distribution. Barnes et al. (1976) used four years of Landsat imagery to compile statistics on nearshore ice concentration relative to the distance from the Alaskan coast for August, September and October (figures 66-68). The lines indicate distance offshore in nautical miles.



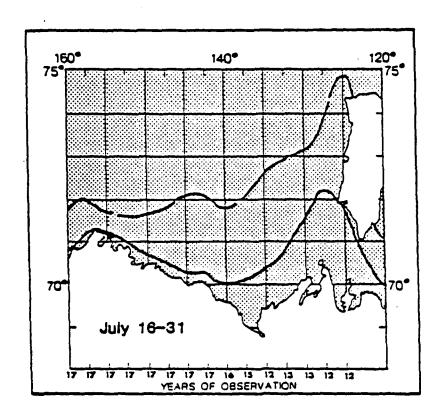
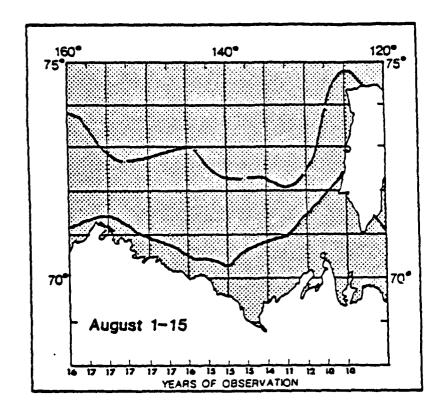


FIGURE 58 - EXTREME NORTHERN AND EXTREME SOUTHERN LATITUDE OF THE PACK ICE EDGE - JULY



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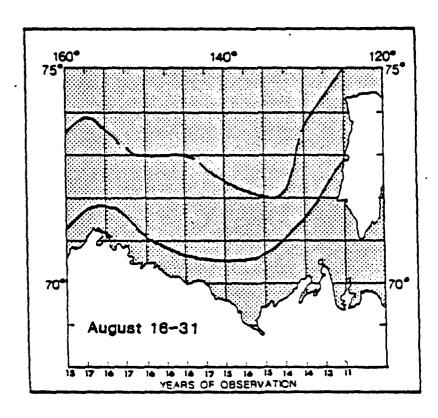
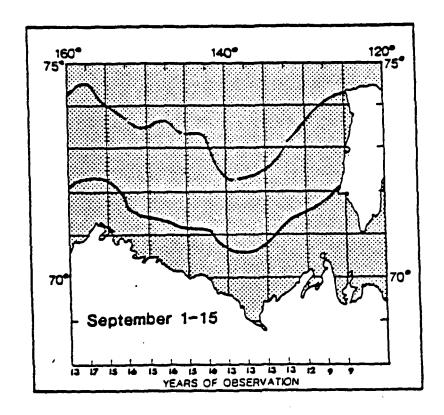


FIGURE 59 - EXTREME NORTHERN AND EXTREME SOUTHERN LATITUDE OF THE PACK ICE EDGE - AUGUST



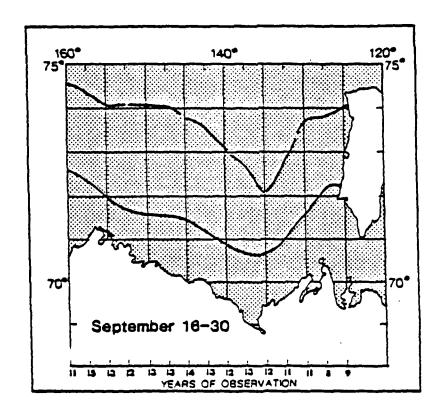
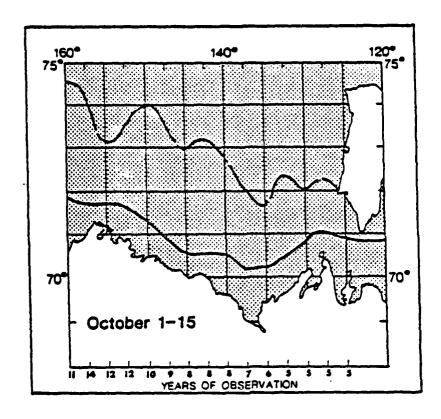


FIGURE 60 - EXTREME NORTHERN AND EXTREME SOUTHERN LATITUDE OF THE PACK ICE EDGE - SEPTEMBER



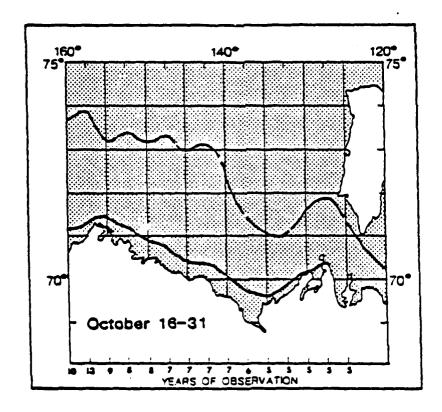
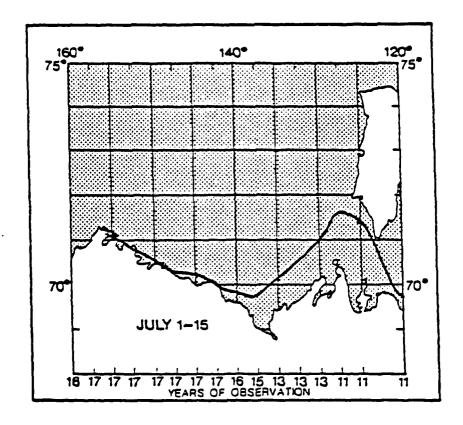


FIGURE 61 - EXTREME NORTHERN AND EXTREME SOUTHERN LATITUDE OF THE PACK ICE EDGE - OCTOBER



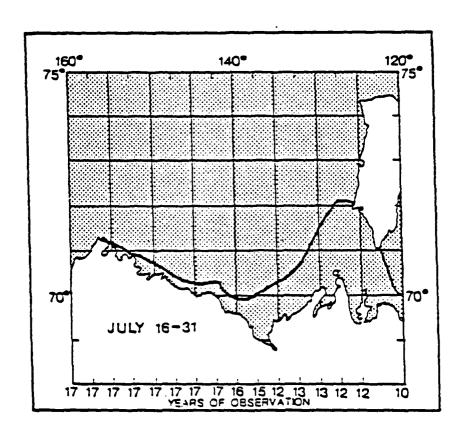
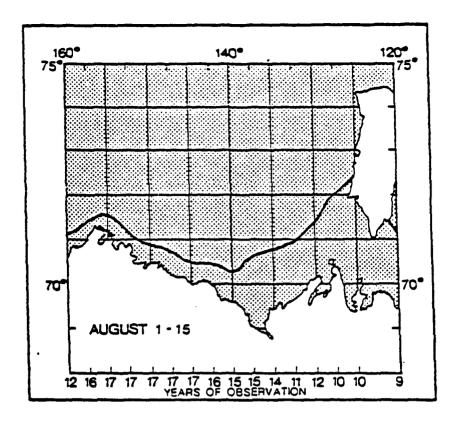


FIGURE 62 - MEAN LATITUDE OF THE PACK ICE EDGE - JULY



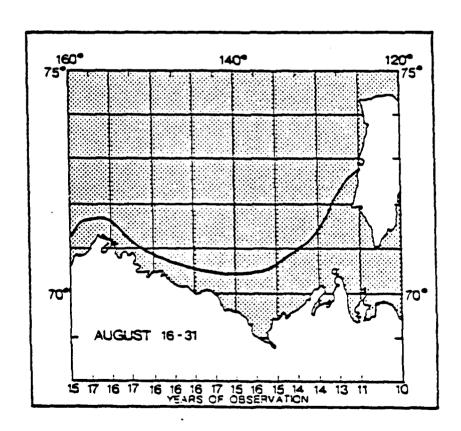
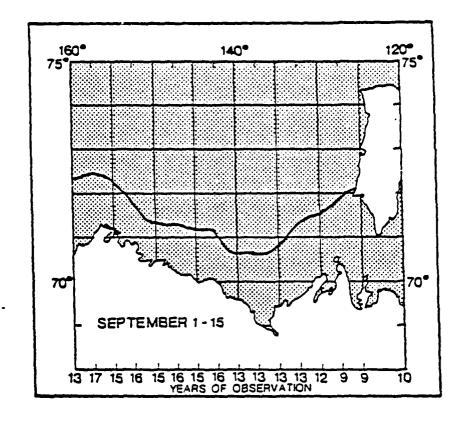


FIGURE 63 - MEAN LATITUDE OF THE PACK ICE EDGE - AUGUST C-14



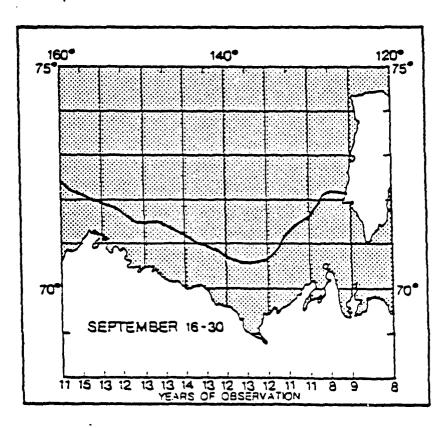
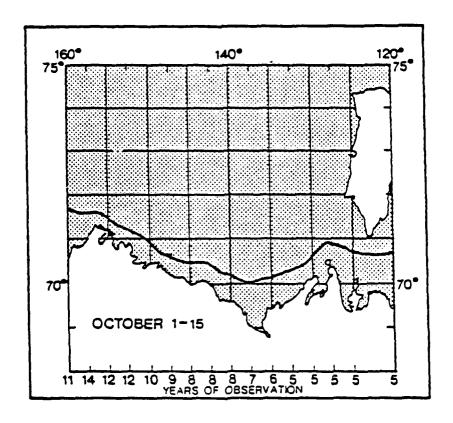


FIGURE 64 - MEAN LATITUDE OF THE PACK ICE EDGE - SEPTEMBER
C-15



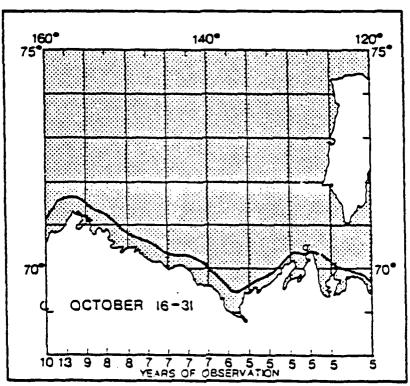


FIGURE 65 - MEAN LATITUDE OF THE PACK ICE EDGE - OCTOBER

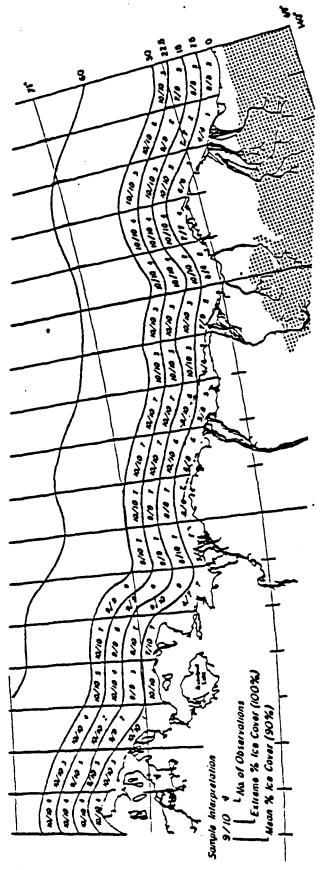


FIGURE 66 - MEAN AND EXTREME ICE COVER (%) ALONG THE BEAUFORT SEA COAST - AUGUST

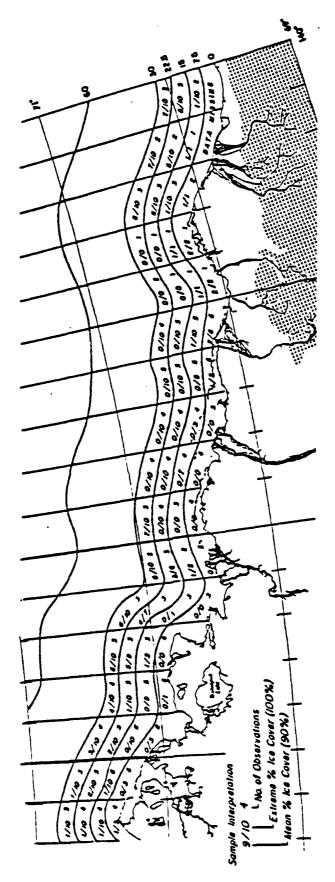


FIGURE 67 - MEAN AND EXTREME ICE COVER (2) ALONG THE BEAUFORT SEA COAST - SEPTEMBER

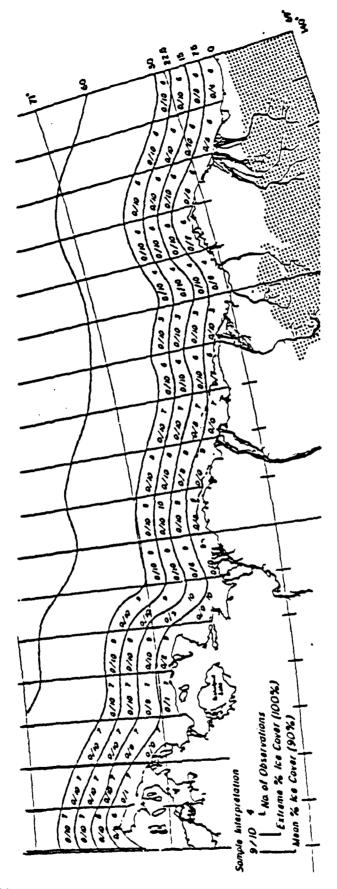


FIGURE 68 - MEAN AND EXTREME ICE COVER (%) ALONG THE BEAUFORT SEA COAST - OCTOBER

SECTION D CLIMATOLOGY

17.0 ARCTIC CHARACTERISTICS

The term Arctic is generally associated with ideas of cold, snow, ice and high latitudes. There are a number of commonly used definitions, each valid in its own right depending on the particular function for which it was developed. Definitions may use climate, conditions, or some other variable to define the polar areas. Several definitions using climate as a criterion are:

- The area lying north of the boundary fixed by the isotherm of 50°F (10°C) for the warmest month, or the isotherm of 14°F (minus 10°C) for the coldest month).
- 2. Areas having a mean annual temperature of 32°F (0°C) or below.
- 3. All areas where the sum of the average temperature in degrees centigrade of the warmest month plus one-tenth of the temperature of the coldest month is less than 9°C.

Regardless of the criteria used for the various definitions, there are certain characteristics common to the region. These are:

- 1. Short, cool summers.
- Long, cold winters.
- 3. Low annual mean temperature.
- 4. Long periods of semi-darkness.
- 5. Periods of continual daylight and darkness.
- 6. Absence of forests.
- 7. Freezing in winter of lakes, rivers, bays, and parts of the sea.
- 8. Scant precipitation.
- 9. Low absolute humidity.
- 10. Low evaporation rate.
- 11. Moist soils when thawed.
- 12. Presence of permanently frozen ground.
- 13. High windchill factor.
- 14. High latitude position.

SUNLIGHT	

18.0 SUNLIGHT

The Beaufort Sea coast receives a majority of its sunlight during the summer months. The Arctic Circle separates the area to the north which receives continuous sunlight for part or all of the summer and no sunlight during part or all of the winter (figure 69). The length of the day during the summer months is extended significantly if twilight is considered (figure 70). At latitude 70°N, twilight occurs approximately 11% of the time and therefore the yearly percentage of sunlight and twilight is 63% at the Arctic Circle latitude. These charts can be used to determine amount of darkness if a winter response to a spill is necessary.

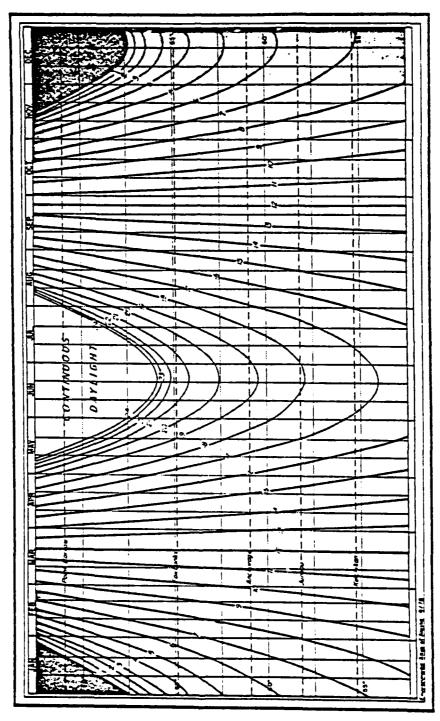


FIGURE 69 - HOURS OF CONTINUOUS SUNLIGHT ON A MONTHLY BASIS FOR NORTHERN LATITUDES (FROM HARTMAN & JOHNSON, 1978)

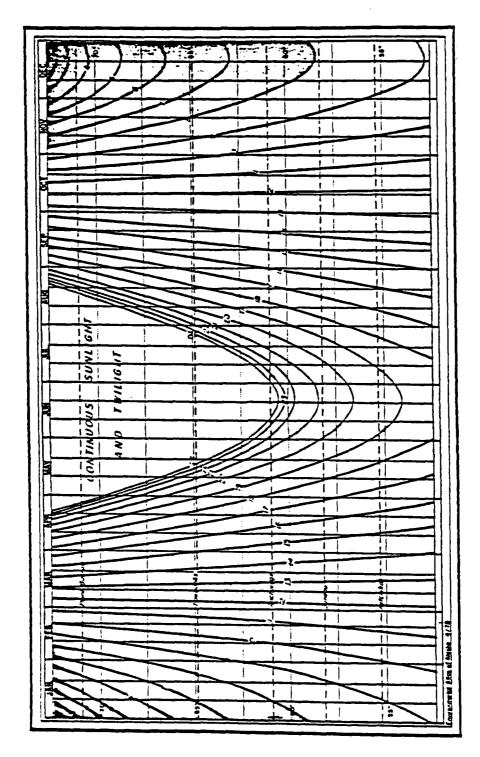


FIGURE 70 - HOURS OF CONTINCOUS SUNLIGHT AND TWILIGHT ON A MONTHLY BASIS FOR NORTHERN LATITUDES (FROM HARTMAN & JOHNSON, 1978)

TEMPERATURE	

19.0 TEMPERATURE

Temperatures in the Arctic, as one might expect, are very cold most of the year. There are large differences in temperature between the interior and coastal areas. In the interior during the summer days, temperatures often climb to the mid 60's or low 70's and occasionally rise to the high 70's or low 80's. Temperatures in the 90's are also recorded on rare occasions.

The Arctic coastal regions are characterized by relatively cool, short summers. During the summer months the temperatures normally climb to the 40's or low 50's and occasionally reach the 60's. There is almost no growing season along the coasts, and the temperatures will fall below freezing during all months of the year. At Point Barrow, Alaska, the minimum temperature fails to fall below freezing on only about 42 days a year. Over the Arctic Ocean the temperatures are very similar to those experienced along the coast; however, the summer temperatures are somewhat colder.

Winter temperatures along the Arctic coast are very cold but not nearly so cold as those observed in certain interior areas. Only on rare occasions does the temperature climb to above freezing during the winter months. The coldest readings for these coastal areas are in the -60's and -70's (degrees Fahrenheit).

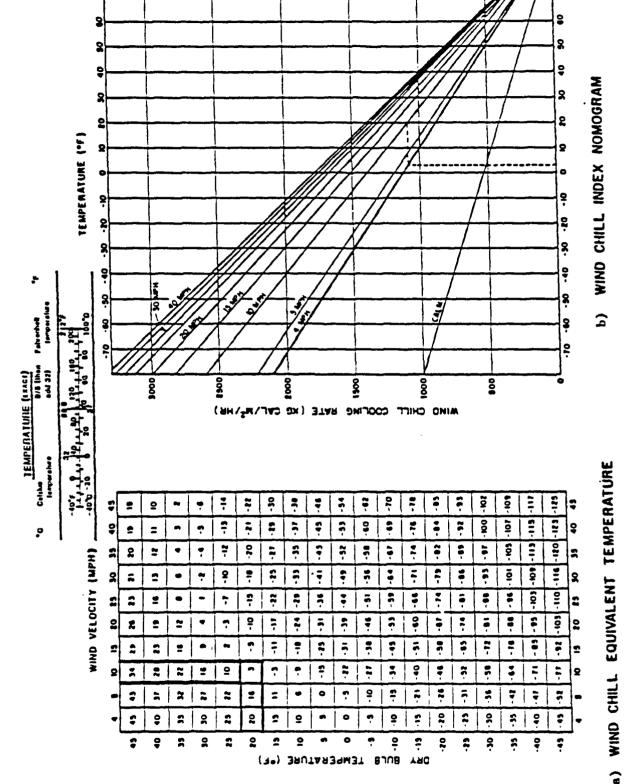
The air temperature alone is not the best indicator of how cold one feels if there is any appreciable wind. Using wind speed and temperature the concept of wind chill index or equivalent temperature can be derived (Arkin, 1971). Table 20 shows a wind chill nomogram and a table of wind chill equivalent temperatures which were prepared from the nomogram. To use the nomogram the temperature and wind speed must be known. For example, assume the wind speed is 10 mph and the air temperature is 20°F. Move along the 10 mph line until it intersects the 20°F vertical line. Then move horizontally to the 4 mph line and read the wind chill temperature vertically below this point. (Note: The 4 mph line is always used as the baseline to estimate the wind chill for any set of wind speed and temperature.) (Example: The data line shows that for a 20°F air temperature and a 10 mph wind, the equivalent temperatures is 3°F). The highlighted area of the wind chill equivalent temperature chart shows the same result as the nomogram example.

Table 21 is another example of a wind chill chart. This chart gives the approximate boundaries, for properly clothed persons, where wind chills would be of little danger, considerable danger or very great danger.

The final temperature charts (figures 71 and 72) provide the percentage probability of occurrence of free air temperature for all coastal stations (North Slope) and percentage probabilty of occurrence equivalent chill temperature developed for Barrow, Alaska (Searby, 1971) but valid for all coastal stations (North Slope). The equivalent chill temperature chart indicates the severity of the Arctic environment. Fifty percent of the time during the months of November, December, January, February and March the equivalent chill temperature poses considerable danger to properly clothed persons (table 21).

Table 20

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Table 21
WIND CHILL CHART DEFINING TEMPERATURES HAZARDOUS FOR PROPERLY CLOTHED PERSONS

			WIND	CHIL	L CH	ART					
		ro	CAL T	EMPE	RATU	REC	°F)				
WIND SPEED (MPH)	32	23	14	5	-4	-13	-22	-31	-40	-49	-51
		EQUI	VALEN	IT TE	MPER	ATURE					
CALM	32	23	14	5	-4	-13	-22	-31	-40	-49	-58
5	29	20	10	1	-9	-18	-28	-37	-47	-56	-6
10	18	7	-4	-15	-26	-37	-48	-59	-70	-81	-9:
15	13	-1	- 13	-25	-37	-49	-61	-73	-85	-97	-10
20	7	-6	-19	-32	-44	-57	-70	-83	-96	-109	-12
25	3	-10	-24	-37	-50	-64	-77	-90	-104	-117	-13
30	1	-13	-27	-41	-54	-68	-82	-97	-109	-123	-13
35	-1	-15	-29	-43	-57	-71	-85	-99	-113	-127	-14
40	-3	-17	-31	-45	-59	-74	-87	-102	-116	-131	-14
45	-3	-18	-32	-46	-61	-75	-89	-104	-118	-132	-14
50	-4	-18	-33	-47	-62	-76	-91	-105	-120	-134	-14
FOR PROPERLY CLOTH		NGER		SIDER		VE	RY G	REAT	DANG	ER	

(

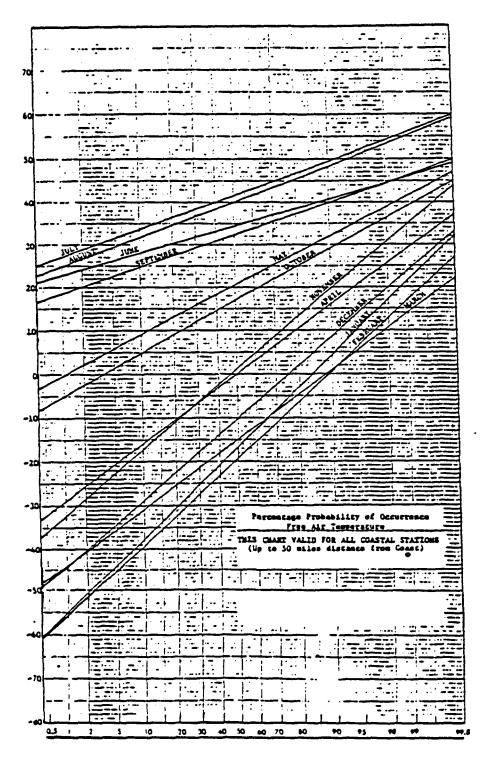


FIGURE 71 - MONTHLY GRAPHS OF THE PERCENTAGE PROBABILITY OF OCCURRENCE OF FREE AIR TEMPERATURES (FROM HARTMAN & JOHNSON, 1978)

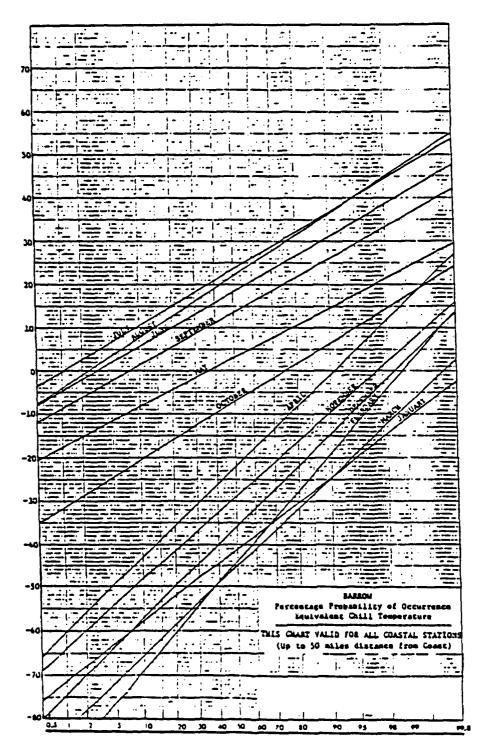


FIGURE 72 - MONTHLY GRAPHS OF THE PERCENTAGE PROBABILITY OF OCCURRENCE OF EQUIVALENT CHILL TEMPERATURE

20.0 VISIBILITY

Two conflicting factors make the subject of visibility in the polar regions very complex. Arctic air being cold and dry, is exceptionally transparent and because of this extreme ranges of visibility are possible. On the other hand, there is a lack of contrast between objects, particularly when all distinguishable objects are covered by a layer of new snow. Limitations to visibility in the Arctic are primarily blowing snow and fog.

- 1. Blowing snow. Blowing snow constitutes a more serious hazard to operations in the Arctic than in midlatitudes, because the snow is dry and fine and is easily picked up by gentle and moderate winds. Winds in excess of 8 to 12 knots may raise the snow several feet off the ground, and the blowing snow may obscure surface objects.
- 2. Fog. The two types of fog most frequently found in the polar regions are advection and radiation fog. Fog is found most frequently along the Arctic coastal areas during summers and usually lies in a belt parallel to the shore. Tables 22 and 23 provide monthly percentage frequency of occurrence of visibility and ceiling height for Barter Island and Point Barrow.

Table 22

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PERCENTAGE FREQUENCY OCCURRENCE OF CEILING HEIGHT AND VISIBILITY FOR BARTER ISLAND (FROM SEARBY & HUNTER, 1971)
(1 foot = 0.3048006 m)
(1 mile = 1.60935 Km)

	₩ 3 m.	11.1	75.3	79.8	80.4	78.5	77.1	78.5	73.0	74.6	79.4	17.9	62.1	77.9
	mi.2 1/2 mi.	.2	.1	7.	\$		\$.			.2	.2	E.	.2	
.117	1- 2 1/4 mf.	7.5	7.7	83	8.7	10.5	0.6	5.3	7.4	6.9	11.2	10.8	8.4	4.
VISIBILITY	5/8- 3/4 ml.	2.0	2.2	2.5	2.4	2.8	. 3.0	2.2	2.9	3.1	3.0	1.9	1.6	2.5
	5/16- 1/2 mi.	3.6	3.4	3.4	2.9	3.5	4.4	4.6	5.0	5.4	3.2	3.0	2.6	3.7
	3/16- 1/4 mi.	3.2	3.0	2.3	2.4	2.8	3.5	3.4	6.3	5.0	2.2	3.5	1.9	3.5
	0- 1/8 mi.	5.8		3.7	2.7	1.2	2.5	3.9	5.0	3.0	•	2.6	3.2	3.5
	>3000 ft.	13.7	17.8	80.5	75.8	33.9	48.3	9. 09	48.6	38.5	37.7	51.9	4.89	57.9
GHT	2100- 3000 ft.	5.9	4.9	5.5	5.1	6.5	5.2	5.5	5.0	6.9	7.6	9.9	7.4	6.2
CEILING HEIGH	1000- 2000 ft.	11.7	8. 0	7.3	11.3	20.1	12.3	7.1	10.3	21.9	. 31.0	25.0	14.9	15.2
2	500- 900 ft.	2.9	1.6	1.2	5.4	28.4	18.2	9.6			15.0	11.2	5.6	10.6
	400 CE.	5.8	1.1	2.5	2.4	11.1	16.0	17.5			6.9	5.3		9.6
	HINOM	6	02	6	70	6	8	ô	96	6	Ę	=	13	1 2

Table 23

E

PERCENTAGE FREQUENCY OCCURRENCE OF CEILING HEIGHT AND VISIBILITY FOR POINT BARROW

(FROM SEARBY & HUNTER, 1971)

(1 foot = 0.3048006 m) (1 mile = 1.60935 km)

		ت ا	CETLING HEIGHT	IGHT					VISIBILITY	ITY		
HING	ONTH	500- 900 ft.	1000- 2000 ft.	2100- 3000 ft. > 3000 ft	> 3000 ft	0- 1/8 mi.	3/16- 1/4 mi.	5/16- 1/2 m1.	5/8- 3/4 ml.2	1- 1/4 mi.	2 1/2 ml.	V) at.
5 2	2.7	4.8	•	3.9	79.6	1.5	2.6	2.5	1.4	5.5	6.	86.2
03	3.1	3.8	10.9	3.7	78.5	2.5	1.8	2.9	1.6	8.1		82.8
63	•.	2.9	7.9	5.4	82.9	ີ. ===	s.	1.9	1.0	4.8		91.2
04	2.6	7.8	12.4	5.7	71.5	٥.	2.0	2.4	1.1	5.0	.2	88.4
05	11.0	29.8	23.3	5.2	30.7	٠.	2.0	3.1	1.0	5.6	•	87.7
90	24.9	21.1	12:8	.3.8	37.4	2.5	4.0	4.3	1.8	7.1	.2	80.1
6	16.4	14.9	9.0	4.3	55.4	1.7	2.5	3.8	1.3	5.9		84.5
80	19.4	23.9	19.7	7.9	33.1	1.8	2.3	3.2	1.2	6.0	€,	85.2
8	13.1	31.0	30.9	6.1	18.9	٠.	9.	1.1	æ.	4.9		89.5
2	7.2	22.1	27.1	8.5	35.1	æ.	1.6	2.1	1.5	8.0	٠.	85.6
=======================================	4.5	12.0	30.0	7.4	46.1	3.1	2.0	4.4	1.4	8.7	•	80.0
12	3.8	7.6	17.0	5.2	4.99	1.7	2.5	2.8	€.	8.3	s.	83.4
!	9.2	15.1	17.3	5.2	53.2	1.5	2.0	2.9	1.2	6.6	6.	65.5
								,				

21.0 PRECIPITATION

Precipitation amounts are small varying from 3 to 7 inches (\sim 7 to 18 cm) along the Arctic coastal area and over the ice pack. The climate over the Arctic Ocean and adjoining coastal areas is as dry as some of the desert regions in the United States. Most of the annual precipitation falls as snow during the winter months. Precipitation statistics are found in tables 24 and 25.

PERCENTAGE FREQUENCY OF OCCURRENCE OF DAILY AMOUNTS OF SNOWFALL AND PRECIPITATION BARTER 1SLAND (FROM SEARBY & HUNTER, 1971)
(1 1nch = 2.54 cm)

i i	11		1	25	12	55	•	9	5				œ		~	~	
	ET/I	זל ען	mmlxs!	2.25	1.22	0.55	, 0.44	0.76	1.15	1.17	1.11	2.23	1.98	0.43	0.55	1.15	
	l_1	tonch	Comminated	0.01	-	-	-	-	90.0	0.15	0.16	0.07	0.13	0.04	-	-	
	T L	tanoi	: mumixal:	4 .08	2.53	1.44	1.22	1.51	2.09	2.19	3.40	4.91	3.62	1.50	1.17	4.91	
	Hazu	Days	.oK neak 0 ≤ nqaq	6.7	6.2	9.9	7.6	4.9	6.5	9.1	11.1	6.6	13.8	6.3	6.5	99.3	
1101			Hean Sons	0.40	0.35	0.20	0.17	0.25	0.51	0.88	1.05	0.94	0.84	07.0	n. 29	f. 28	
PRECIPITATION	PRINTILLY VALUES BELOW ARE PER- CENTACE FREQUENCY OF OCCURRENCE OF DATLY A'HOUNTS	SJIICS	85. ≤	~	4	•	-	•	-	4	•	_		•	•	-	
PREC	ARE L	7.1	ςz·-ττ·	•	•	-	-	-	~	•	•	•	^	•	-	•	
	BELOW ARE PER- Y OF OCCURRENC ATOUNTS	BELOW ARE IN INCHES	OT 90 ·	•	~	~	-	~	•	9	~	•	•	•	~	•	
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	FREQUENC OF DAILY		to.	\$	~	•	2	•	•	•	•	•	•	•	-	1	f snov.
	HONTHLY VALUES ENTAGE FREQUENC OF DAILY	CATICORIES	93877	35	32	39	35	15	*	7	333	40	32	38	17	39	nt of
	HONT	ATLC	эпой	3	46	39	40	29	34	30	32	27	24	74	3.8	34	Ž.
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		_	TTRIMONC	9	_	•	-	_	•	•	-						
	изти	Days	Nean No. Listwond	9.9	6.4	6.6	7.8	6.3	2.8	9.0	2.0	4.9	13.3	9.2	6.5	73.3	iter
	43174	PYSG	Snowfall Nean No.						7.3 2.	2.5	7.4 2.0			<u>.</u>		15.8 73.3	he vater
	d3.1V	lonch	.ok nask	6.2 35.0 6.0	3.1 15.3 6.	2.8 15.0 6.0	2.7 12.2 7.0	3.3 11.11 6.				6.2 35.8 6.4	9.2 32.1 13.3	5.5 14.9 8.2	3.8 12.9 6.5	35.8	the v
	43174	lonch	Snowfall Maximum H Snowfall Mean No.	35.0	15.3	15.0	12.2	11.11	7.3	2.5	7.4	35.8	32.1	14.9 8.	12.9		u
	AT AT	.hly lonch.	Hean Hong Snowfall Haximum H Snowfall Hean No.	35.0	15.3	15.0	12.2	3.3 11.1	7.3	2.5	7.4	35.8	32.1	14.9 8.	12.9	46.6 35.8	u
	RE PER- CURRENCE	.hly lonch.	Hean Hong Snowfall Haximus H Snowfall Snowfall	35.0	15.3	15.0	12.2	3.3 11.1	7.3	2.5	7.4	35.8	32.1	14.9 8.	12.9	46.6 35.8	u
SNOWFALL	NA ARE PER- OCCURRENCE	RE IN INCIES	2.3-2.3 2.3 — Hean Hone Showtall Haximum H Snowtall Snowtall	35.0	15.3	15.0	12.2	3.3 11.1	7.3	2.5	7.4	35.8	32.1	14.9 8.	12.9	8.66 35.8	recipitation includes
	NA ARE PER- OCCURRENCE	RE IN INCIES	3.3-2.4 2.3-2.4 2.3-2.4 3.3-2.4 3.3-2.4 4.3-2.4 Amminant Heariman Maximum Maxim	35.0	15.3	15.0	12.2	3.3 11.1	7.3	2.5	7.4	35.8	32.1	14.9 8.	12.9	8.66 35.8	cipitation includes
	NA ARE PER- OCCURRENCE	RE IN INCIES	3.5-2.4 3.5-6.4 2.6-6.4 2.6-6.4 Mean Hone Itanon Heamixah Hone Hone Hone Hone Hone Hone Hone Hone	35.0	5 3 4 3.1 15.3	15.0	12.2	3.3 11.1	7.3	2.5	7.4	35.8	32.1	A A 5.5 14.9 8.	12.9	8.66 35.8	· Precipitation includes
	NA ARE PER- OCCURRENCE	RE IN INCIES	2.5-2.4 2.5-2.4 2.5-6.4 2.6-6.4 Mean Hons Showfall Heximum P Snovfall	11 6 3 4 1 4 6.2 35.0	15 5 3 4 3.1 15.3	17 4 4 4 2.8 15.0	21 4 1 • • 2.7 12.2	16 3 1 4 4 3.3 11.1	5 3 * * 1.6 7.3	2.5	1.7 7.4	2 * * 1 * 6.2 35.8	22 16 3 1 * * * 9.2 32.1 1	17 8 2 4 4 5.5 14.9 8.	1 4 4 3.8 12.9	13 5 1 4 4 4 4 46.6 35.8	0.5 . Precipitation includes
	NA ARE PER- OCCURRENCE	RE IN INCIES	3.5-2.0 3.0-1.0 3.5-2.1 3.5-2.2 3.5-2.2 5.6-2.2 5.6-2.2 6.5-2.2 6.5-2.2 6.5-2.2 7.5-2.2 6.5-2.2 7.5-2.2 6.5-2.2 7.5-2.2 7.5-2.2	34 11 6 3 4 1 4 6.2 35.0	31 15 5 3 4 3.1 15.3	39 17 4 ^ • 2.8 15.0	32 21 4 3 4 4 2.7 12.2	45 16 3 1 4 4 3.3 11.1	23 5 3 • • 1.6 7.3	2.5	8 3 2 4 4 1.7 7.4	25 12 5 2 * * 1 * 6.2 35.8	30 22 16 3 1 * * * 9.2 32.1 1	38 17 8 2 A A 5.5 14.9 8.	41 16 4 1 4 4 3.8 12.9	5 1 4 4 4 46.6 35.8	< 0.5 . Precipitation includes
	N ARE PER- OCCURRINCE	IN INCIES	4.0-1.0 4.0-2.0 4.5-2.1 4.5-2.2 4.5-2.2 4.6-2.2 4.6-2.2 4.6-2.2 4.6-2.1 4.10-2.0	11 6 3 4 1 4 6.2 35.0	15 5 3 4 3.1 15.3	17 4 4 4 2.8 15.0	21 4 1 • • 2.7 12.2	16 3 1 4 4 3.3 11.1	5 3 * * 1.6 7.3	2.5	3 2 * * 1.7 7.4	12 5 2 4 4 1 4 6.2 35.8	22 16 3 1 * * * 9.2 32.1 1	17 8 2 4 4 5.5 14.9 8.	16 4 1 4 4 3.8 12.9	13 5 1 4 4 4 4 46.6 35.8	means < 0.5 . Precipitation includes

PERCENTAGE FREQUENCY OF OCCURRENCE OF DAILY AMOUNTS OF SNOWFALL AND PRECIPITATION
POINT BARROW
(FROM SEARBY & HUNTER, 1971)
(1 Inch = 2.54 cm)

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		. ezil	זט זל	munikak.	0.00	0.36	0.71	0.42	0.30	0.82	0.86	0.83	0.56	1.00	0.41	0.26	1.3	
			;ןסטנץ	Minimia	0.00	0.00	0.00	0.00	-	-	-	-	0.01	0.12	-	9.0	9.9	
	Inches	4 1	(dsnot	muaixel!	1.04	0.81	1.49	1.36	0.81	1.15	7.44	2.81	1.56	1.65	1.15	0.76	2.81	
		433A	"IO.	.ok. nask 0 == nqɔ٩	4.7	5.1	4.1	4.6	4.3	5.8		13.8	6.6		9.1	5.3	- 1	
¥.			nolls	Precipite	0.18	0.17	0.11	0.11	0.12	0.36	n.77	0.90 1	0.64	n.sn 12.n	0.23	9.17	4.26 87.0	
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	17	-	CATECORIES	Trace	20	64	28	1	1	2	2	7	55	7	46	8	⋝	10
	INO	NTA	TECC	Hone	25	33	28	33	15	34	ž	7	22	11	7.	=	2	aler
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		I. ATEP	· × × × × × × × × × × × × × × × × × × ×	.oW mask [[s]von2	5.1	5.3	4.8	5.4	4.4	1.6	1.0	1.4	6.3	12.7	9.1	5.8	62.0	cipitation includes the water equivalent of
	Γ.		Tuauoi	Kaximum :	11.9	7.6	15.8	15.4	12.9	6.6	0.6	6.0	7.9	21.2	19.0	6.7	21.2	2 4
	Inches		(4,000	Snowfall	2.4	2.3				~			•			- {	1	
		1	¥ţų:	Hean Hong	7	<u>~</u>	2.0	2.2	2.0	0.5	0.7	0.7	2.9	7.0	3.7	2.8	29.2	clude
		643		٤.،ه ≤							• u	он						n in
3		ENC!	Ş	7.8-2.4							• •	о қ				-		at to
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	THE BEE	QUENCY C	BELOW	7.1-2.0	-	•	-	•								4		ĺ
	V VATURE BELL	I VALUES BLLOW FREQUENCY OF C OF DAILY AMOUNT	RIES BELOW		12 3	14 5	1,4	13 5	:	•	-	•	16	28	×	17	13	0.5.
	VALUE OF PARTIES BEE	TAGE FREQUENCY C	TLCORIES BELOW	7"7-5"0	48 12 3	49 14 5	56 14 1	45 13 5	68 11	38 4	11 1	155 4	48 16	40 28	72 77	55 17	44 13	< 0.5.
	COLLEGE STATES OF STATES	CENTAGE FREQUENCY OF OF DALLY AMOUN	CATLGORIES BELOW ARE	7"I-5"0 7"U-T"0	4.0		29 56 16 1	37 45 13 5		56 38 4	88 11 1	11 25 4				,		means < 0.5.

22.0. CLIMATOLOGICAL PARAMETERS

Climatological parameters for Point Barrow and Barter Island are summarized in tables 26-29.

Table 26

CLIMATOLOGICAL SUMMARIES FOR POINT BARROW AND BARTER ISLAND (1 knot = 0.5148 m/s) (1 inch = 2.54 cm)

BARROW, ALASKA (Airport Station), 71"18"N., 158"47"W.; Elevation (ground) 22 foot.

		Air man	P.)				cipita (inche			COME)		Wind (knat						Mess	number	ol days		
		omu!		Ext	-		•	total	m.	ě							هادمین معیده ه					\lceil
Month	Daily mealmum	Delly minimum	Monthly	Record highest	Record lowest	Normal total	Maximum in 24 hrs.	Snow, altel, mean	8:00 a.m. Lecel time	2:00 p.m. Local tim	Mean speed	Prevailing direction	Meximum speed and direction	Perceit of possible sunshing	Mean aky cover	Geer	Prtly cloudy	Cloudy	Precipitation .01 fact or more	Snow, elect 1,0 inch or more	Thuaderstorms	Heavy fog
(a)				46	41		40	•1	17	17	32	12			20	27	27	27	41	17	41	20
88.	لعف	-21 6	-13.1	122	-11	0.16	0.70	2.1	1 53	62		25E		1		3	,	١,		! .		١.
• 6	لفطل	-24.3	-17 9	13:	- 36	9 13	10.36	2.2	43	61	9. 6	ENE		† 	3.3	12		10			8-	┿
	. 7.1							1.6		64	9.6	ENE		Τ	3 0	14	7	10	,		à	H
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\uc	43.7	33.0	38 4	73		0.80	0.43		74			E		} 	9.0	1 2	 	24	10	 •	-	113
×9.	34.2		10.6			0.55		2.9	92	80		E		† 	9.3	 -	 	25	10		0	1
×:.	22 0		_			0.52		6.9		144	11.3	Ē		+	8.8		 	25	16	 	- 8	+ 3
Y		• 5.3		38	-40			3.5	115	114	110.9			1.	0	4	1 3	10	1	-	0	1
Dec.	- 4.2	-16.5	-10.	34	-34	0. 29	0. 26	2.7	65	45	10.0	ENE		T	0	0	0	0	3	•	ě	1
Yesr	15.5	4.2	10.	78	-54	4. 11	1.00	27.0	79	17	10.3	E		1	-	62	31	184	69	i .	·	185

BARTER ISLAND, ALASKA (Airport Station), 70°08'N., 143°18'W.; Elevation (ground) 38 (est.

0

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Table 27

PERCENTAGE FREQUENCY OF OCCURRENCE OF WEATHER CONDITIONS FOR POINT BARROW, BARTER ISLAND AND UMIAT (AN INLAND STATION)
(FROM SEARBY & HUNTER, 1971)

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	arreer is.	6.9	8.0	9.7	11.8	25.1	26.6	25.3	31.5	26.6	13.5	9.8	0.0	
9 =	382#	35.7	37.9	34.7						17.5	19.4	47.7	50.6	
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Table 28

C

MONTHLY VALUES OF PERCENTAGE FREQUENCY OF OCCURRENCE OF SNOW DEPTH, VISIBILITY AND SKY COVER POINT BARROW (FROM SEARBY & HUNTER, 1971)

(1 inch = 2.54 cm)

				_												
	3		TU LEU VAE 2K	4.4	4.7	5.0	9.5	9.4	6.3	8.2	9.3	9.3	8.6	7.6	5.7	7.1
	1	rs)	οτ	33.2	34.9	34.5	42.1	75.2	0.89	62.2	82.2	85.0	73.6	63.0	45.4	58.3
VER	OF SKY	LOW ARE	6-8	5.3	6.2	8.7	7.8	6.2	10.0	14.3	8.5	6 .8	8.3	7.4	5.9	9.0
SKY COVER	CE FRI	OF TOTAL	۷-9	4.3	4.8	S.8	4.7	2.8	4.7	9.6	2.8	2.1	4.1	4.8	4.3	4.2
	PERCENTAGE FREQUENCY OF OCCURRENCE OF SKY COVER (BASED ON HRLY OBSERVATIONS)	CATECORIES BELOW ARE TENTHS OF TOTAL SKY COVER	5-7	4.0	3.7	5.2	4.2	2.1	4.1	9.4	1.8	1.7	3.0	3.8	4.0	3.5
	2005	2 1 2	€-0	53.2	\$0.4	45.8	41.2	13.7	13.2	13.3	4.7	4.4	11.0	21.0	40.4	26.0
	-	sned l l l l	Total o	.7.9	8.7	3.7	6.3	9.9	12.7	7.6	8.4	4.1	6.1	10.8	7.8	7.7
7	EQUENCE NCE U Vaby d by:	1	Precip-	.2	ŗ.	:	9.	ŗ.	7.		-		1.0	ż.		
VISIBILITY	CENTAGE FREQU OF OCCURRENCE IT Obans/w V	von2 3rud	Bloving and/or	4.1	9.9	2.5	2.7	.3					1.7	6.8	5.2	2.5
17	PERCENTAGE FREQUENCY OF OCCURRENCE Hourly Obans/v Vsby	92 8 H	Smoke and/or													•
	V		303	3.5	1.8	1.0	3.0	6.1	12.4	7.6	8.3	3.7	3.4	3.6	2.5	6.9
	T., ATEP	J = 0.	Mean No Snowfal	31.0	28.0	31.0	30.0	31.0	14.8	0.0	0.0	3.5	27.2	30.n	31.0	6.7 257.5
	чз		Hean Sn	9.8	12.1	12.5	13.7	11.4	2.2	-	j	0.2	3.2	9.9	8.7	6.1
	<u>ယ</u> 1		87-18											-		1
	10		9E-SZ													
	(BASED ON DAS INCHES		77-57	2	20	20	20	30	7					4	10	18
FPT	= -		27-2	2	48	20	20	55	=======================================				13	43	57	32
SNOW DEPTH	S BILLOW ARE PERCENTA- OCCURRENCE (BASED ON DBSEBVATIONS		9-7	2	7			14	•				24	77	33	12
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Table 29

e

MONTHLY VALUES OF PERCENTAGE FREQUENCY OF OCCURRENCE OF SNOW DEPTH, VISIBILITY AND SKY COVER

BARTER 1SLAND
(FROM SEARBY & HUNTER, 1971)

(1 1nch = 2.54 cm)

				COVET	tn Tent	5.3	5.2	5.3	6.1	6.3	G. E	7.9	8.5	4.8	8.2	6.9	5.8	7.0
٠	10	VER	(BASED ON URLY OBSERVATIONS)		UT	37.8	36.4	34.3	42.6	70.5	4.09	55.9	66.2	6.89	68.9	56.1	43.0	53.6
SKY COVER	FREQUENCY OF	SKY CO	BSERVA	JU ARE	6-8	7.6	7.4	6.6	10.5	8.2	13.4	15.3	14.0	10.3	8.1	1.5	7.3	10.0
	E FRE	OCCURRENCE OF SKY COVER	IRLY O	CATECORIES BELOW ARE TENTHS OF TOTAL SKY COVER	L-9	9.9	5.7	7.3	6. 6	4.0	6.2	7.3	5.7	4.8	4.2	4.6	5.8	5.7
	PERCENTAGE	URRENC	NO Q	CATECORIE Tentiis of Cover	5-7	5.6	6.2	6.3	5.6	3.2	5.0	4.1	4.5	3.9	3.8	4.4	5.4	5.0
	PER	200	(BASE	CATEC TENTII COVER	€-0	42.4	44.3	42.2	34.8	14.1	15.1	15.5	9.6	12.2	14.8	27.4	38.4	25.7
		1			Total c	14.6	16.9	11.9	10.4	10.3	13.5	16.2	19.4	16.4	9.1	11.0	9.3	13.2
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רונו	FREQUENCY	RENCE	V Vene		-qisarq molisil	Œ	÷.		ς.	ς.	4.			δ.	1.4		æ	9.
VISIBILITY		OF OCCURRENCE	uequ .	- nous	Blowing To\bac	12.2	14.3	9.0	5.8	€.				€0	7.6	6.5	6.4	4.8
	PERCENTAGE	0F.	Hourly		Smoke and/or		0.	Ċ.							•			0.
		! !	 		203	1.6	2.0	2.7	3.7	9.0	13.0	16.1	19.1	14.7	5.1	æ.	2.1	7.8
		43	.,T MTI	DEYS	Mean No Snowfal	31.0	28.0	31.0	30.0	31.0	15.4	0.5	0.9	6. 6	25.4	29.9	31.0	1 10.0 260.7
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SNOW DEPTH	MONTHLY VALI'ES BELOW ARE PERCENTA	FREQUENCY OF OCCURRENCE (BASED ON	OBSERVATIONS)		9-7	6	2 1	1	9	14 2	7			7	34 2	7 92	16 4	10 2
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23.0. CASE HISTORIES

The following fourteen case histories provide descriptive information for severe storm surges between 1954 and 1977.

STORM PROFILE: 1 REGIONS: West Arctic COMMUNITIES: Barrow

INCLUSIVE DATES: 17/9/1954 TO 18/9/1954
DAMAGE, Barrow: Water washed over the beach into camp, a helium tank from

Barrow village was moved almost to the point.

MAXIMUM SURGE: 9 to 10 ft in Barrow

DATA SOURCES: 8.0

STORM PROFILE: 2 REGIONS: West Arctic

COMMUNITIES: Barrow, Wainwright

INCLUSIVE DATES: 3/10/1954 TO 5/10/1954

DAMAGE: Minor damage. Wainwright had beach erosion, goods were moved from

store, and a scow was washed 4 miles to east and buried in sand.

MAXIMUM SURGE: 9.5 ft

COMMENTS: Only known reference is by Hume. The 1954 storm was the strongest

prior to the October 1963 storm.

DATA SOURCES: 8.0, 5.0

STORM PROFILE: 3 REGIONS: East Arctic COMMUNITIES: Barter Island

INCLUSIVE DATES: 19/9/1957 TO 21/9/1957

DAMAGE: Road damage, 4,400 barrels of fuel washed away. Part of runway

undermined. Navy LST ran aground.

MAXIMUM SURGE: 6-12 ft DATA SOURCES: 8.0

STORM PROFILE: 4

REGIONS: West Arctic, Colville, East Arctic

CONMUNITIES: Barrow, Barter Island, Point Lay, Wainwright

INCLUSIVE DATES: 3/10/1963 TO 3/10/63

DAMAGE: \$3 million to Barrow. 15 homes plus 15 other buildings and contents, 4 airplanes, freshwater supply contaminated with sea water, electrical generating plant received \$100,000 damage. Sediment transport was the equivalent of 20 years normal transport. Water depths 2 ft at ARL Laboratory. 3.5 ft in other areas. Wainwright 50% flooded, 4 ft water to top of bluff. MAXIMUM SURGE: Barrow 11-12 ft; Barter Is. 5.5 ft; Pt. Lay 9 ft; Wainwright

CONTIENTS: A deepening low pressure center 995 to 980 mb moved from 74N 163W to 74N 12OW at 25-30 knots in 24 hours. Surge developed in west to northwest flow, south of storm.

DATA SOURCES: 8.0, 5.0, 3.0

STORM PROFILE: 5

REGIONS: Western Arctic

CONTIUNITIES: Barrow, Pt. Barrow

INCLUSIVE DATES: 21/9/1968 TO 24/9/1968

DAMAGE: \$50,000. Road between Barrow and city dump (3 mi) eroded severely

and a bridge damaged. MAXIMUM SURGE: 8.5 ft

COMMIENTS: A storm moving west to east at 15 kts, 200 miles north of Barrow,

brought 25 ft waves offshore DATA SOURCES: 1.0, 3.0, 5.0

STORM PROFILE: 6

REGIONS: West Arctic, East Arctic

COMMUNITIES: Oliktok, Prudhoe Bay, Barrow INCLUSIVE DATES: 13/9/1970 TO 13/9/1970

DAMAGE: Oliktok lost several hundred ft of runway, driftwood lines indicated

a surge of approximately 3 meters.

MAXIMUM SURGE: 3 m at Oliktok; 3 m approx. at Prudhoe; 2.4 and 3 m at Herschel COMMENTS: No weather reports except Barrow. Winds from reference: Oliktok 80 km/hr (50 mph); Deadhorse 46 km/hr; Cape Haklett 130 km/hr (estimated).

DATA SOURCES: 18.0

STORM PROFILE: 7 REGIONS: East Arctic COMMUNITIES: Prudhoe Bay

INCLUSIVE DATES: 28/11/1970 TO 28/11/1970

COMMENTS: Storm center causing flooding 27/11/70 at Kotzebue curved near

Barrow to move east at 25 kt to 72N, 45W at 28/1200Z.

DATA SOURCES: 15.0

STORM PROFILE: 8 REGIONS: East Arctic

COMMUNITIES:

INCLUSIVE DATES: 2/9/1972 TO 2/9/1972

MAXIMUM SURGE: 3.8 ft

COMMENTS: A 996 mb low pressure center moved from west to east 200 nmi north of the coast. Surge occurred in northwesterly flow southwest of the storm. DATA SOURCES: 6.0

STORM PROFILE: 9 REGIONS: East Arctic

COMMUNITIES:

INCLUSIVE DATES: 27/9/1972 TO 27/9/1972

COMMENTS: Ship at MSQ 268/13 reported 14 kts W with 5 ft waves, 5 sec. period. Same ship next at MSQ 268/12 at 2100Z 9-28-72 reports wind 15 kts W.

with 15 ft waves 6-7 sec. period.

DATA SOURCES: 15.0

STORM PROFILE: 10 REGIONS: West Arctic COMMUNITIES: Point Lay

INCLUSIVE DATES: 9/10/1972 TO 10/10/1972

COMMENTS: A 980 mb low pressure system moved from 62N 177W to 74N 153W at 25

knots in 36 hours. Surge occurred with southwest winds in the southeast

quadrant of the storm. DATA SOURCES: 4.0

STORM PROFILE: 11 REGIONS: East Arctic

INCLUSIVE DATES: 5/1/1974 TO 7/1/1974

COMMENTS: Maximum winds on record at Barter Island occurred with an estimated 990 mb low pressure center moving west to east 300 nmi north of the Beaufort

Sea coast.

(

DATA SOURCES: 6.0

STORM PROFILE: 12

REGIONS: West Arctic, Colville, East Arctic INCLUSIVE DATES: 26/8/1975

DAMAGE: Unknown

MAXIMUM SURGE: 9.5 ft

COMMENTS: The 1975 Prudhoe Bay Sealift fleet was stopped for several days. One barge in the fleet went aground. Driftwood line indicated surge height of 9.5 to 10.0 ft in places. Surge heights highest in sector 1 determined by driftwood lines (Reimnitz and Haurer).

DATA SOURCES: 3.0, 4.0

STORM PROFILE: 13 REGIONS: West Arctic COMMUNITIES: Icy Cape

INCLUSIVE DATES: 26/8/1975 TO 27/8/1975 DAMAGE: High water and flooding at Icy Cape

MAXIMUM SURGE: Unknown

COMMENTS: The remains of tropical storm Rita moved north/northeast at 30-35

kts from 67N, 174W to 77N 149W in 12 hours.

DATA SOURCES: 4.0

STORM PROFILE: 14 REGIONS: West Arctic

COMMUNITIES: Barrow, Barrow Gas Wells INCLUSIVE DATES: 29/12/1977 To 30/12/1977

DAMAGE: On the morning of December 30, rising water lifted the pack ice at Barrow and wind drove it as much as 30 yds inland. Barrow gas well runway partially flooded with 6 to 18 in of water rising through a crack in the ice.

MAXIMUM SURGE: 3.5 ft

COMENTS: A storm moved north at 40-50 knots from the Aleutians thru Bering Strait to northwest of Barrow. Southwesterly winds along the Chuckchi coast

persisted 12 hours. DATA SOURCES: 5.0, 3.0

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